

NOAA Technical Report NOS 125 NGS 42



Project REDEAM: Models for Historical Horizontal Deformation

**Richard A. Snay
Michael W. Cline
Edward L. Timmerman**

**Rockville, MD
September 1987**

**U.S. DEPARTMENT OF COMMERCE
National Oceanic and Atmospheric Administration
National Ocean Service
Charting and Geodetic Services**

NOAA TECHNICAL PUBLICATIONS

National Ocean Service/National Geodetic Survey Subseries

The National Geodetic Survey (NGS), Office of Charting and Geodetic Services, the National Ocean Service (NOS), NOAA, establishes and maintains the basic national horizontal, vertical, and gravity networks of geodetic control, and provides Government-wide leadership in the improvement of geodetic surveying methods and instrumentation, coordinates operations to assure network development, and provides specifications and criteria for survey operations by Federal, State, and other agencies.

NGS engages in research and development for the improvement of knowledge of the figure of the Earth and its gravity field, and has the responsibility to procure geodetic data from all sources, process these data, and make them generally available to users through a central data base.

NOAA geodetic publications and relevant geodetic publications of the former U.S. Coast and Geodetic Survey are sold in paper form by the National Geodetic Information Center. To obtain a price list or to place an order, contact:

National Geodetic Information Center (N/CG17x2)
Charting and Geodetic Services
National Ocean Service
National Oceanic and Atmospheric Administration
Rockville, MD 20852

When placing an order, make check or money order payable to: NOAA, National Geodetic Survey. Do not send cash or stamps. Publications can also be charged to Visa, Master Card, or prepaid Government Printing Office Deposit Account. Telephone orders are accepted (area code 301 443-8316).

Publications can also be purchased over the counter at the National Geodetic Information Center, 11400 Rockville Pike, Room 14, Rockville, MD. (Do not send correspondence to this address.)

An excellent reference source for all Government publications is the National Depository Library Program, a network of about 1,400 designated libraries. Requests for borrowing Depository Library material may be made through your local library. A free listing of libraries in this system is available from the Marketing Office (mail stop MK), U.S. Government Printing Office, Washington, DC 20402 (area code 202 275-3634).

NOAA Technical Report NOS 125 NGS 42



Project REDEAM: Models for Historical Horizontal Deformation

**Richard A. Snay
Michael W. Cline
Edward L. Timmerman
National Geodetic Survey
Rockville, MD
September 1987**

**U.S. DEPARTMENT OF COMMERCE
Bruce Smart, Acting Secretary of Commerce
National Oceanic and Atmospheric Administration
Anthony J. Callo, Under Secretary
National Ocean Service
Paul M. Wolff, Assistant Administrator
Charting and Geodetic Services
R. Adm. Wesley V. Hull, Director**

For sale by the National Geodetic Information Center, NOAA, Rockville, MD 20852

CONTENTS

Abstract.....	1
1. Introduction.....	2
2. Geodetic data.....	5
3. Mathematical model.....	9
4. Parameter estimation.....	13
5. Model implementation.....	15
6. Model evaluation.....	16
7. Summary.....	20
Appendix A. Dislocation equations.....	21
Appendix B. Data distributions.....	27
Appendix C. Parameter values.....	36
Appendix D. Secular strain rates.....	53
Appendix E. Anchorage region.....	57
Appendix F. Hilo region.....	65
Acknowledgments.....	73
References.....	74

Project REDEAM: Models for
Historical Horizontal Deformation in
Tectonically Active Regions of the United States

Richard A. Snay
Michael W. Cline
Edward L. Timmerman

National Geodetic Survey
Charting and Geodetic Services
National Ocean Service, NOAA
Rockville, Maryland 20852

ABSTRACT. Models for historical horizontal crustal deformation have been derived for a total of 19 tectonically active regions, 16 of which combine to cover all of California, with one model each representing parts of Nevada, Alaska, and Hawaii. Model parameters were estimated from geodetic data (direction, azimuth, and distance observations). These data span the 1850-1980 interval; however, over 90 percent postdate 1930. The models were developed in support of the redefinition of the North American horizontal reference system. In particular, the models were used to update all appropriate geodetic observations to the values that would be obtained if the observations were remeasured on December 31, 1983. These updated observations were then used to estimate, for this date, the positional coordinates (latitude and longitude) of the geodetic marks that comprise the horizontal reference network.

The models address both the secular and episodic components of motion. For secular motion, each modeled region is partitioned into a mosaic of districts that are individually allowed to translate, rotate, and deform homogeneously as a linear function of time. Episodic movement corresponds to displacements associated with large earthquakes ($M \geq 6$), and is modeled in accordance with elastic dislocation theory. The preponderance of triangulation data for the 1930-80 interval well define the time-averaged pattern for secular shear-strain rates. Much of the pre-1970 distance data, however, are contaminated by systematic errors on the order of several parts in 10^6 . As such, an insufficient time span of accurate distances exists to well constrain the scale dependent aspects of the secular deformation which are characterized by strain rates of a few parts in 10^7 per year. Also the rotational component of the secular deformation is hidden within the 0.1-microradian-per-year noise level obtained with existing azimuth data. Episodic motion is well resolved for those earthquakes with $M > 7$; earthquakes with $6 < M < 7$ are modeled with only mixed success because the corresponding ground movements are near the threshold of pre-1970 geodetic sensitivity.

1. INTRODUCTION

The title REDEAM (Regional Deformation of the Earth Models) identifies the project to model horizontal deformation for various tectonically active regions in the United States. Individual models were developed for 19 mutually disjoint geographic regions. Sixteen of these regions cover, in combination, the State of California (fig. 1.1). The three other regions are located in Nevada (fig. 1.1), Alaska (fig. 1.2), and Hawaii (fig. 1.3).

The REDEAM models were generated in support of the North American Datum (NAD) project, an international project to redefine the geodetic reference system used by map makers, engineers, land surveyors, and others for horizontal positioning in North America.

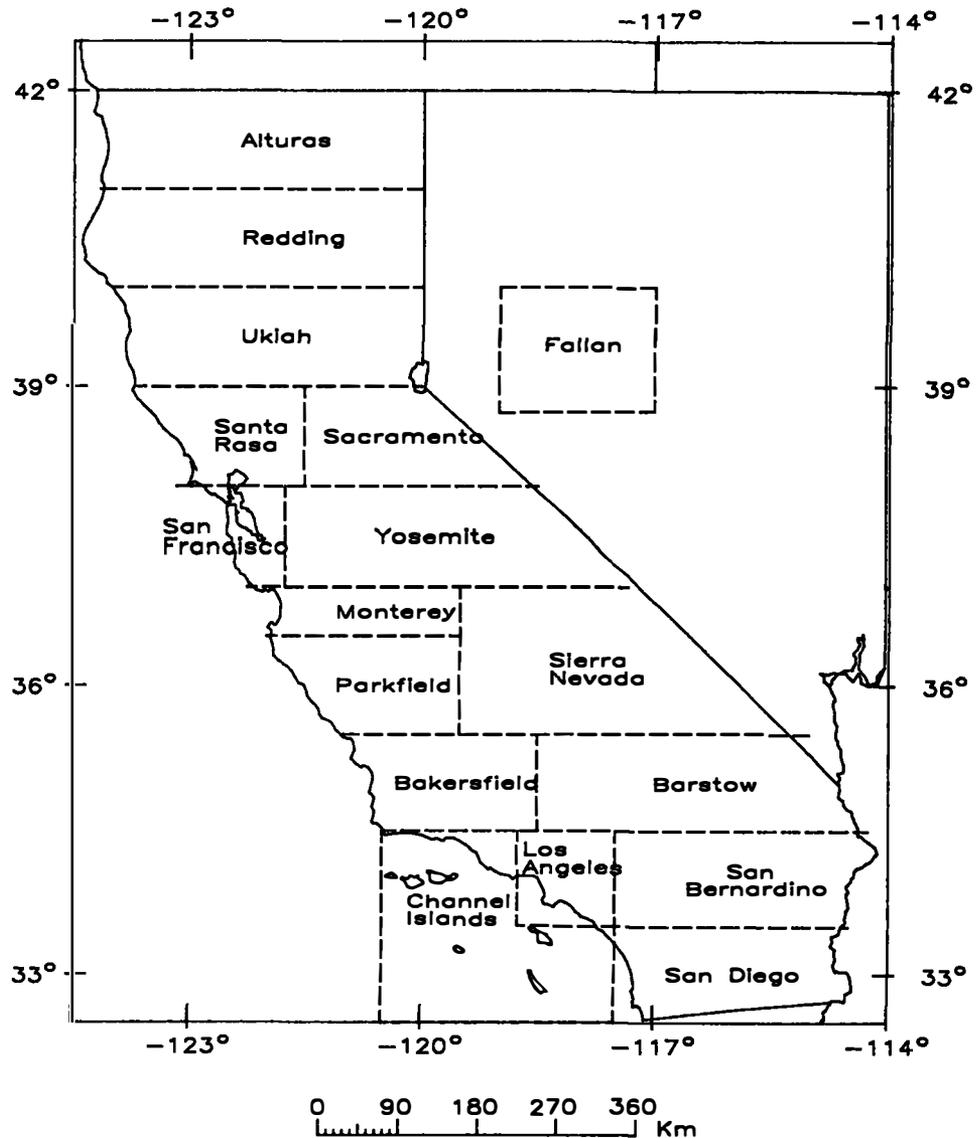


Figure 1.1.--A model for historical crustal deformation was developed for each of 19 regions. Seventeen of these regions are pictured above. Other regions are located in Alaska and Hawaii.

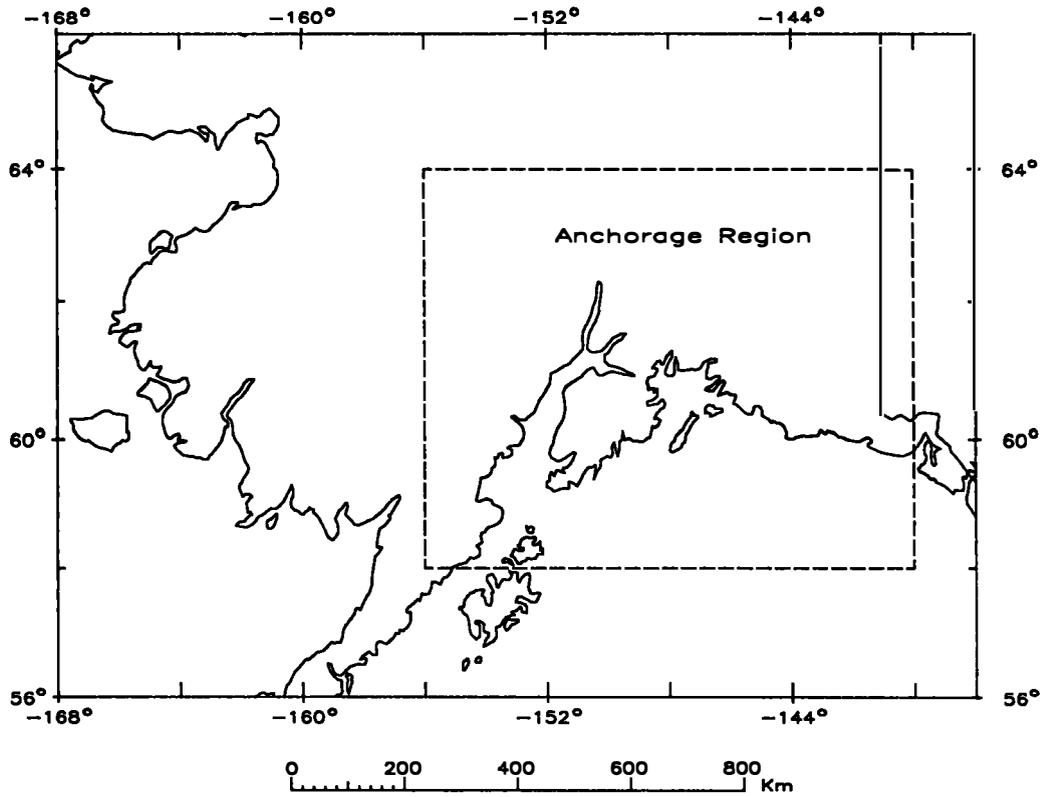


Figure 1.2.--The crustal deformation model for Alaska's Anchorage region characterizes horizontal displacements associated with the 1964 Prince William's Sound earthquake.

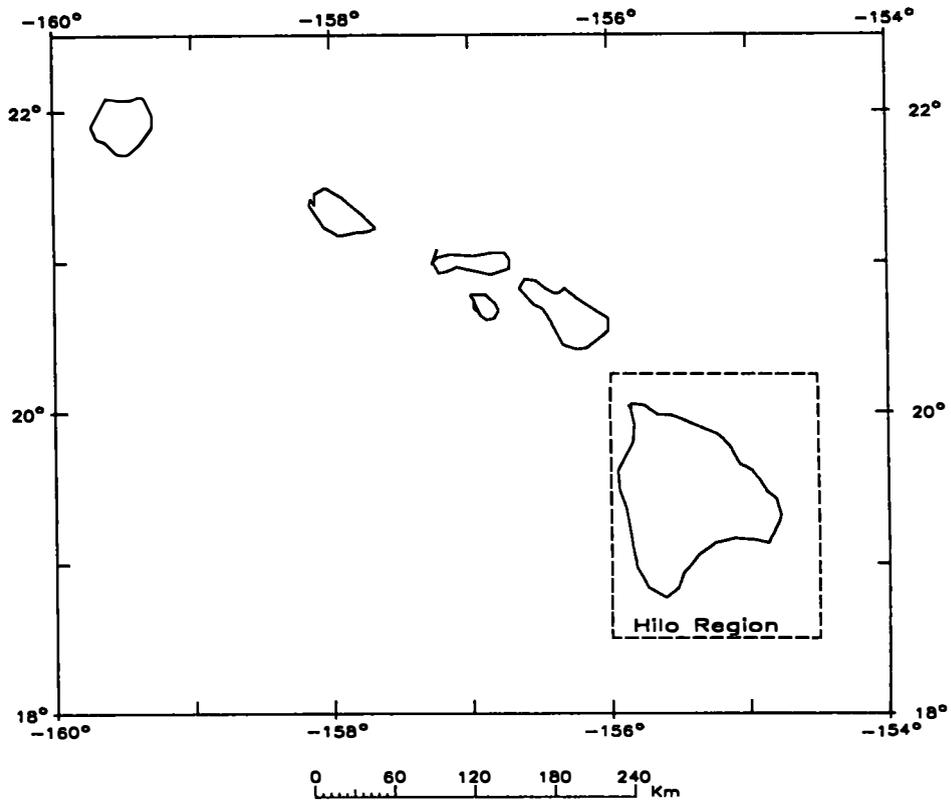


Figure 1.3.--The crustal deformation model for Hawaii's Hilo region addresses horizontal motion associated with volcanic and seismic activity, especially the 1975 Kalpana earthquake (M=7.1).

For the NAD project, all appropriate geodetic observations were entered into a simultaneous solution to estimate positional coordinates (latitude and longitude) for the several hundred thousand monumented stations that comprise the North American horizontal reference network. Prior to entry into the solution, these geodetic data were reduced, using REDEAM models, to a common date (December 31, 1983) to account for temporal variations of the positional coordinates. That is, for each geodetic observation, the REDEAM models served to estimate the value that would be obtained if the observation were remeasured on December 31, 1983. The newly derived NAD positional coordinates thus correspond in time to this date.

Since 1807 the National Geodetic Survey (NGS) and its predecessor agencies have determined latitudes and longitudes of more than 200,000 monumented stations covering the United States at spacings from 0.01 to 100 km. These monuments or marks are 'positioned' through a network of geodetic observations that include directions, taped and electronically measured distances, and astronomically determined azimuths. Currently, geodetic observations for positioning are also obtained by techniques that employ artificial satellites and extragalactic radio sources such as quasars. Today a 10-km length can be measured to a precision of a few millimeters, and a transcontinental distance to within a few centimeters. As instrument accuracies continue to improve, the need to improve the geodetic reference system becomes greater. Anticipating this need, NGS and other national geodetic agencies representing the North American countries and Denmark initiated a joint project to recompute the positions of all marks in North America and Greenland. Such a recomputation was last accomplished for the United States in 1927. The recent effort was characterized by greater rigor as advances in computer technology permitted both the use of more sophisticated mathematical techniques and of more up-to-date scientific theories of the Earth and its environment.

Contributing to the rigor of the new computations, the REDEAM models describe the time variability of station coordinates attributed to horizontal crustal motion. According to plate tectonic theory, the latitudes and longitudes of monumented stations continually change. The rates of these motions have been estimated from geologic and seismic data by using models that assume that the Earth's surface consists of several rigid plates each rotating at a constant rate about a specific pole (Minster and Jordan 1978). Although these models are acceptable on a global scale for motions averaged over millions of years, significant regional deviations develop when the motions are considered over a time period of decades. In particular, friction between adjacent plates retards relative plate motion and causes a gradual bending of the Earth's crust over a zone hundreds of kilometers in width. This regional bending is occasionally interrupted by the sudden displacements associated with earthquakes as elastic crustal elements rebound from their distorted states. The REDEAM models address both this slow regional bending and the rapid coseismic displacements.

The purpose of this report is to document the development and implementation of the REDEAM models. Chapter 2 describes the geodetic data used to estimate model parameters. Chapter 3 describes the mathematical model, and Chapter 4, the relevant software. Chapter 5 specifies the algorithm used for applying the models to update existing geodetic observations. Chapter 6 contains an evaluation of the models, and Chapter 7 summarizes the report. So as not to interrupt the flow of the presentation, this report also contains six appendices: Appendix A describes the application of dislocation theory to the characterization of earthquake displacements, appendix B shows the spatial and temporal distributions of the geodetic data, appendix C tabulates values for model parameters, appendix D tabulates secular rotation rates and strain rates derived from model parameters, appendix E describes the model for Alaska's Anchorage region (fig. 1.2), and appendix F describes the model for Hawaii's Hilo region (fig. 1.3). Appendices E and F are included because the Anchorage and Hilo models differ in form from each other and from the other 17 regional models.

2. GEODETIC DATA

Parameters for REDEAM models were estimated from geodetic data (directions, distances, and azimuths) contained in the archives maintained by NGS. This data base incorporates contributions from various Federal, state, and local organizations. The archives include numerous geodetic measurements in California which were performed explicitly to measure crustal motion. These crustal motion measurements include those performed by NGS and its predecessor agencies following most of the major earthquakes in the United States, including the San Francisco earthquake of 1906. These agencies have also repeatedly surveyed several geographic areas to monitor aseismic strain rate (fig. 2.1) and secular fault slip (fig. 2.2). The archived crustal motion measurements also include the regularly repeated line length determinations performed by the California Department of Water Resources (1968) from 1959 to 1969 and the California Division of Mines and Geology from 1969 to 1979 (Bennett 1980) and the U.S. Geological

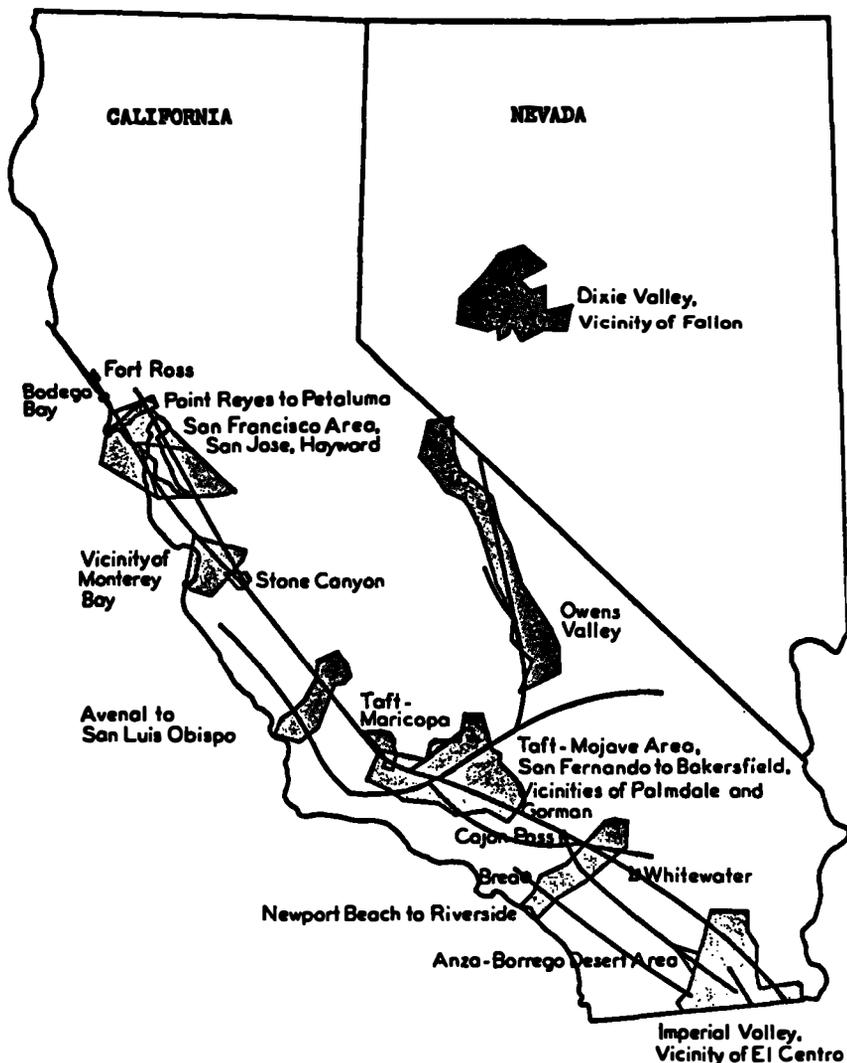


Figure 2.1-- Meade (1971) published this figure identifying several areas that were being monitored before and during the 1960s using direction or triangulation observations. In the early 1970s, the more precise electronic-distance-measuring technique became sufficiently operational for monitoring deformation over large areas.

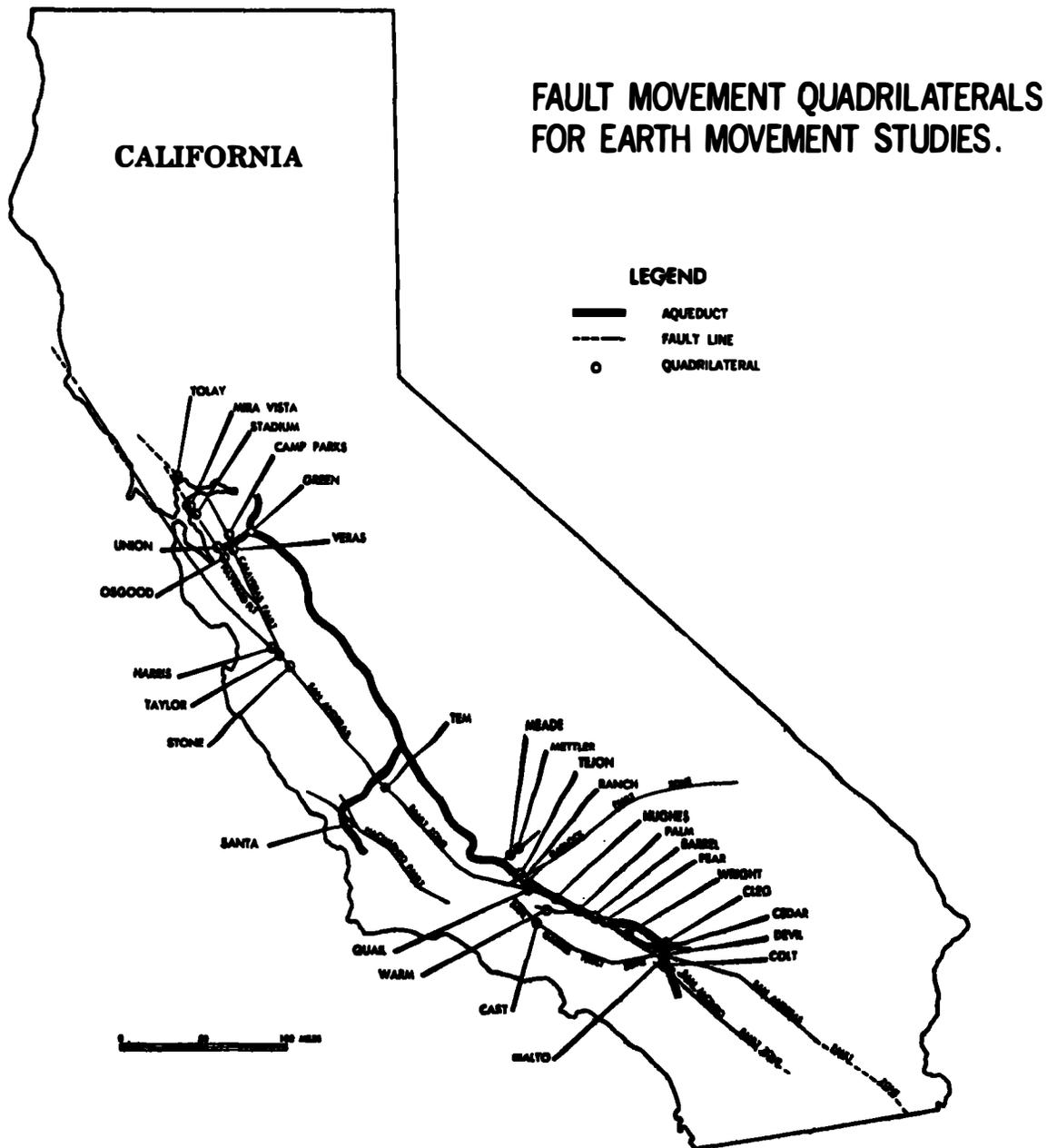


Figure. 2.2.--In the 1960s and 1970s, Federal agencies repeatedly surveyed 30 small geodetic networks to determine fault-slip rates in the vicinity of the California aqueduct. Most of these networks contained six to eight stations located within a kilometer of each other and with half of the stations to either side of the straddled fault (from Meade 1971).

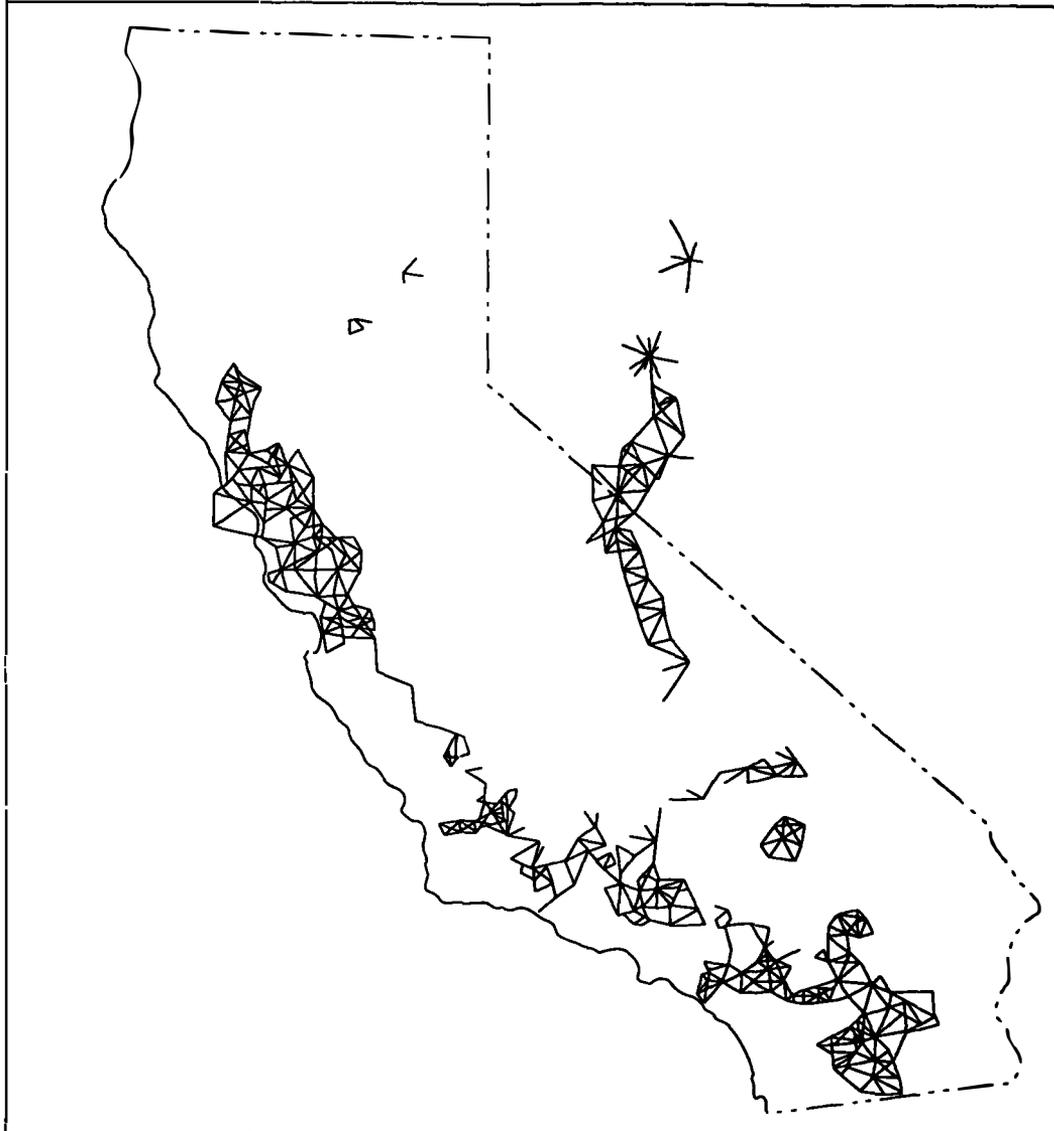


Figure 2.3.--Since the early 1970s, federal and state agencies have monitored crustal deformation using electronic-distance-measuring instrumentation. The figure identifies frequently measured lines located in California and Nevada.

Survey (USGS) from 1970 to the present (Savage 1983). (See fig. 2.3.) It is important to note, however, that most of the geodetic data used for project REDEAM were observed not to measure crustal motion but simply to position horizontal reference marks that comprise the geodetic reference network. The crustal motion information contained in this majority of the data results largely from past demands for additional reference marks whereby previously established marks were resurveyed to position the newer marks.

California's first geodetic data date back to the time of statehood, 1850. Most nineteenth century surveys, however, are concentrated along the coast as they were performed to aid navigation. California's interior network remained sparse until the introduction of Bilby towers around 1930. Because these 20- to 40-m tall observation platforms are transportable and can be erected or dismantled in less than a day, they provided an economical means for seeing over trees, buildings, and other obstacles. Consequently, the 1930s represent the original epoch of data for much of California. With the exception of the San Diego region, only pre-1980 data were included in the

modeling effort. This cutoff date reflects the status of NGS's automated data base in early 1982--the time when the data were organized for project REDEAM. The San Diego data set was updated subsequent to 1982 to model coseismic deformation associated with the Imperial Valley earthquake (M=6.6) of 1979.

Direction observations represent the bulk of the archived data. Direction observations are organized into 'bundles'. Each bundle corresponds to a set of measurements made during a single observing session (usually spanning a few hours) from a particular instrumented station-mark to several targeted station-marks that are located within a few tens of kilometers. The standard error of a single direction observation ranges from 0.6 arc-seconds for first-order surveys to 1.2 arc-seconds for third-order surveys, that is, from 3 to 6 microradians.

Distance observations include both those that were taped and those that were obtained with electro-optical distance measuring (EDM) instrumentation. There are relatively few taped distance observations exceeding 1 km in length because the taping process is labor intensive and requires level, unobstructed terrain. The number and length of distance measurements increased drastically around 1960 with the advent of EDM equipment. Consequently, the set of distance observations included for REDEAM essentially spans only about two decades. The standard error of a distance observation is expressed by the empirical formula

$$\sigma = (a^2 + b^2 L^2)^{1/2} \quad (2.1)$$

where a and b are specified parameters and L is line length. For first-order surveys performed by NGS, $a = 0.015$ m and $b = 10^{-6}$ (Gergen 1975). For USGS distance measurements, $a \approx 0.003$ m and $b = 0.2 \times 10^{-6}$ (Savage and Prescott 1973). The USGS achieves the better precision by flying aircraft along observed lines-of-sight to obtain profiles of temperature and humidity needed to correct for refraction. For routine geodetic work, only endpoint (and sometimes midpoint) readings of temperature and humidity are recorded.

Astronomic azimuth observations also contributed to the REDEAM solutions. NGS's archives contain approximately 400 such observations for California, with each observation having a standard error of 7 microradians or greater. Azimuth observations, however, are distributed poorly through time with approximately 80 percent observed after 1960.

Appendix B contains a number of illustrations, two for each region, presenting the spatial and temporal distributions of the data. They reveal:

- The regional data sets overlap. In particular, the model for each region was derived from not only data within the region but also extending to a distance of 16 km beyond the region's geographic span. This data overlap was engineered to provide a measure of spatial continuity among the various models.
- A significant increase in the number of direction observations occurred around 1930 with the introduction of Bilby towers.
- A significant increase in the number of distance observations occurred around 1960 with the introduction of EDM instrumentation and again around 1970 with the start of USGS's monitoring program.

- Only three regions (Channel Islands, Los Angeles, and Bakersfield) include data that predate the San Francisco earthquake of 1906. For some regions (San Diego, San Bernadino, and Barstow) the pre-1906 data had not been automated when the corresponding models were derived. For the other 10 California regions, the pre-1906 data were intentionally excluded to avoid modeling the coseismic movement associated with the San Francisco earthquake.
- Only the San Diego region includes post-1980 data.

3. MATHEMATICAL MODEL

The mathematical formulation of the REDEAM models includes parameters for both the secular and episodic components of motion.

Secular motion is represented by dividing the region to be modeled into a mosaic of districts. The words, region and district, convey specific meanings in this report. The geographic area pertaining to a specific model is called a region. A district is one of several specifically designated areas associated with a region. Each district is allowed to translate, rotate, and undergo spatially homogeneous deformation at a constant rate with respect to time. By approximating the known geologic faults with district boundaries, the relative motion between districts represents the relative movement across these faults. This is not to say that all district boundaries correspond to faults. Some districts have been introduced simply to increase the spatial resolution of the secular motion. Figure 3.1 identifies the 10 districts that comprise the San Diego region. Similar maps for the other regions are located in appendix B.

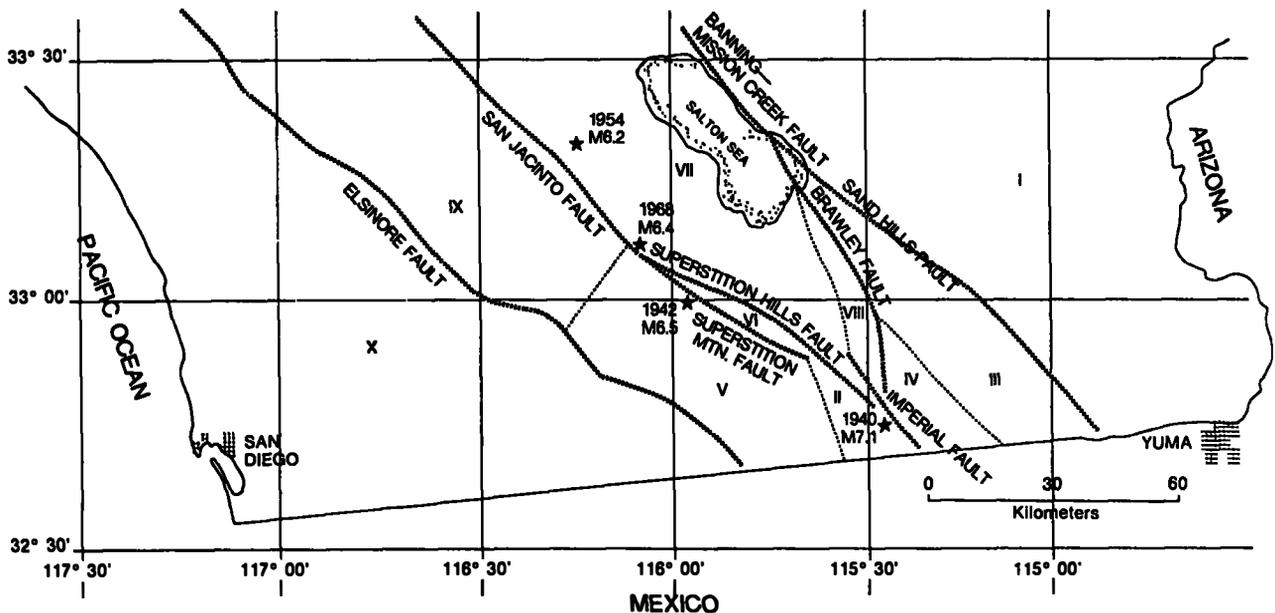


Figure 3.1.--A mosaic of 10 districts is used to model the secular motion of the San Diego region. Each Roman numeral identifies a district that can individually translate, rotate, and undergo homogeneous deformation at a constant rate with respect to time. District boundaries usually approximate geologic faults (hatched lines) so that relative motion between the districts corresponds to secular slip. The dashed lines denote district boundaries that do not correspond to faults. Dislocation theory is used to model the episodic motion associated with major earthquakes. The stars locate four modeled earthquakes identified by year of occurrence and magnitude.

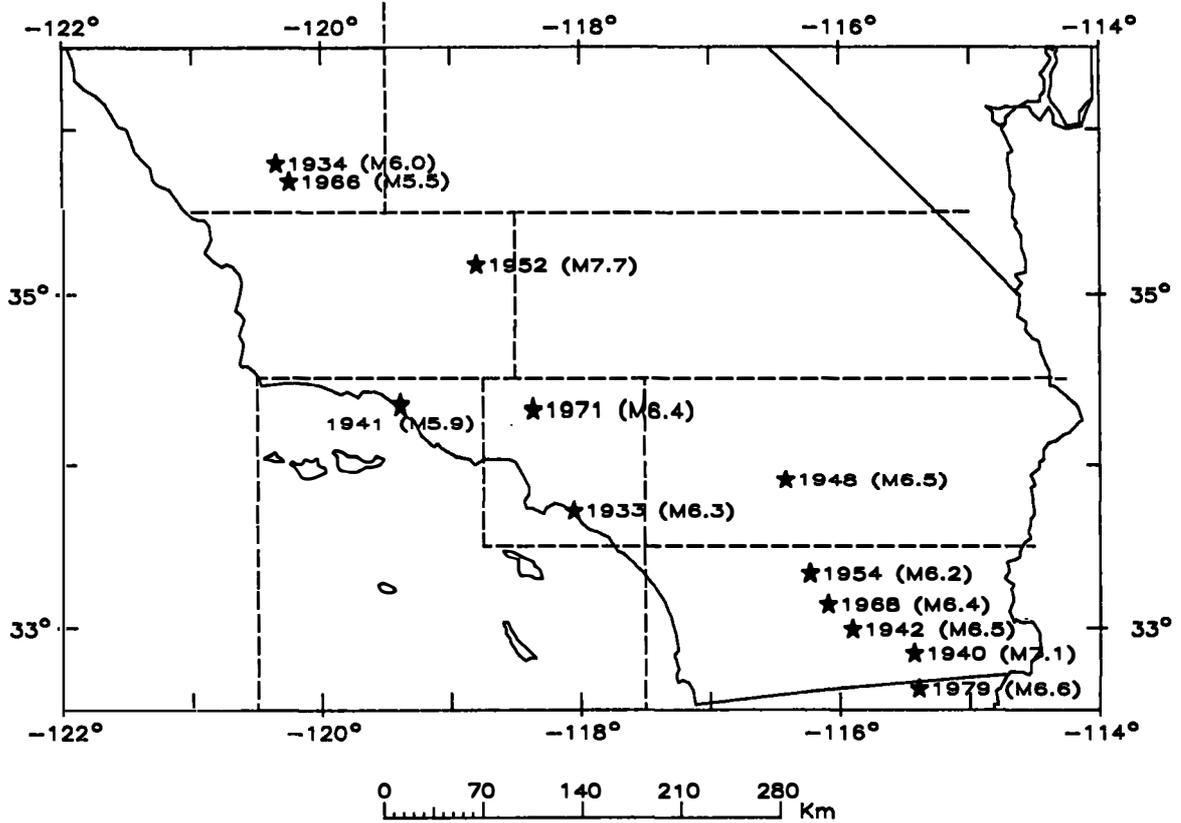


Figure 3.2.--Modeled earthquakes in California identified by year of occurrence and magnitude.

Modeled episodic motion corresponds to displacements associated with large earthquakes. For the episodic motion the Earth is considered to be an isotropic, homogeneous, elastic halfspace whose bounding plane represents the Earth's surface; that is, the Earth is represented as the set of points (x,y,z) with $z \leq 0$. Rectangular planes of finite dimensions are embedded in the halfspace to represent seismically active faults. The motions associated with an earthquake correspond to the displacements that the elastic halfspace undergoes in response to slip along the rectangular surfaces. This motion is given by the equations of dislocation theory (appendix A). The displacements are a function of the location, size, and orientation of the rectangles, as well as the amount and sense of the slip. Figure 3.2 identifies the earthquakes modeled for the 16 California regions.

More specifically, the mathematical model expresses the geodetic latitude $\phi_M(t)$ (positive north) and longitude $\lambda_M(t)$ (positive west) of a station M in district i at time t by the equation

$$\begin{bmatrix} \phi_M(t) \\ \lambda_M(t) \end{bmatrix} = \begin{bmatrix} \phi_M(t_0) \\ \lambda_M(t_0) \end{bmatrix} + \begin{bmatrix} a_{\phi^*}(i) & a_{\phi\phi}(i) & a_{\phi\lambda}(i) \\ a_{\lambda^*}(i) & a_{\lambda\phi}(i) & a_{\lambda\lambda}(i) \end{bmatrix} \begin{bmatrix} 1 \\ \phi_M(t_0) - \bar{\phi} \\ \lambda_M(t_0) - \bar{\lambda} \end{bmatrix} (t - t_0) + \sum_k r(t, t_k) \begin{bmatrix} A_k & B_k \\ C_k & D_k \end{bmatrix} \begin{bmatrix} s_k \\ d_k \end{bmatrix} \quad (3.1)$$

Here t_0 is a fixed time of reference. The preceding equation states that a station's coordinate values at time t equal its coordinate values at time t_0 plus a secular term, plus an episodic term.

In using eq. (3.1), we presume that secular motion is linear in time. In the secular term the variables $(\bar{\phi}, \bar{\lambda})$ represent reference coordinates that are selected prior to adjusting the model to the geodetic data. The variables $a_{\phi\phi}(i)$, $a_{\phi\lambda}(i)$, $a_{\lambda\phi}(i)$, $a_{\lambda\lambda}(i)$ are parameters to be estimated for district i . Equation (3.1) implies that the secular motion interior to each district is essentially homogeneous strain plus rotation at a constant rate with respect to time. (See appendix D.)

The expression that involves the summation sign in eq. (3.1) corresponds to the episodic motion and gives the change of the station's latitude $\Delta_k\phi$ and longitude $\Delta_k\lambda$ caused by strike slip s_k and dip slip d_k at time t_k on the k -th rectangle for each value of k . That is,

$$\begin{bmatrix} \Delta_k\phi \\ \Delta_k\lambda \end{bmatrix} = r(t, t_k) \begin{bmatrix} A_k & B_k \\ C_k & D_k \end{bmatrix} \begin{bmatrix} s_k \\ d_k \end{bmatrix} \quad (3.2)$$

The episodic time dependence is embedded in the step function $r(t, t_k)$ defined by the conditions:

$$\begin{array}{l} \text{for} \\ t_k < t_0 \end{array} \quad r(t, t_k) = \begin{cases} -1 & \text{if } t < t_k \\ 0 & \text{if } t > t_k \end{cases} \quad (3.3)$$

$$\begin{array}{l} \text{for} \\ t_k > t_0 \end{array} \quad r(t, t_k) = \begin{cases} 0 & \text{if } t < t_k \\ 1 & \text{if } t > t_k \end{cases}$$

Equation (3.3) prescribes that the slip on the k -th rectangle occurs instantaneously at time t_k . In eqs. (3.1) and (3.2) the quantities A_k , B_k , C_k , and D_k represent mathematical expressions involving the coordinates of station M as well as the location, orientation, and size of the k -th rectangle. Formulas for calculating these quantities are developed in the following paragraphs.

The fault center of the k -th rectangle is defined as the midpoint of the line segment that is the projected image of the rectangle onto the Earth's surface. (See appendix A.) Relative to no displacement at the fault center, let $U_{1k}(\phi, \lambda)$ denote the displacement due to slip on the k -th rectangle in the direction parallel to this rectangle's strike as realized at the point of the Earth's surface with coordinates (ϕ, λ) . Then by dislocation theory, functions $F_{1k}(\phi, \lambda)$ and $G_{1k}(\phi, \lambda)$ exist such that

$$U_{1k}(\phi, \lambda) = F_{1k}(\phi, \lambda)s_k + G_{1k}(\phi, \lambda)d_k \quad (3.4)$$

Similarly, functions $F_{2k}(\phi, \lambda)$ and $G_{2k}(\phi, \lambda)$ exist such that

$$U_{2k}(\phi, \lambda) = F_{2k}(\phi, \lambda)s_k + G_{2k}(\phi, \lambda)d_k \quad (3.5)$$

where $U_{2k}(\phi, \lambda)$ represents the displacement perpendicular to the strike at (ϕ, λ) . The functions $F_{jk}(\phi, \lambda)$ and $G_{jk}(\phi, \lambda)$ for $j = 1, 2$ are given in appendix A. In addition to being functions of position, these functions depend on the k -th rectangle's size, location, orientation (strike and dip angle), and on Poisson's ratio which characterizes crustal elasticity. For this study, Poisson's ratio is set equal to 0.24.

A differential approximation gives the displacements $\Delta_k \phi$ and $\Delta_k \lambda$ on the curved surface of the Earth relative to the assumption that the position $(\bar{\phi}, \bar{\lambda})$ does not move. That is,

$$\begin{bmatrix} \Delta_k \phi \\ \Delta_k \lambda \end{bmatrix} = r(t, t_k) \left\{ [T_k][H_k(\phi, \lambda)] \begin{bmatrix} s_k \\ d_k \end{bmatrix} - [T_k][H_k(\bar{\phi}, \bar{\lambda})] \begin{bmatrix} s_k \\ d_k \end{bmatrix} \right\} \quad (3.6)$$

where

$$[H_k(\phi, \lambda)] = \begin{bmatrix} F_{1k}(\phi, \lambda) & G_{1k}(\phi, \lambda) \\ F_{2k}(\phi, \lambda) & G_{2k}(\phi, \lambda) \end{bmatrix} \quad (3.7)$$

and

$$[T_k] = \begin{bmatrix} R^{-1} & 0 \\ 0 & (R \cos \phi_k)^{-1} \end{bmatrix} \begin{bmatrix} \cos \alpha_k & -\sin \alpha_k \\ -\sin \alpha_k & -\cos \alpha_k \end{bmatrix} \quad (3.8)$$

Here R = the Earth's radius;

ϕ_k = the latitude of k -th fault center;

α_k = the strike of the k -th rectangle measured clockwise from north. The matrix $[T_k]$ transforms displacements from a locally defined reference system, referred to the k -th fault plane, to the globally defined ellipsoidal reference system.

Eqs. (3.2) and (3.6) imply that

$$\begin{bmatrix} A_k & B_k \\ C_k & D_k \end{bmatrix} = [T_k] [H_k(\phi, \lambda) - H_k(\bar{\phi}, \bar{\lambda})] \quad (3.9)$$

With eq. (3.1) observations are entered into a least-squares process to estimate the unknown coordinates $(\phi_M(t_0), \lambda_M(t_0))$ for all M, the unknown parameters $a_{\phi\phi}(i), a_{\phi\lambda}(i), \dots, a_{\lambda\lambda}(i)$ for all values of i, and the slips s_k and d_k for all values of k. In the least-squares process an observation β_t at time t (a direction, a distance, or an azimuth) is first corrected for known systematic errors such as refraction and it is then projected onto an ellipsoidal reference surface so that the 'reduced' observation β_t' is expressible solely as a function h of mark coordinates. That is,

$$\beta_t' = h[\phi_P(t), \lambda_P(t), \phi_Q(t), \lambda_Q(t)] \quad (3.10)$$

where P and Q denote marks associated with the observation. Substituting into this equation from eq. (3.1), β_t' becomes a function of the coordinates $(\phi_P(t_0), \lambda_P(t_0))$ and $(\phi_Q(t_0), \lambda_Q(t_0))$, the parameters $a_{\phi\phi}(i), a_{\phi\lambda}(i), \dots, a_{\lambda\lambda}(i)$ for all i corresponding to the district(s) containing P and Q, and also a function of the slips s_k and d_k for all values of k. These expressions constitute the so-called 'observation equations' of the least squares process. The observations are weighted in the solution equal to the squared inverse of their respective standard errors.

4. PARAMETER ESTIMATION

Once the boundary of a particular region had been identified, we assembled all available geodetic data in that region into a single computer file called a TRAVDECK. Model parameters were estimated from these data via the least squares process. Here we briefly describe the software employed to obtain these estimates. The estimates themselves are tabulated in appendix C.

In addition to observational records, a TRAVDECK contains a positional record for each geodetic mark in the region. This positional record contains, among other information, preliminary estimates of the mark's latitude and longitude and adopted values for its orthometric height and geoid height. These positional quantities are needed for reducing observations to corresponding values on a mathematical reference ellipsoid. Preliminary latitude and longitude estimates are also needed because of the nonlinear relationship (eq. 3.10) between the observables and the estimated quantities.

The processing software, entitled MOTION III, was fashioned after the computer program TRAV10 which NGS employed during the late 1970s and early 1980s. TRAV10 performed a static network adjustment of a collection of horizontal data; that is, TRAV10 only estimated horizontal coordinates. For developing MOTION III, we modified the TRAV10 code to include the capability of estimating secular motion coefficients and earthquake-related slip vectors in addition to positional coordinates. MOTION III consists of six separate routines: QUIKR, REDUCE, REORD, ENDRY, CRUNCH, and POMOT.

QUIKR assigns a unique identification number to each geodetic mark and modifies the TRAVDECK so that updated positional records and updated observational records contain the numbers assigned to the mark(s) associated with those records. Subsequent routines use these identification numbers for internal computer reference.

REDUCE computes and applies "corrections" required to project observations onto the reference ellipsoid. The corresponding formulas are given by Schwarz (1978).

REORD determines an ordering of the geodetic marks to minimize computer-storage requirements of the array containing the least-squares normal-equations matrix when it is formed and manipulated in the subsequent routine CRUNCH. CRUNCH uses the same profile storage scheme as TRAV10 (Schwarz 1978). In particular, REORD employs the "banker's" algorithm (Snay 1976) to obtain a near optimal ordering of the marks.

BNDRY assigns each geodetic mark to its proper district. Recall that each region is partitioned into a mosaic of districts and, within a particular district, secular horizontal velocity is constrained to be a linear function of latitude and longitude per eq. (3.1). Also recall that district boundaries often, but not always, correspond to geologic faults. District boundaries were selected with the help of published geodetic diagrams for California (NOAA 1971) which show the relative locations of geodetic marks and geologic faults. The 1:250,000 series of the Geologic Map of California (Olaf P. Jenkins edition available from the California Division of Mines and Geology, San Francisco) was also frequently consulted. The boundary of a particular district is defined by specifying the coordinates for the vertices of a simple closed polygon, that is, a polygonal loop that does not intersect itself. Such a loop bounds two disjoint areas on the Earth's surface. In BNDRY the district corresponding to a particular polygon is the bounded area that does not contain the North Pole. In BNDRY, every mark is assigned to one and only one district. Appendix C tabulates vertices defining the districts for the various regions. See also appendix B for maps of the various regions that illustrate the spatial extent of the districts.

CRUNCH performs the least squares adjustment for estimating user-designated model parameters. (See chapter 3.) In particular, CRUNCH computes the coefficients of the observation equations, forms and solves the normal equations, and then computes residuals. The observation-equation coefficients are computed by an application of the chain rule. That is, if β'_t is a reduced observation, say a distance, observed at time t , and if α is a parameter to be estimated, say a secular motion coefficient, then from eq. (3.10) the coefficient $(\partial\beta'_t)/(\partial\alpha)$ may be computed by the equation:

$$\begin{aligned} \frac{\partial\beta'_t}{\partial\alpha} = & \left(\frac{\partial h}{\partial\phi_P(t)} \right) \left(\frac{\partial\phi_P(t)}{\partial\alpha} \right) + \left(\frac{\partial h}{\partial\lambda_P(t)} \right) \left(\frac{\partial\lambda_P(t)}{\partial\alpha} \right) + \left(\frac{\partial h}{\partial\phi_Q(t)} \right) \left(\frac{\partial\phi_Q(t)}{\partial\alpha} \right) \\ & + \left(\frac{\partial h}{\partial\lambda_Q(t)} \right) \left(\frac{\partial\lambda_Q(t)}{\partial\alpha} \right) \end{aligned} \quad (4.1)$$

The partials $\frac{\partial h}{\partial\phi_P(t)}$, $\frac{\partial h}{\partial\lambda_P(t)}$, $\frac{\partial h}{\partial\phi_Q(t)}$, $\frac{\partial h}{\partial\lambda_Q(t)}$ in eq. (4.1) are exactly the coefficients that would be computed for a static horizontal network adjustment, and the appropriate formulas are given by Schwarz (1978). The other four partials on the right-hand-side of eq. (4.1), namely those involving partial derivatives with respect to α , may be computed by differentiating eq. (3.1) with respect to α . To compute these latter four partials, the following information must be provided initially to CRUNCH: (1) the time of reference t_0 , (2) the coordinates $(\bar{\phi}, \bar{\lambda})$ for the origin of reference, and (3) the dates of the earthquakes together with the various parameter values defining location, size, and orientation for the corresponding dislocation surfaces. (See appendix A.)

POMOT prints the parameter estimates obtained in CRUNCH with the associated correlation matrix for the crustal motion parameters. POMOT will also compute certain derived quantities if requested; namely, (1) station positions at specified times, (2) displacements attributed to earthquakes at specified locations, and (3) secular strain-rate parameters for specified districts. In addition POMOT prints residuals and normalized residuals for the observations.

5. MODEL IMPLEMENTATION

The derived crustal motion models were used to "update" all archived horizontal data in the United States to the common date December 31, 1983. Thus an observation measured in 1940, for example, was updated to approximate the value that would be expected if it were remeasured on December 31, 1983. These updated observations were then entered into a static horizontal network adjustment to determine latitudes and longitudes for the North American Datum of 1983 (NAD 83) geodetic reference system. This chapter describes the algorithm used for computing the crustal motion "corrections" for updating observations.

Crustal motion corrections were applied to all observations that involve two stations, namely, direction, azimuth, and distance observations. Crustal motion corrections were not applied to observations involving only a single station, namely, Doppler positioning observations, because the data used to generate the models are insensitive to 'absolute' motion. The effect of not correcting the Doppler observations should be insignificant because all archived Doppler observations were performed after 1970 and because the horizontal components of these observations have meter-level uncertainties.

To correct an observation between two stations P and Q measured at time t_1 to its corresponding value at time t_2 , approximate horizontal coordinates for P and Q at the preselected reference time $t_0 = \text{January 1, 1950}$ are needed. Because we are interested only in changes to observations from one time to another, the selection of t_0 was rather arbitrary (1950 corresponds to the approximate weighted midpoint in the observation dates), and the station coordinates at time t_0 did not need to be extremely accurate. For our purpose, the NAD 27 coordinates of P and Q constitute sufficiently accurate estimates for $\phi_P(t_0)$, $\lambda_P(t_0)$, $\phi_Q(t_0)$, and $\lambda_Q(t_0)$.

Step 1: Determine the regions R_P and R_Q that contain stations P and Q, respectively. Recall that crustal motion models were formulated for 19 mutually disjoint regions. Points in the United States that are not contained in any of these 19 regions are assigned to a twentieth, complementary region where the crustal motion model for region 20 is defined as the null model, that is, no motion.

Step 2: Let T denote the type of observation to be corrected. Use the model for R_P to compute $\phi_P(t_1)$, $\lambda_P(t_1)$, $\phi_Q(t_1)$, and $\lambda_Q(t_1)$ according to eq. (3.1), and let $b(t_1, R_P)$ denote the hypothetical observation of type T that would be measured at time t_1 between P and Q given these coordinates. Similarly use the model for R_P to compute $\phi_P(t_2)$, $\lambda_P(t_2)$, $\phi_Q(t_2)$, and $\lambda_Q(t_2)$ according to eq. (3.1), and let $b(t_2, R_P)$ denote the hypothetical observation of type T that would be measured at time t_2 between P and Q given these coordinates.

Step 3: Use the model for R_Q to compute, as in step 2, the hypothetical observations $b(t_1, R_Q)$ and $b(t_2, R_Q)$ of type T between P and Q at times t_1 and t_2 , respectively.

Step 4: If b denotes the value of the actual observation measured at time t_1 , then

$$b' = b + 1/2 [b(t_2, R_P) - b(t_1, R_P)] + 1/2 [b(t_2, R_Q) - b(t_1, R_Q)] \quad (5.1)$$

is the corrected observation corresponding to time t_2 .

Note that the correction in eq. (5.1) represents the average of two estimates for the crustal motion between times t_1 and t_2 : one estimate from the model for region R_P and the other from the model for region R_Q . This averaging process minimizes possible discrepancies between different models for observations that cross regional boundaries. Recall that regional boundaries are artifacts which, unlike most district boundaries, do not correspond to geologic faults.

Note also that the algorithm does not require that observation b be projected onto the reference ellipsoid to compute b' , even though $b(t_1, R_P)$, $b(t_2, R_P)$, $b(t_1, R_Q)$, and $b(t_2, R_Q)$ may correspond to hypothetical observations on the reference ellipsoid. Moreover, any other data correction, for example, refraction, can be applied either before or after the crustal motion correction, provided the value of the correction does not strongly depend on the date of observation or on the horizontal coordinate changes at the level of expected crustal motion.

Finally note that REDEAM models should not be used to compute positions that predate the 1906 San Francisco earthquake (April 18) for marks in California which are north of the 35.5° parallel of latitude.

6. MODEL EVALUATION

Evaluation of the REDEAM models constitutes an ongoing process. Five papers, in addition to this one, have already appeared in print, and other studies are anticipated. Four of these five publications discuss specific models and compare these models both with results derived by independent investigators and with current geophysical theories. In particular, these publications discuss the model for the San Diego region (Snay et al. 1983), the model for the Los Angeles region (Cline et al. 1984), the models for the San Francisco, Santa Rosa, and Ukiah regions (Cline et al. 1985), and the model for the Fallon, Nevada, region (Snay et al. 1985). The fifth publication (Snay et al. 1986) presents an overall evaluation of the 16 regional models spanning California. The fifth paper also discusses ideas for improving the models. This chapter recaps some of the material appearing in these five publications.

Figure 6.1 portrays the derived shear strain pattern for California. (See appendix D for a discussion of strain.) Because we have chosen to model the 16 California regions independently, artificial discontinuities in secular velocity occur across regional boundaries. These discontinuities have an rms value of ~ 5 mm/yr even after individual models are adjusted to be mutually compatible. We partitioned California into 16 regions to limit the size of the individual data sets. Even with 16 regions the data sets were often too large for a person to become adequately familiar with

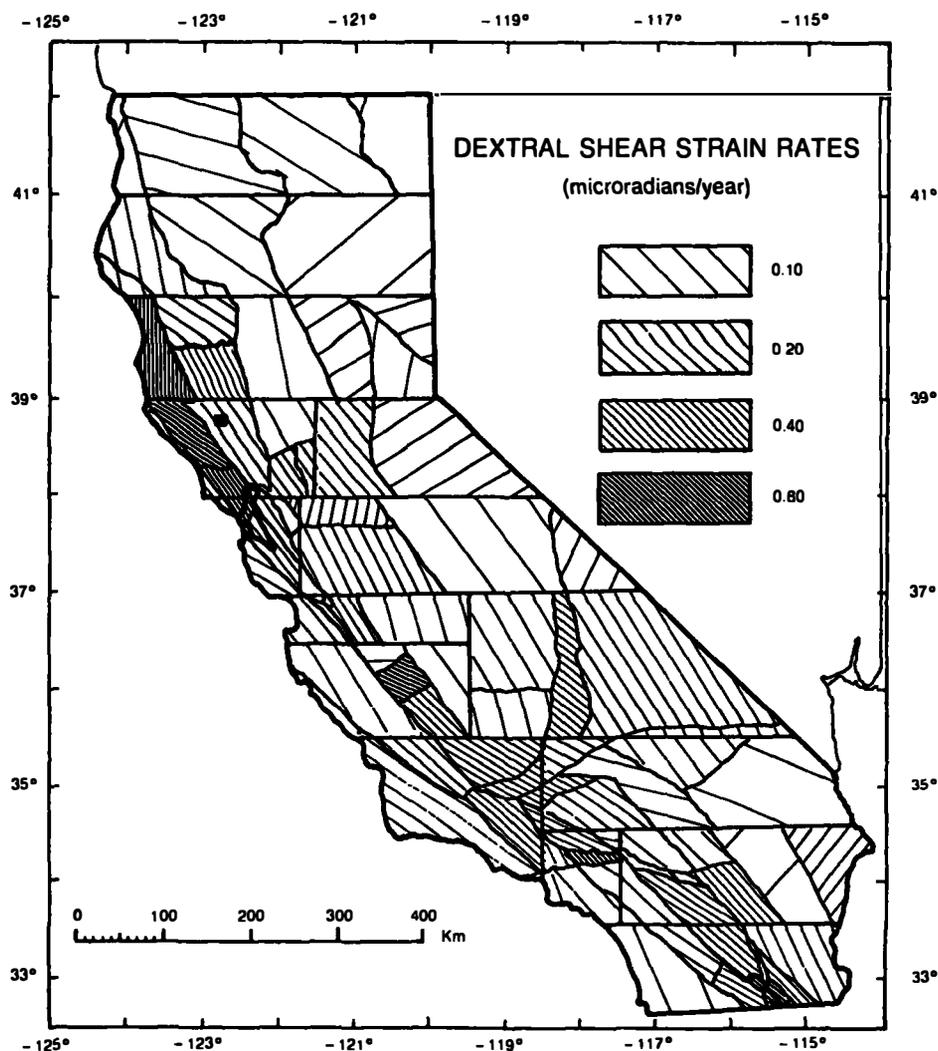


Figure 6.1.--Secular shear strain pattern for California as derived from historical geodetic data. The line patterns designate the directions and magnitudes (engineering units) of maximum dextral (right-lateral) shear strain rates for the mosaic of districts. Shearing between the North American and Pacific plates dominates the regional stress field producing an overall northwest-southeast trend for the direction of maximum dextral shear strain. The secular motion is assumed to be linear in time and excludes the movements associated directly with earthquakes of magnitude six and greater.

all the data contained in a region. Also, we limited the sizes of the data sets because model development was largely a trial-and-error procedure: (1) the data sets had to be screened for blunders, (2) district boundaries had to be resolved, and (3) appropriate earthquake fault parameters had to be determined. It was not uncommon for us to perform as many as 10 adjustments of a region's data to various models in search of the best representation for regional crustal movement. Moreover, each such adjustment strained computer resources to the extent that frequently the computer could only execute the adjustment over the weekend. Now, having modeled the 16 regions, we feel sufficiently familiar with the data, the techniques, and the geophysics to undertake a simultaneous adjustment of all California data to a single model. We need only develop the appropriate software and allocate the necessary resources.

Figure 6.1 presents only the shear components of the secular deformation pattern. Data limitations render our estimates of the other components (rotation and dilatation) questionable.

The rotation estimates depend on the astronomic azimuth data. The geodetic archives contain less than 400 such azimuth observations for California, with each observation having a standard deviation of 7 microradians or greater. Moreover, the azimuth observations are distributed poorly through time, with approximately 80 percent observed since 1960. Consequently, rotation uncertainties (σ) for districts are about 0.1 microradians per year (appendix D). Said differently, for every 100 km separating two stations, an uncertainty of 1 cm/yr exists in the transverse component of the secular velocity between these stations. These velocity uncertainties are similar in magnitude to the expected secular velocities between stations on opposite sides of the state (Minster and Jordan 1978). The near-future availability of space-based data, providing 3-dimensional coordinate differences between stations to centimeter-level precision over lines exceeding 100 km in length, will enhance our estimates of regional rotation.

The dilatation component of the deformation depends on the collection of distance observations. Prior to the initial deployment of electro-optical distance measuring equipment in 1959, distances were laboriously taped. Consequently, the set of distance observations essentially spans only the last two decades. Moreover, unmodeled systematic errors, having magnitudes on the order of several parts in 10^6 , are thought to contaminate much of the distance data [see table 2 in Snay et al. (1983) and table 4 in Cline et al. (1984)]. Considering the short time base, such errors could easily bias our estimated dilatation rates at the level of a few parts in 10^7 per year--a level that approximates in magnitude the dilatation rates that have been accurately measured with EDM by the U.S. Geological Survey (USGS) for selected areas of California (Savage 1983). These USGS EDM data are more precise than most of the other archived distance measurements because the USGS flew aircraft over the observed lines-of-sight to obtain temperature and humidity profiles to better correct for refraction. For routine geodetic work, only endpoint meteorological readings are recorded. The highly precise USGS EDM data measured before 1979 were included in the REDEAM modeling project. These data profoundly helped to subdue biases in our dilatation rate estimates, yet their effect was understandably limited to the areas that these data cover. The USGS monitoring program continues through time to the present and has expanded gradually to cover a greater area. Our planned second generation model will benefit from the inclusion of the USGS data measured since 1979.

Also, to better address the problem of systematic errors in the distance data, a future generation model will probably include several scale parameters. These parameters would be introduced on the premise that a large part of the systematic error manifests itself as scale factors, each such factor being common to a group of distances observed with the same instrument. Such an error could be caused, for example, by instrumental miscalibration. These scale parameters would be estimated simultaneously with other model parameters via the least squares process.

The secular shear strain pattern (fig. 6.1) is well determined for California because of the preponderance of triangulation data. The earliest data are from the 1850s, but most pre-1900 data are concentrated along the coast where they primarily supported navigational charting. For most California areas, then, the first geodetic data were

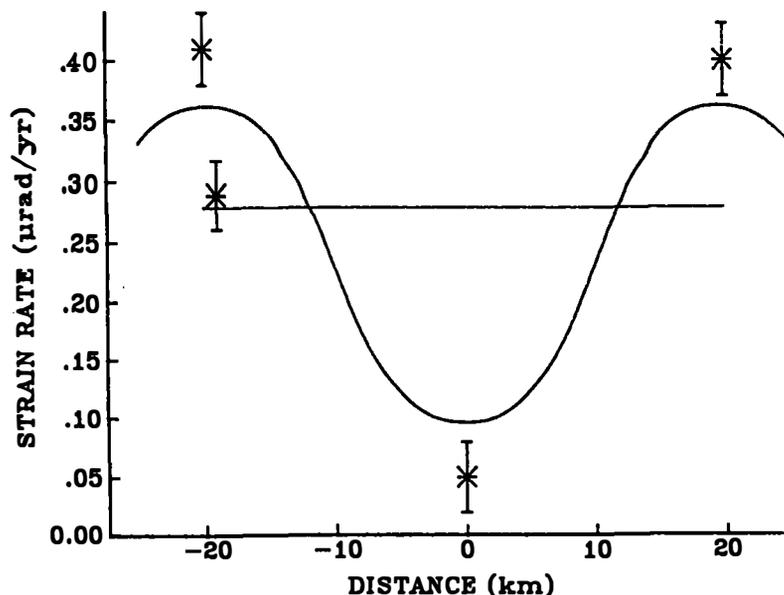


Figure 6.2.--The San Bernardino regional model contains a district corresponding in area to the 40-km wide strip of land between the San Andreas and San Jacinto fault zones. The above graph represents a cross section of this strip with the San Andreas fault zone located at +20 km and the San Jacinto at -20 km. The REDEAM model presumes that the secular shear-strain rate is homogeneous across the strip. The horizontal line represents the REDEAM rate. Asterisks (with 1σ error bars) represent four localized strain-rate estimates that reveal significant spatial variation across the strip. The curve corresponds to a hypothetical representation of this variation as proposed by King and Savage (1983).

observed in the 1930s, corresponding in time to the introduction of the Bilby tower. Consequently, the shear strain-rate estimates correspond essentially to a 50-year time interval, 1930-80. These estimates also represent spatial averages over several tens of kilometers; that is, the models presume that secular strain is spatially homogeneous within each district. A localized study of USGS EDM data in southern California (King and Savage 1983) demonstrates the need for a model allowing greater spatial resolution (fig. 6.2).

Table C.3 presents parameters for modeled earthquakes in California and Nevada. In addition the parameters for some of these earthquakes have been discussed by Snay et al. (1983 and 1985) and Cline et al. (1984). The locations, dimensions, and orientations of the various rectangles that represent the fault planes were specified after reviewing the geologic and seismic literature. We scanned this literature for hypocenters, aftershock zones, focal mechanisms, and surface ruptures. The uncertainties (1σ) associated with the estimated components of the coseismic slip vectors imply decimeter-level resolution. We believe these uncertainties, however, are overly optimistic because the locations, dimensions, and orientations of the dislocation planes were introduced into the solution as if they were perfectly known. Also, the derived uncertainties are optimistic because of the ambiguities that exist in discriminating between coseismic and secular fault slip. Consequently, we contend that the existing geodetic data in general suffice only to resolve those coseismic slips exceeding a meter in magnitude. A meter of slip essentially requires an event with $M \geq 7$ (Bonilla and Buchanan 1970), such as the 1940 Imperial Valley earthquake ($M = 7.1$) or the 1952 Kern County earthquake ($M = 7.7$). Only a few seismic events with $6 \leq M \leq 7$ have their coseismic slips well determined by existing horizontal data.

It may be unfair to attribute this poor resolution of coseismic slip completely to data limitations. Some of the problem, more than likely, rests with our employed mathematical representation of earthquake movement. Our technique assumes that coseismic slip is constant over a rectangle whose dimensions are typically on the order of tens of kilometers. Current theory for earthquake mechanics, however, favors the existence of significant spatial variations in slip over the rupture surface. The newer theory promotes the concepts of "asperities" and "barriers" that strongly influence the distribution of coseismic slip (Aki 1984). Both terms refer to strong patches on the fault that are resistive to breaking. Considerable study is yet needed to identify and classify these fault features and then develop analytic expressions that will more realistically represent the episodic motion.

7. SUMMARY

Crustal motion models were produced for 19 total regions, 16 of which combine to cover all of California, with one model each for parts of Nevada, Hawaii, and Alaska. We view this accomplishment not as an end but as the beginning of a new approach to the description of crustal motion--an approach that integrates both large volumes of data as well as various types of data. In producing these first models, we have gained such an increased understanding of the data and of the geophysics that work on the second generation model is already in the planning stage. Table 7.1 summarizes the contemplated improvements to the original regional models. Some of these improvements will be incorporated into the second generation model; others will come only with later generations. The requirement for model improvement is also driven by the continual accession of new and often more accurate data. One of our goals is to synthesize the space-based data into the modeling process as techniques like the Global Positioning System, mobile Very Long Baseline Interferometry, and Satellite Laser Ranging become more operational. Other geophysical data types might also be synthesized into the modeling process. Our original models indirectly used processed seismic data to define certain coseismic fault parameters. These models also indirectly used geologic information to delineate district boundaries. Future models could possibly benefit from additional data types, such as measurements of fault creep, gravity, in situ stress, or heat flow. The synthesis of these various data types will necessarily lead to the development of physically more realistic models.

Table 7.1.--Proposed improvements to original crustal motion models

Improvement	Purpose
Adjust California to a single model	To remove artificial discontinuities along regional boundaries and to "push back" boundary effects.
Incorporate space-based data	To constrain the accumulation of systematic errors over distance and to better resolve regional rotation.
Introduce scale parameters	To remove systematic scale biases suspected among many groups of distance observations.
Incorporate more recent USGS EDM data	To better resolve secular dilatation rates.
Replace current expressions for secular motion with individual station velocity parameters	To provide better spatial resolution of the secular motion.
Refine earthquake models	To represent episodic motion more realistically.

APPENDIX A.--DISLOCATION EQUATIONS

From the analysis of the surface displacements associated with the 1906 San Francisco earthquake, Hayford and Baldwin (1908) found: (1) that points on opposite sides of the disrupted San Andreas fault moved in opposite directions, (2) that displacements were less the greater the distance from the fault, and (3) that the directions of the displacements correlated well with fault orientation. These discoveries helped establish the then controversial theory that fault slip causes the vibrations in the ground that we know as earthquakes. The association between fault slip and earthquakes extends even further. Numerous studies, for example, Chinnery (1961), Savage and Hastie (1966), Hastie and Savage (1970), Stein and Thatcher (1981), and Snay et al. (1982), have since demonstrated that coseismic crustal deformation relates quantitatively to fault slip. Steketee (1958a,b) constructed the mathematical foundation for this relationship by translating relevant concepts from the theory that deals with dislocations or offsets in the lattice structure of crystals. Even the name, dislocation theory, has been carried from crystallography to the realm of geophysics. There is a basic difference, however. Whereas crystallography deals with dislocations in a discrete structure, geophysics is concerned with dislocations in a continuum.

A dislocation surface in a continuous 3-dimensional body is, briefly, a surface of discontinuity in displacement. It may be understood conceptually by visualizing its formation in the following process:

- a) make a cut inside the body to form an arbitrary new surface S;
- b) apply relative displacement of slip-type to the two faces of the cut;
- c) rejoin the faces in their new positions.

In this way the body regains its stress continuity, but undergoes deformation to accommodate the displacement discontinuity across S.

In geophysical applications the body encompassing the dislocation surface is customarily a homogeneous, isotropic, elastic halfspace whose free surface represents the Earth's surface. For the special case of uniform slip over a rectangular dislocation surface whose upper edge parallels the free surface (fig. A.1), closed analytical formulas for computing the associated displacement field have been published by several authors. In particular, see Okada (1985) for an overview of existing literature. REDEAM software incorporates the formulas published by Dunbar (1977).

The equations assume a right-hand coordinate system (fig. A.1) with the 1-axis and 2-axis defining a horizontal plane and the 3-axis corresponding to the vertical dimension. The 1-axis parallels the strike of the rectangle, the 2-axis is positive in the direction of dip, and the 3-axis is positive into the halfspace. The origin of the coordinate system corresponds to a specific point on the line where the plane of the rectangle intersects the free surface. If the rectangle is projected orthogonal to its strike onto this line, then the origin corresponds to the midpoint of the projected rectangle. If $X(x_1, x_2, x_3)$ denotes an arbitrary point in the halfspace, the equations yield the displacements $U_1(X)$ and $U_2(X)$ which parallel the 1-axis and the 2-axis, respectively, as a function of

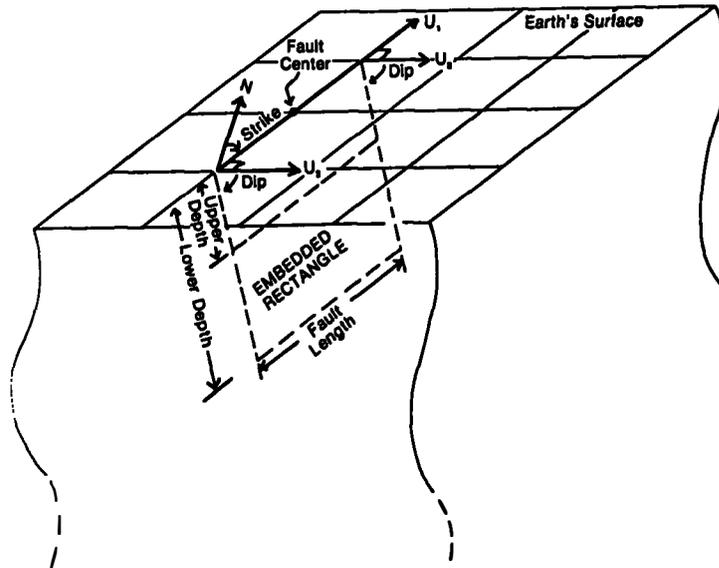


Figure A.1.--Given a rectangle embedded in an isotropic homogeneous elastic halfspace, dislocation theory describes how each point of the halfspace moves in response to relative slip between opposing rectangular faces. It is assumed that the upper and lower edges of the rectangle are parallel to the Earth's surface and that the slip is pure translation in the plane of the rectangle, in particular, no rotation. The figure identifies a set of parameters sufficient to specify the size, location, and orientation of the rectangle. The motion depends on these parameters, the Poisson ratio of the crust, and the slip vector. A constant slip vector over the rectangle is required by the mathematical formulation employed in this report. Spatial variation of fault slip is introduced by approximating the fault with multiple disjoint rectangles. The axes U_1 and U_2 define a local reference system for horizontal motion. The U_1 axis is positive along the fault strike. The U_2 axis is oriented 90° clockwise from U_1 and is positive in the direction of the dip.

- s strike slip
- d dip slip
- θ fault dip
- D_U upper fault depth (measured along dip)
- D_L lower fault depth (measured along dip)
- H fault half-length
- ν Poisson's ratio .

The equations are of the form

$$U_1(X) = F_1(X) s + G_1(X) d \quad (A.1)$$

where

$$F_1(X) = \int_{D_U}^{D_L} \int_{-H}^H f_1(X, \xi) d\xi \quad (A.2)$$

for $\xi(\xi_1, \xi_2, \xi_3)$ denoting points on the rectangle.

Consequently,

$$\begin{aligned} F_1(X) = & f_1(X, \xi) \Big|_{\xi = \xi(+H, D_L \cos \theta, D_L \sin \theta)} \\ & - f_1(X, \xi) \Big|_{\xi = \xi(-H, D_L \cos \theta, D_L \sin \theta)} \\ & - f_1(X, \xi) \Big|_{\xi = \xi(+H, D_U \cos \theta, D_U \sin \theta)} \\ & + f_1(X, \xi) \Big|_{\xi = \xi(-H, D_U \cos \theta, D_U \sin \theta)} \end{aligned} \quad (A.3)$$

Similarly

$$G_1(X) = \int_{D_U}^{D_L} \int_{-H}^H g_1(X, \xi) d\xi \quad (A.4)$$

To express the equations $f_1(X, \xi)$ and $g_1(X, \xi)$, it is convenient to introduce the notation:

(Mansinha and Smylie 1971):

$$\xi_0 = (\xi_2^2 + \xi_3^2)^{1/2}$$

$$r_2 = x_2 \sin \theta - x_3 \cos \theta \quad q_2 = x_2 \sin \theta + x_3 \cos \theta$$

$$r_3 = x_2 \cos \theta + x_3 \sin \theta \quad q_3 = -x_2 \cos \theta + x_3 \sin \theta$$

$$R^2 = (x_1 - \xi_1)^2 + (x_2 - \xi_2)^2 + (x_3 - \xi_3)^2$$

$$= (x_1 - \xi_1)^2 + r_2^2 + (r_3 - \xi_0)^2$$

$$Q^2 = (x_1 - \xi_1)^2 + (x_2 - \xi_2)^2 + (x_3 + \xi_3)^2$$

$$= (x_1 - \xi_1)^2 + q_2^2 + (q_3 + \xi_0)^2$$

$$k^2 = (x_1 - \xi_1)^2 + q_2^2$$

$$h^2 = q_2^2 + (q_3 + \xi_0)^2$$

$$\xi_2 = \xi_0 \cos \theta \quad \xi_3 = \xi_0 \sin \theta$$

Thus the expressions involving strike slip are

$$f_1(X, \bar{c}) = \frac{1}{8\pi(1-\nu)} \left[\begin{aligned} & 2(1-\nu) \tan^{-1} \left\{ \frac{(x_1 - \xi_1)(r_3 - \xi_0)}{r_2 R} \right\} + \frac{r_2(x_1 - \xi_1)}{R(R + r_3 - \xi_0)} \\ & - 2(1-\nu) \tan^{-1} \left\{ \frac{(x_1 - \xi_1)(q_3 + \xi_0)}{q_2 Q} \right\} - (x_1 - \xi_1) \left[\frac{(3-4\nu)x_2 \sin \theta + x_3 \cos \theta}{Q(Q + q_3 + \xi_0)} \right. \\ & \left. - 2q_2 x_3 \sin \theta \left\{ \frac{1}{Q^3} - q_3 \frac{(2Q + q_3 + \xi_0)}{Q^3(Q + q_3 + \xi_0)^2} \right\} \right] \\ & - 4(1-\nu)(1-2\nu) \tan \theta \left\{ 2 \tan \theta - \tan^{-1} \left[\frac{(k - q_2 \cos \theta)(Q - k) + k(q_3 + \xi_0) \sin \theta}{(x_1 - \xi_1)(q_3 + \xi_0) \cos \theta} \right] \right. \\ & \left. + \frac{x_1 - \xi_1}{Q + x_3 + \xi_3} \right\} \end{aligned} \right]$$

$$\text{As } \theta \rightarrow \pi/2, \quad - \tan \theta \left(\right) \rightarrow + \frac{x_2(x_1 - \xi_1)}{2(Q + x_3 + \xi_3)^2}$$

$$f_2(X, \xi) = \frac{1}{8\pi(1-\nu)} \left[\begin{aligned} & - (1-2\nu)\sin\theta \ln(R+r_3-\xi_0) + \frac{r_2 \cos\theta}{R} + \frac{r_2^2 \sin\theta}{R(R+r_3-\xi_0)} \\ & - (1-2\nu)\sin\theta \ln(Q+q_3+\xi_0) + \frac{((3-4\nu)x_2 \sin\theta + x_3 \cos\theta) \cos\theta - 2x_3 \sin^2\theta}{Q} \\ & - \sin\theta \cdot \frac{[q_2((3-4\nu)x_2 \sin\theta + x_3 \cos\theta) + 2x_3(q_2 \cos\theta - q_3 \sin\theta)]}{Q(Q+q_3+\xi_0)} \\ & + 2q_2 x_3 \sin\theta \left\{ \frac{[x_2 - \xi_2 + q_3 \cos\theta]}{Q^3} - q_2 q_3 \sin\theta \cdot \frac{(2Q+q_3+\xi_0)}{Q^3(Q+q_3+\xi_0)^2} \right\} \\ & + 4(1-\nu)(1-2\nu) \tan\theta \left\{ \tan\theta \cdot [\ln(Q+x_3+\xi_3) - \sin\theta \ln(Q+q_3+\xi_0)] \right. \\ & \left. - \frac{x_2 - \xi_2}{Q+x_3+\xi_3} \right\} . \end{aligned} \right]$$

$$\text{As } \theta \rightarrow \pi/2, \tan\theta(\cdot) + \frac{1}{2} \left\{ \frac{x_3 + \xi_3}{Q+x_3+\xi_3} + \ln(Q+x_3+\xi_3) + \frac{x_2^2}{(Q+x_3+\xi_3)^2} \right\} .$$

and the expressions involving dip slip are

$$g_1(X, \xi) = \frac{1}{8\pi(1-\nu)} \left[\begin{aligned} & \frac{r_2}{R} + \frac{(3-4\nu)x_2 \sin\theta + x_3 \cos\theta}{Q} - \frac{2q_2 x_3 \xi_3}{Q^3} - \frac{4(1-\nu)q_2 x_3 \sin\theta}{Q(Q+q_3+\xi_0)} \\ & + 4(1-\nu)(1-2\nu) \left\{ \tan\theta \cdot [\sin\theta \ln(Q+x_3+\xi_3) - \ln(Q+q_3+\xi_0)] \right. \\ & \left. - \frac{(x_2 - \xi_2) \sin\theta}{Q+x_3+\xi_3} \right\} \end{aligned} \right]$$

$$\text{As } \theta \rightarrow \pi/2, (\cdot) \rightarrow 0$$

$$g_2(X, \xi) = \frac{1}{8\pi(1-\nu)} \left[\begin{aligned} & - (1-2\nu)\sin\theta \ln(R+x_1-\xi_1) + 2(1-\nu)\cos\theta \cdot \tan^{-1} \left\{ \frac{(x_1-\xi_1)(r_3-\xi_0)}{r_2^R} \right\} \\ & + \frac{r_2(x_2-\xi_2)}{R(R+x_1-\xi_1)} + (1-2\nu)\sin\theta \ln(Q+x_1-\xi_1) - 2(1-\nu)\cos\theta \cdot \tan^{-1} \left\{ \frac{(x_1-\xi_1)(q_3+\xi_0)}{q_2 Q} \right\} \\ & + \frac{(3-4\nu)r_2(x_2-\xi_2) + 2\xi_3 x_3 \sin\theta + 4(1-\nu)x_3[r_3 + \xi(\sin^2\theta - \cos^2\theta)]}{Q(Q+x_1-\xi_1)} \\ & + \frac{4(1-\nu)x_3(x_1-\xi_1)\sin^2\theta}{Q(Q+q_3+\xi_0)} - 2q_2 x_3 \xi_3 (x_2-\xi_2) \frac{(2Q+x_1-\xi_1)}{Q^3(Q+x_1-\xi_1)^2} \\ & + 4(1-\nu)(1-2\nu)\sin\theta \left\{ 2\tan\theta \cdot \tan^{-1} \left[\frac{(k-q_2\cos\theta)(Q-k) + k(q_3+\xi_0)\sin\theta}{(x_1-\xi_1)(q_3+\xi_0)\cos\theta} \right] \right. \\ & \left. + \frac{x_1-\xi_1}{Q+x_3+\xi_3} \right\} . \end{aligned} \right]$$

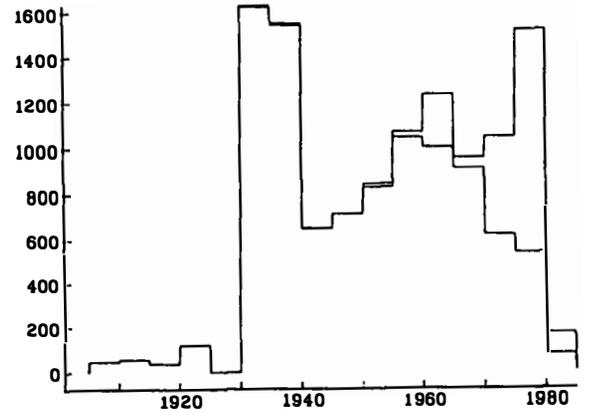
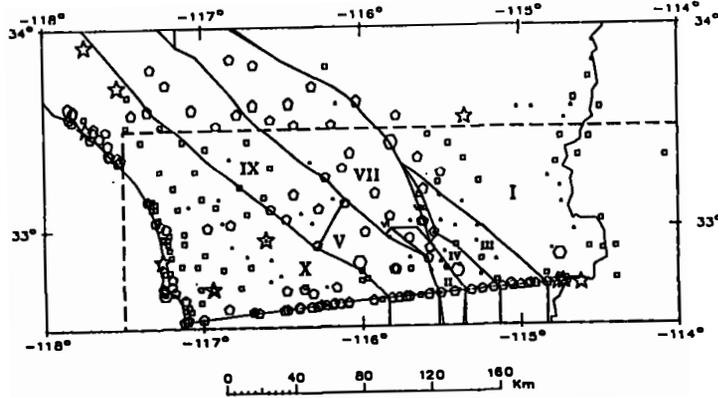
As $\theta \rightarrow \pi/2$, $\cdot(\cdot) \rightarrow 0$.

APPENDIX B.--DATA DISTRIBUTIONS

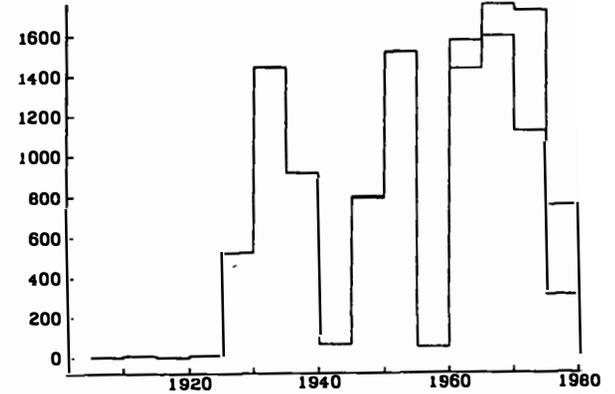
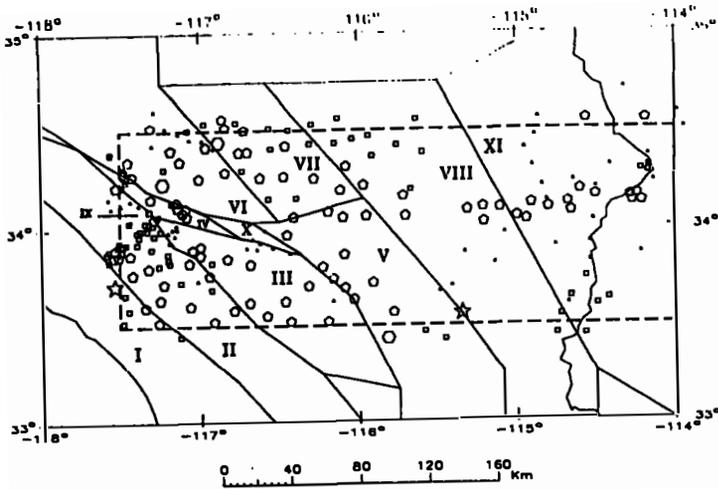
This appendix contains various maps and histograms, a pair for each of the 16 REDEAM regions of California. Together, a region's map and histogram depict the spatial and temporal distributions of the geodetic data used in developing the corresponding model.

On the maps, regional boundaries are plotted as dashed lines and district boundaries as solid lines. Roman numerals identify the various districts. Map symbols (triangles, squares, pentagons, hexagons, and stars) locate geodetic marks and specify the time interval between the earliest and latest observations at these marks.

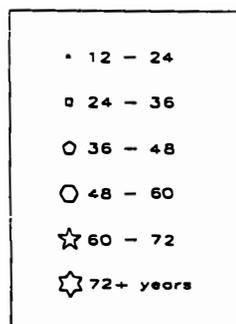
In the histograms the numbers of direction and distance observations for various 5-year intervals are plotted as a function of time. The lower column height represents the number of direction observations, and the height difference between the lower and upper columns represents the number of distance observations.

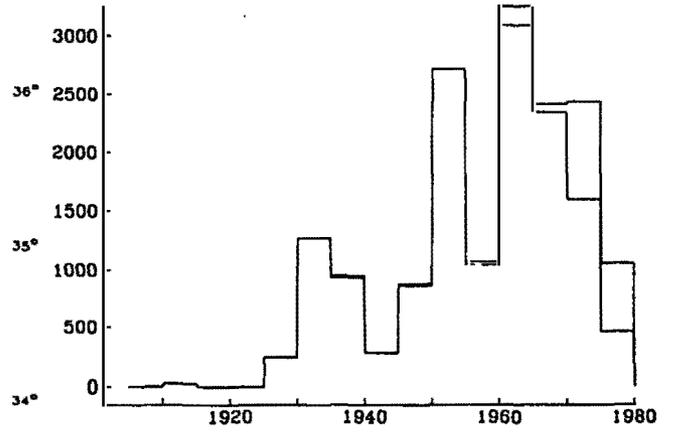
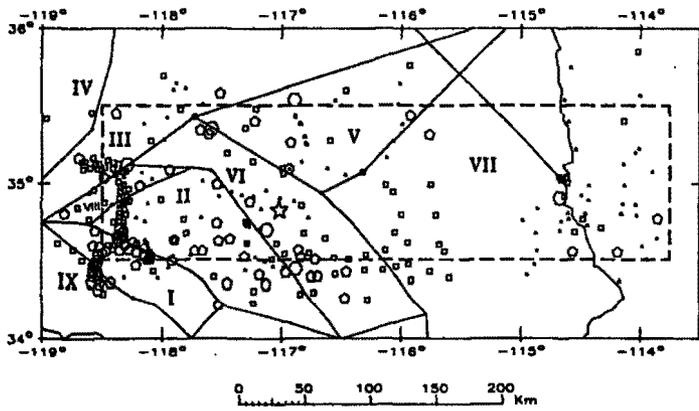


Region 1. San Diego

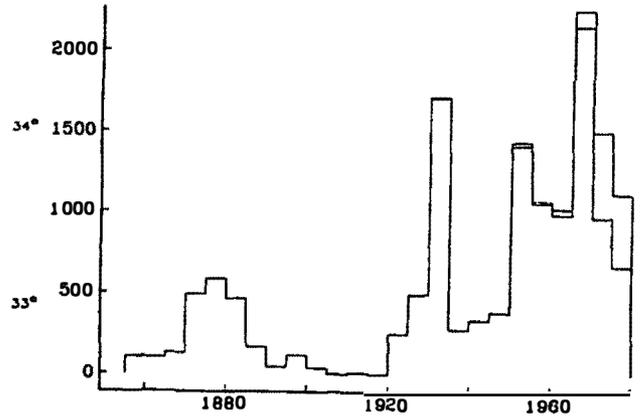
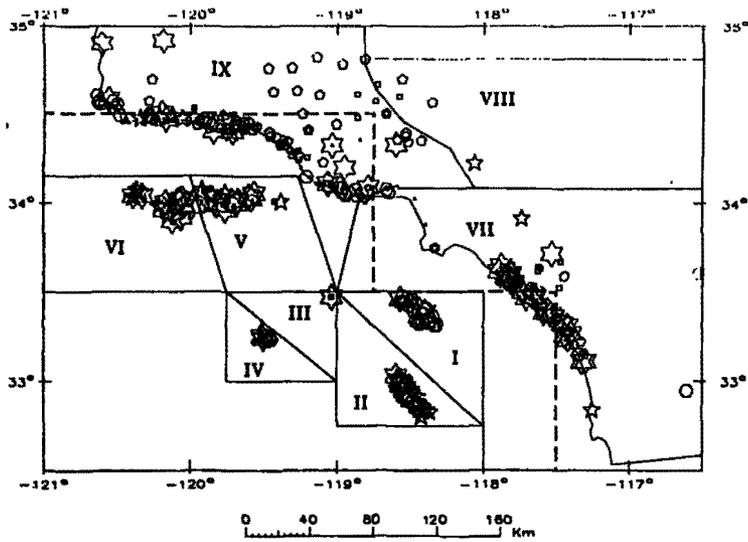


Region 2. San Bernardino

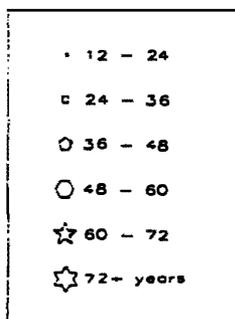


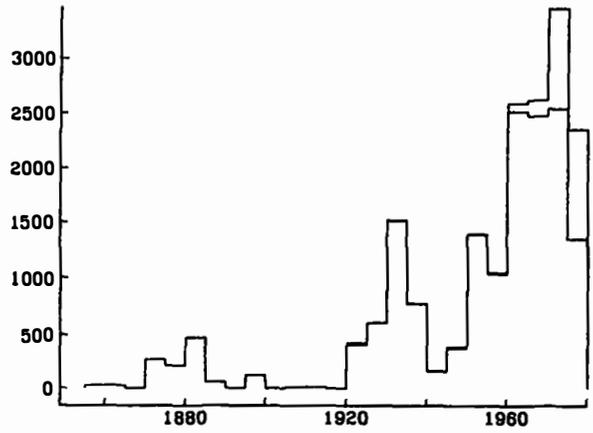
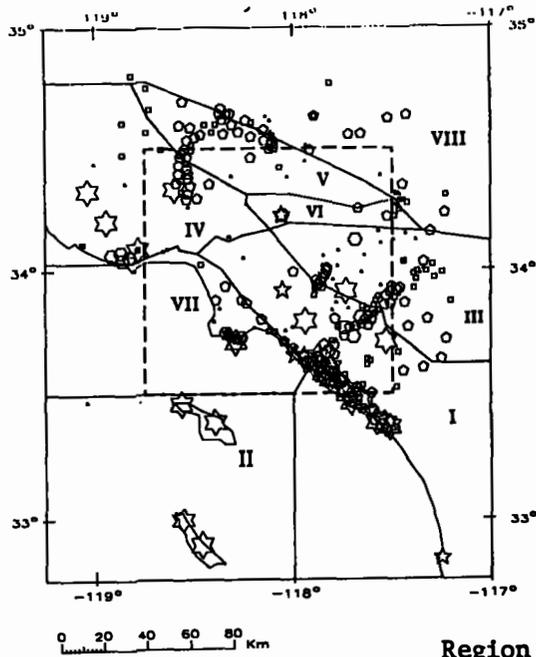


Region 3. Barstow

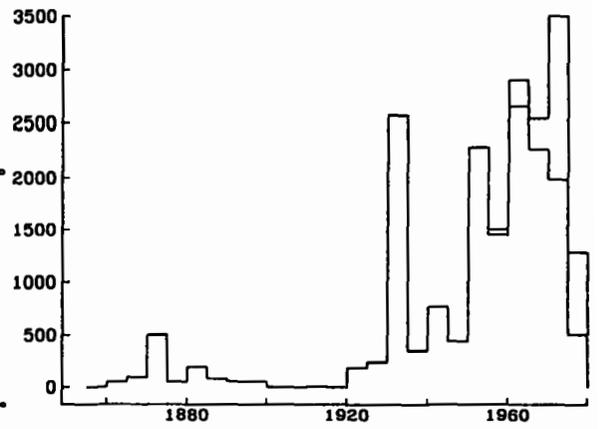
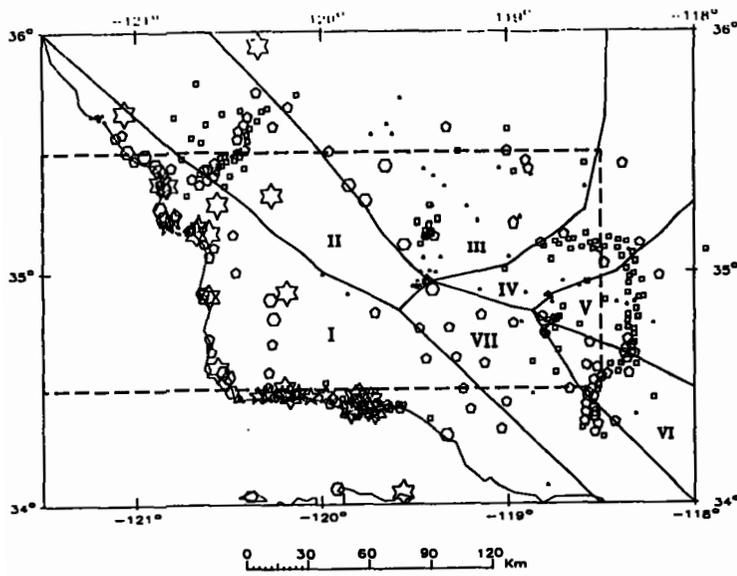


Region 4. Channel Islands

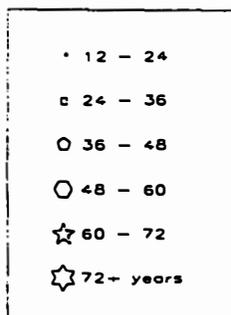


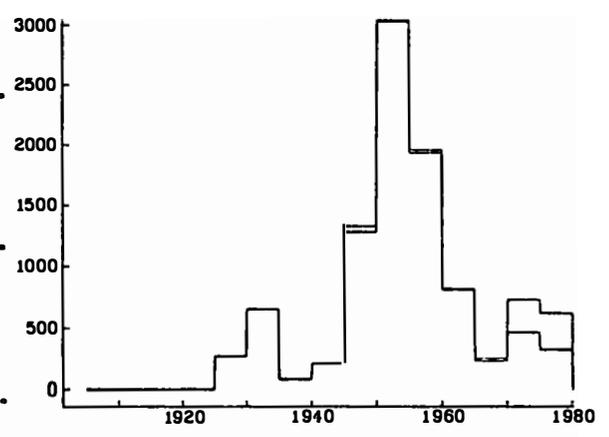
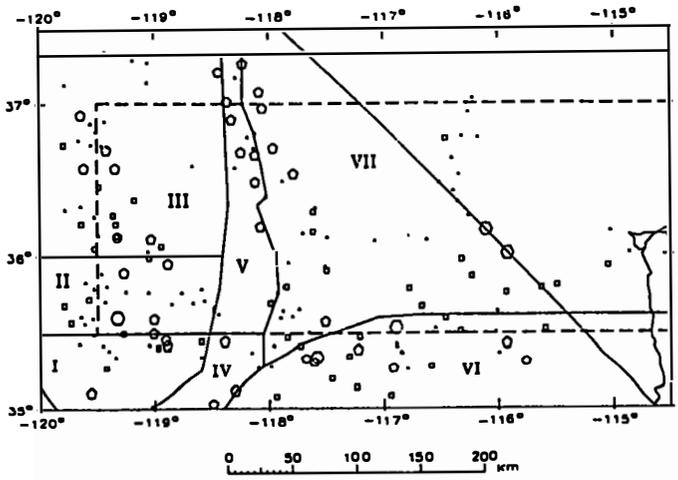


Region 5. Los Angeles

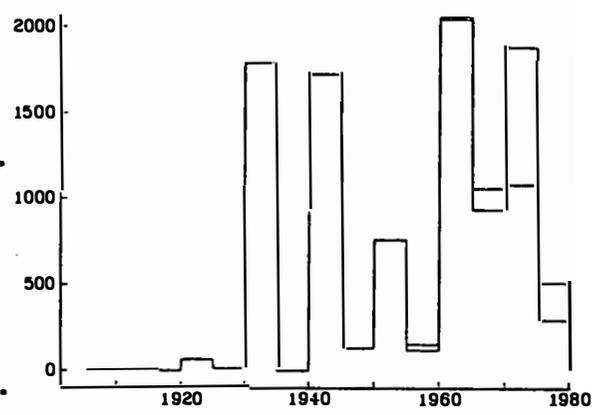
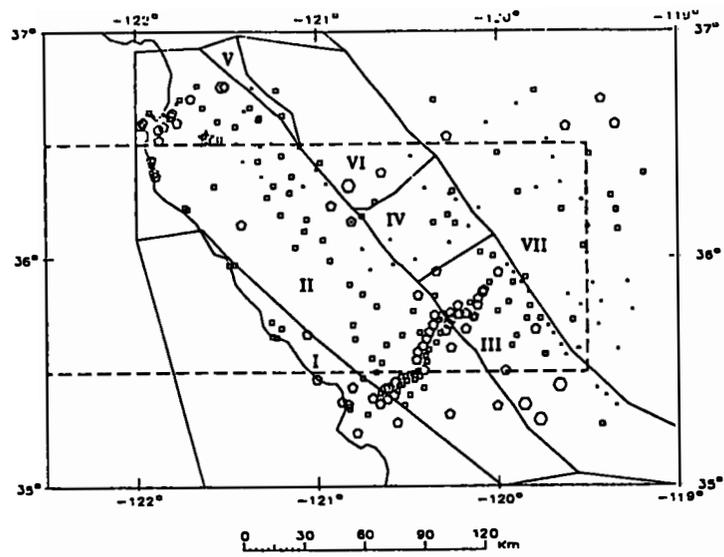


Region 6: Bakersfield

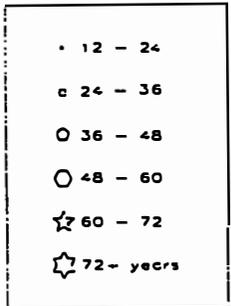


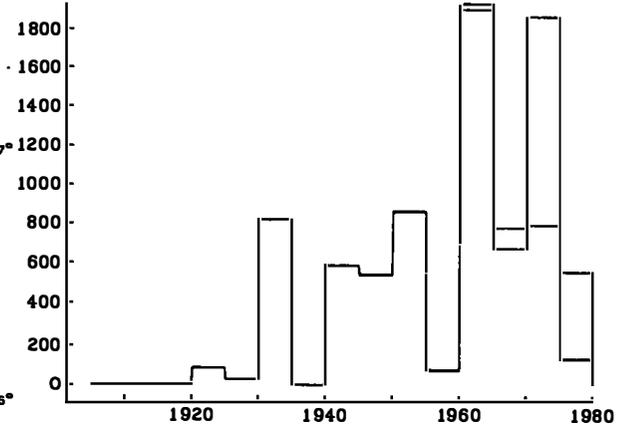
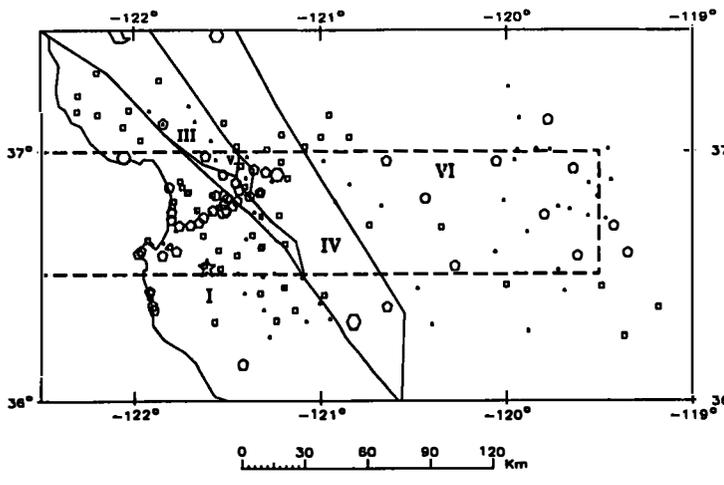


Region 7. Sierra Nevada

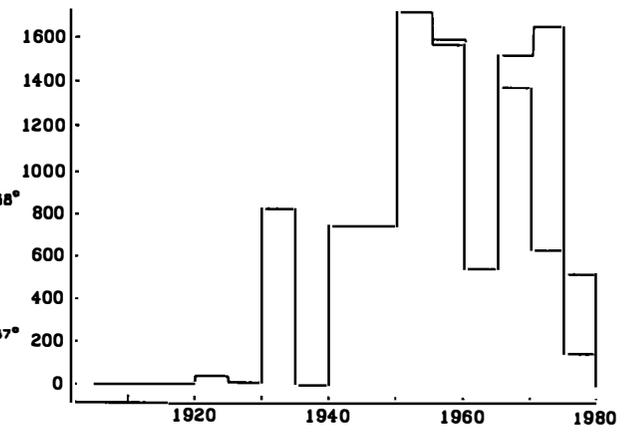
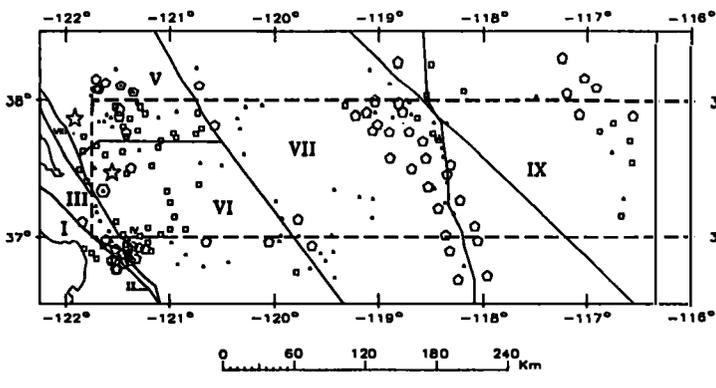


Region 8. Parkfield

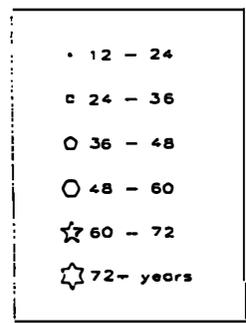


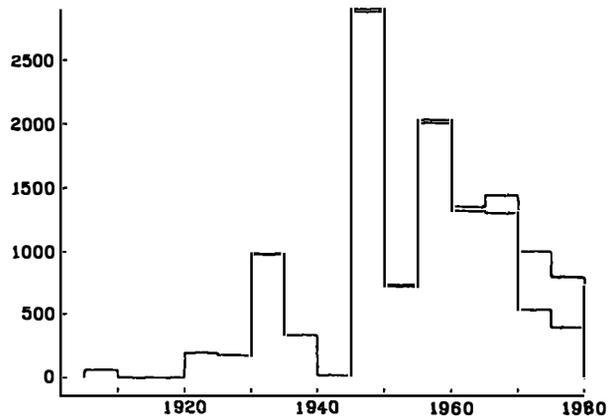
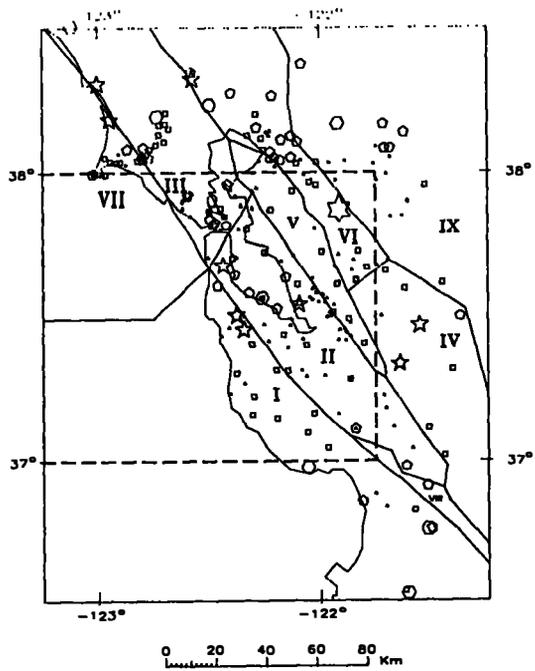


Region 9. Monterey

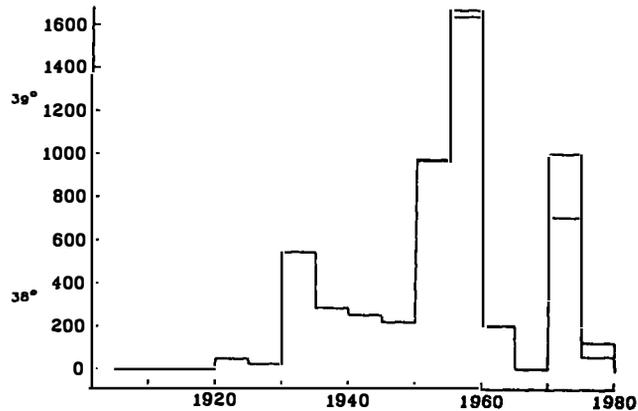
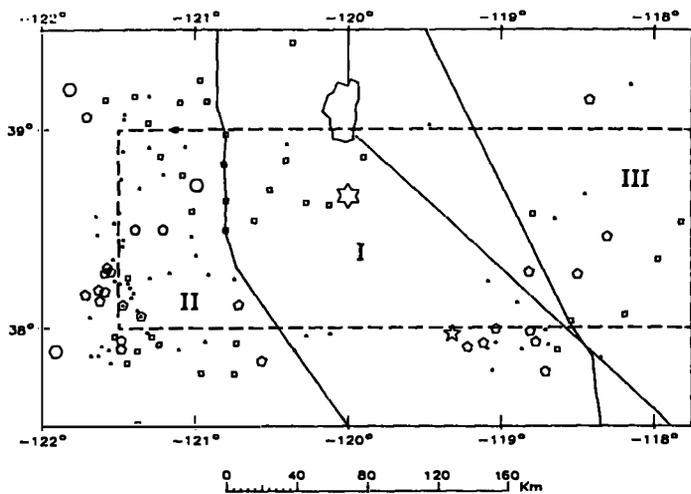


Region 10. Yosemite

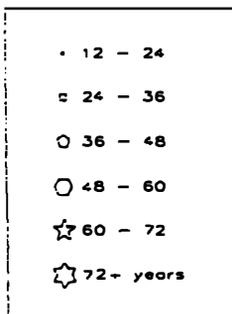


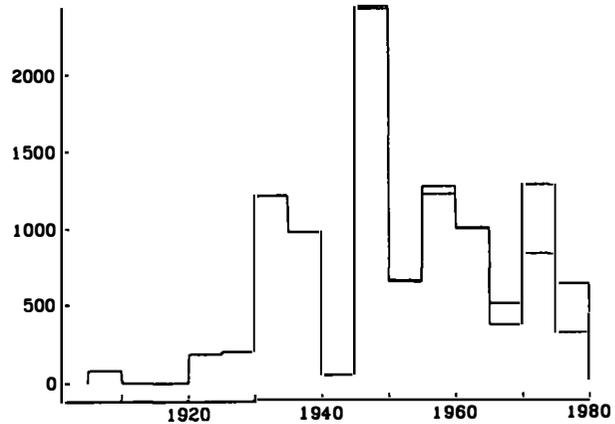
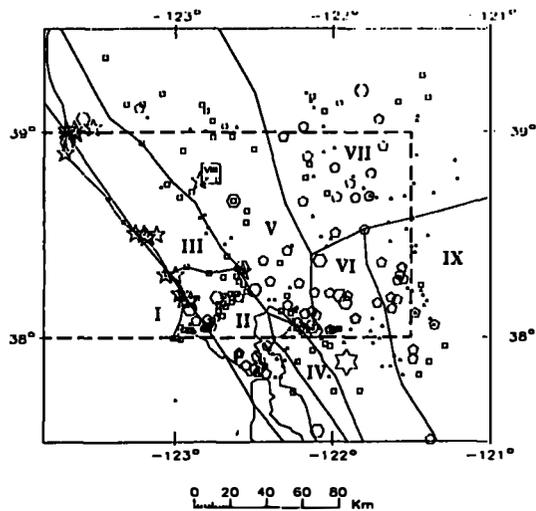


Region 11. San Francisco

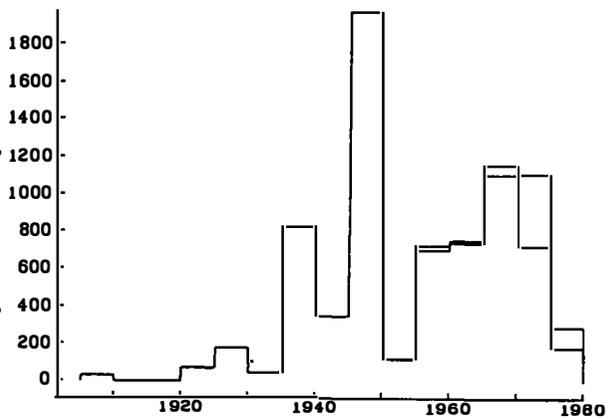
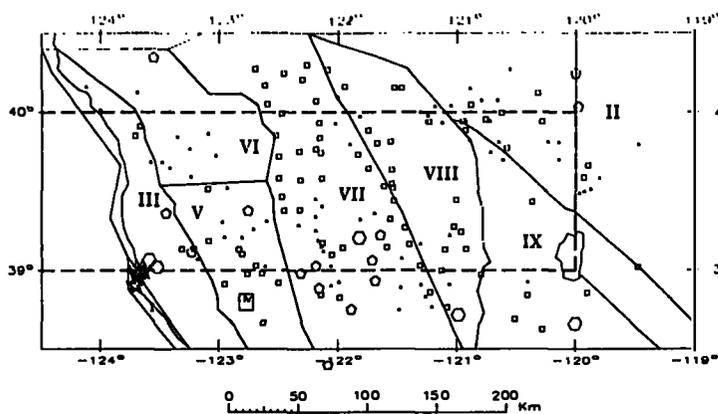


Region 12. Sacramento

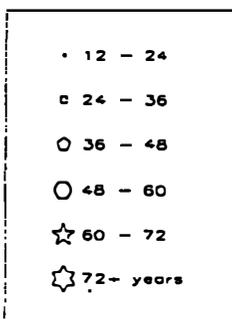


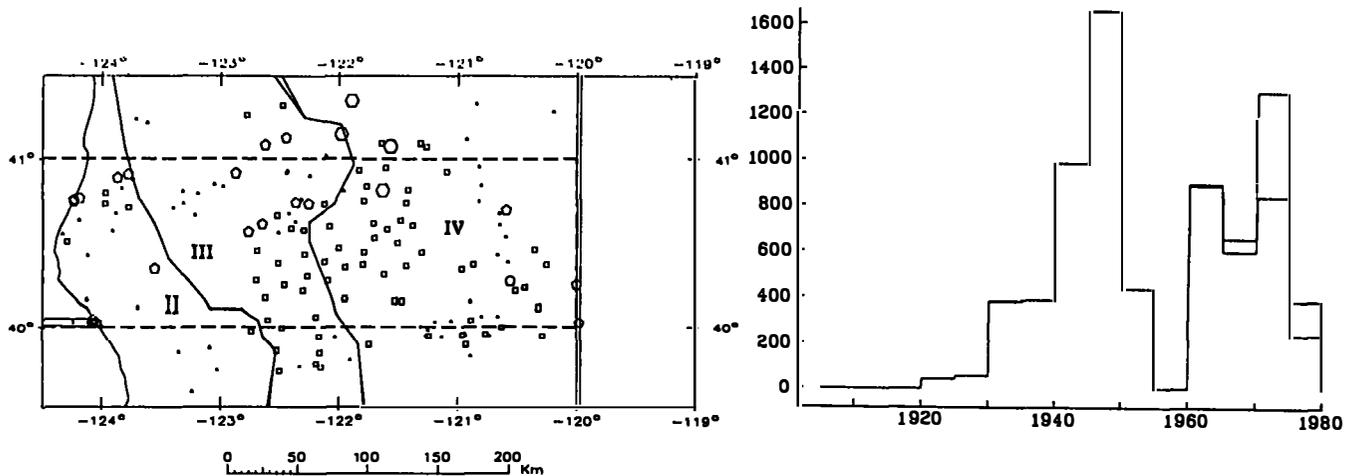


Region 13. Santa Rosa

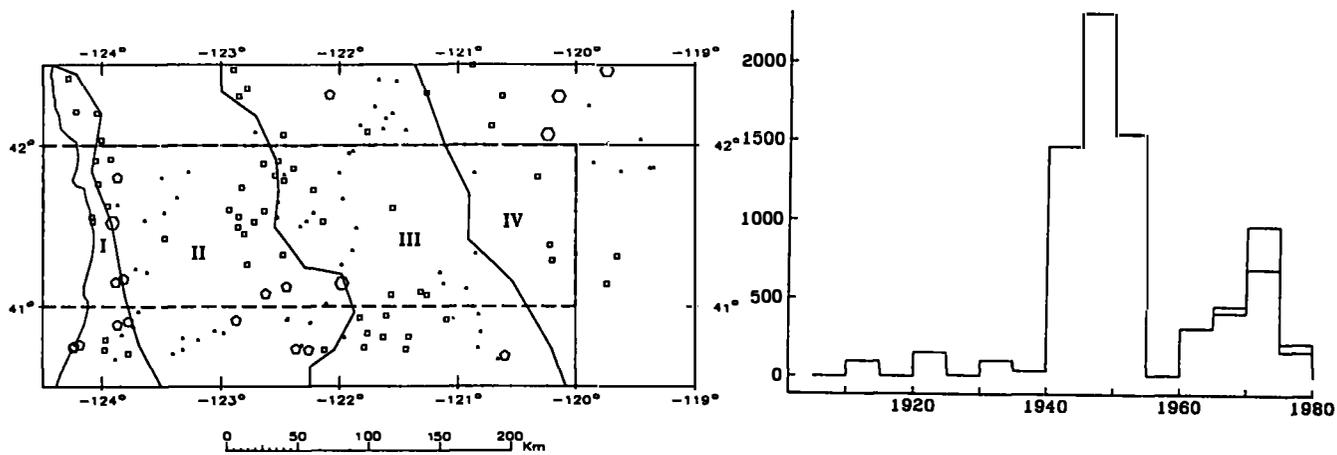


Region 14. Ukiah

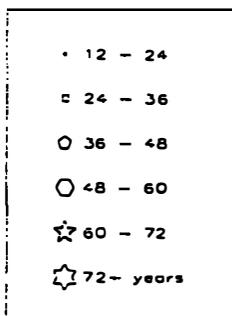




Region 15. Redding



Region 16. Alturas



APPENDIX C.--PARAMETER VALUES

This appendix contains four tables that document the parameter values for REDEAM regions in California and Nevada. For the regions in Alaska and Hawaii certain parameter values are found in other appendices because the representation of episodic motion differs for these latter two States. In particular, parameters for the Anchorage, Alaska, region are documented in appendix E while parameters pertaining to episodic motion for the Hilo, Hawaii, region are documented in appendix F. Secular motion parameters for the Hilo region are, however, documented in this appendix.

The tabulated parameters enable one to compute $(\phi_M(t), \lambda_M(t))$ for a point M at time t according to eq. (3.1) if the coordinates $(\phi_M(t_0), \lambda_M(t_0))$ are given where $t_0 =$ January 1, 1950. Recall that for our application of the REDEAM models (chapter 5), we compute only corrections to observations, and thus the choice of t_0 could be and was rather arbitrary. Moreover for our application, the NAD 27 positions provide sufficiently accurate values for coordinates at time t_0 .

Table C.1 contains the values for the reference coordinates (latitude and longitude) of the various regions. These parameters are referred to as $\bar{\phi}$ and $\bar{\lambda}$ in eq. (3.1). Table C.2 contains the values for the estimated secular motion parameters, and table C.3 contains values for episodic motion parameters. Finally, table C.4 contains the vertices that define the boundaries for the various regions and their districts. These vertices are represented as coordinate pairs (latitude and longitude). A simple closed polygon is formed by connecting consecutive vertices with edges and by similarly connecting the first and last vertices. For operational purposes, a geodetic mark is considered to be located in a region or district if it is located inside the corresponding polygonal loop. (A simple closed polygon bounds two disjoint areas on the Earth's surface. A point is said to be inside this polygon if it is located in the bounded area that does not contain the North Pole.)

Table C.1.--Reference coordinates

Region	Latitude (deg-min-sec)	Longitude (deg-min-sec)
San Diego (CA)	33 16 44.384	115 30 01.121
San Bernardino (CA)	34 11 45.849	115 40 15.762
Barstow (CA)	35 04 15.283	116 19 21.480
Channel Islands (CA)	34 19 32.41086	119 02 17.689
Los Angeles (CA)	34 13 25.51664	118 03 38.66333
Bakersfield (CA)	34 45 14.69362	119 28 27.34339
Sierra Nevada (CA)	37 00 38.382	118 21 37.546
Parkfield (CA)	35 25 49.913	119 26 55.226
Monterey (CA)	36 46 39.990	121 28 13.088
Yosemite (CA)	37 09 25.683	116 40 03.267
San Francisco (CA)	37 28 38.609	121 33 17.529
Sacramento (CA)	38 26 50.125	121 23 48.071
Santa Rosa (CA)	38 40 10.00120	122 37 56.60606
Ukiah (CA)	39 12 21.300	121 49 10.350
Redding (CA)	40 45 05.030	124 13 51.856
Alturas (CA)	41 31 42.127	122 08 22.245
Fallon (NV)	39 34 59.23375	118 14 04.07787
Anchorage (AK)	NA	NA
Hilo (HI)	19 43 10.832	155 05 57.572

Table C.2.--Parameter values for secular motion

District	$\overset{a}{\phi}^*$ 10^{-8} rad/yr	$\overset{a}{\phi\phi}$ $10^{-6}/\text{yr}$	$\overset{a}{\phi\lambda}$ $10^{-6}/\text{yr}$	$\overset{a}{\lambda}^*$ 10^{-8} rad/yr	$\overset{a}{\lambda\phi}$ $10^{-6}/\text{yr}$	$\overset{a}{\lambda\lambda}$ $10^{-6}/\text{yr}$
SAN DIEGO REGION, CALIFORNIA						
I	0.0000*	-0.1120 (0.027)	0.1020 (0.105)	0.0000*	-0.0066 (0.149)	0.0412 (0.022)
II	0.1870 (0.036)	-0.2330 (0.039)	-0.0170 (0.112)	-0.2530 (0.047)	-0.5060 (0.153)	0.6260 (0.063)
III	-0.0514 (0.016)	-0.2430 (0.049)	0.1240 (0.110)	-0.0518 (0.018)	-0.1690 (0.157)	0.2120 (0.054)
IV	-0.6250 (0.069)	-1.0200 (0.093)	0.5210 (0.120)	0.4280 (0.084)	0.4670 (0.183)	0.0183 (0.093)
V	0.2770 (0.035)	-0.1630 (0.028)	-0.0079 (0.107)	0.0141 (0.043)	-0.2480 (0.149)	0.1420 (0.034)
VI	0.3200 (0.061)	0.0000*	0.0000*	0.2550 (0.086)	0.0000*	0.0000*
VII	0.1560 (0.027)	-0.2310 (0.039)	0.0590 (0.107)	0.0368 (0.028)	-0.1700 (0.152)	0.1200 (0.031)
VIII	0.0675 (0.041)	-0.2000 (0.071)	0.0000*	0.0160 (0.025)	-0.1970 (0.156)	0.0000*
IX	0.3210 (0.072)	-0.1550 (0.037)	-0.0192 (0.108)	0.2390 (0.060)	-0.0759 (0.151)	0.0148 (0.030)
X	0.3270 (0.042)	-0.1430 (0.025)	-0.0223 (0.104)	0.5000 (0.053)	0.0778 (0.146)	-0.0991 (0.022)
SAN BERNARDINO REGION, CALIFORNIA						
I	0.4960 (0.243)	0.0693 (0.114)	0.0000*	0.0592 (0.115)	-0.2140 (0.118)	0.0000*
II	-0.0008 (0.077)	-0.1690 (0.024)	0.0865 (0.056)	0.1080 (0.084)	-0.1480 (0.081)	-0.0057 (0.023)
III	-0.0060 (0.079)	-0.1810 (0.041)	0.0682 (0.057)	-0.1360 (0.080)	-0.2380 (0.090)	0.0702 (0.028)
IV	-0.2250 (0.176)	-0.1590 (0.100)	0.1470 (0.076)	-0.2820 (0.198)	-0.3170 (0.132)	0.1170 (0.066)
V	-0.1470 (0.039)	-0.2040 (0.034)	0.1630 (0.064)	-0.1570 (0.046)	-0.1950 (0.090)	0.1170 (0.033)
VI	-0.1990 (0.090)	-0.1340 (0.043)	0.1350 (0.061)	-0.1610 (0.103)	-0.2240 (0.097)	-0.0365 (0.038)
VII	-0.0914 (0.041)	-0.0950 (0.034)	0.0857 (0.060)	-0.0283 (0.049)	-0.1340 (0.093)	0.0280 (0.040)
VIII	0.0000*	-0.0282 (0.046)	-0.0356 (0.068)	0.0000*	0.0535 (0.095)	-0.0606 (0.053)
IX	0.1210 (0.170)	-0.7050 (0.300)	0.0000*	-0.0092 (0.082)	-0.7370 (0.310)	0.0000*
X	0.1930 (0.102)	0.0000*	0.0000*	0.0750 (0.059)	0.0000*	0.0000*
XI	-0.0130 (0.059)	0.0446 (0.057)	-0.0579 (0.070)	-0.0296 (0.071)	0.1290 (0.102)	-0.1240 (0.054)
BARSTOW REGION, CALIFORNIA						
I	-0.1780 (0.092)	-0.1890 (0.029)	0.0723 (0.103)	-0.5300 (0.101)	-0.3870 (0.153)	0.1220 (0.019)
II	-0.1570 (0.067)	-0.1610 (0.022)	0.0724 (0.103)	-0.1050 (0.073)	-0.1850 (0.151)	0.0442 (0.016)
III	0.0790 (0.080)	-0.1980 (0.047)	0.0070 (0.102)	0.0547 (0.121)	-0.1800 (0.156)	0.0073 (0.035)
IV	-1.2300 (0.800)	1.3400 (0.455)	0.0810 (0.196)	2.7200 (1.380)	-3.2000 (0.386)	-0.2120 (0.345)
V	0.0000*	-0.1630 (0.054)	0.0416 (0.104)	0.0000*	-0.0164 (0.155)	-0.0472 (0.039)
VI	0.0888 (0.061)	-0.0523 (0.028)	-0.0247 (0.103)	0.0192 (0.070)	-0.1000 (0.151)	0.0100 (0.039)
VII	-0.0431 (0.035)	-0.2020 (0.062)	-0.0057 (0.113)	-0.0301 (0.047)	-0.0229 (0.163)	-0.1750 (0.070)
VIII	0.1480 (0.095)	-0.2100 (0.047)	-0.0158 (0.104)	-0.2910 (0.108)	-0.1750 (0.157)	0.1020 (0.028)
IX	-0.7770 (0.232)	-0.4070 (0.047)	0.1580 (0.115)	0.4780 (0.239)	0.0131 (0.168)	-0.0240 (0.047)
CHANNEL ISLANDS REGION, CALIFORNIA						
I	0.4610 (0.656)	0.1240 (0.411)	0.1510 (0.178)	0.2060 (0.588)	-0.1810 (0.268)	0.3910 (0.402)
II	-0.1180 (2.290)	-0.2220 (0.993)	0.3280 (0.130)	-0.5150 (0.861)	-0.3330 (0.175)	0.1830 (0.995)
III	0.1950 (0.107)	0.0000*	0.0000*	0.1280 (0.160)	0.0000*	0.0000*
IV	0.1780 (0.795)	-0.1590 (0.400)	-0.1790 (0.206)	0.6960 (0.607)	0.5930 (0.274)	0.4740 (0.409)
V	-0.2710 (0.106)	-0.2370 (0.143)	0.1600 (0.083)	0.0791 (0.156)	-0.0339 (0.173)	-0.1440 (0.118)
VI	0.4910 (0.264)	0.1640 (0.259)	-0.1010 (0.102)	-0.5990 (0.435)	-1.1400 (0.203)	-0.1420 (0.224)
VII	-0.1820 (0.102)	-0.1880 (0.123)	-0.1810 (0.085)	0.0718 (0.133)	0.2740 (0.123)	-0.1070 (0.121)
VIII	0.1920 (0.193)	-0.3800 (0.212)	0.0584 (0.138)	0.3490 (0.254)	-0.2440 (0.238)	0.2100 (0.211)
IX	0.0000*	-0.1590 (0.026)	-0.0006 (0.042)	0.0000*	0.0789 (0.065)	-0.0742 (0.025)

*Parameter constrained to this value.

Table C.2.--Continued

LOS ANGELES REGION, CALIFORNIA

I	0.0974 (0.029)	-0.0722 (0.023)	0.0995 (0.048)	-0.0507 (0.025)	-0.1310 (0.068)	0.0492 (0.022)
II	0.0611 (1.070)	0.0843 (0.616)	0.2140 (0.341)	0.0468 (0.363)	0.0356 (0.191)	0.3060 (0.375)
III	0.0328 (0.030)	-0.0884 (0.034)	0.0448 (0.048)	-0.0480 (0.026)	-0.1470 (0.076)	0.0474 (0.021)
IV	-0.0387 (0.041)	-0.3060 (0.034)	0.0757 (0.052)	0.1020 (0.039)	0.0887 (0.079)	-0.1330 (0.027)
V	-0.0206 (0.035)	-0.1620 (0.033)	0.0236 (0.049)	0.0576 (0.030)	-0.2610 (0.079)	0.0751 (0.019)
VI	0.0000*	-0.5400 (0.117)	0.0185 (0.056)	0.0000*	0.0148 (0.182)	0.0408 (0.032)
VII	-0.0722 (0.133)	-0.2160 (0.141)	0.1020 (0.089)	0.0103 (0.192)	-0.0300 (0.215)	-0.1030 (0.120)
VIII	-0.1020 (0.042)	-0.0297 (0.044)	-0.0294 (0.051)	-0.0713 (0.042)	-0.0435 (0.090)	-0.0036 (0.029)

BAKERSFIELD REGION, CALIFORNIA

I	0.0535 (0.013)	-0.1190 (0.032)	0.1640 (0.041)	-0.0241 (0.015)	-0.3240 (0.065)	0.0492 (0.026)
II	-0.0374 (0.019)	-0.1750 (0.039)	0.2380 (0.044)	0.0286 (0.024)	-0.3840 (0.073)	0.0601 (0.029)
III	0.0263 (0.022)	-0.3030 (0.031)	0.1640 (0.046)	0.0375 (0.025)	-0.4310 (0.069)	0.0016 (0.027)
IV	0.0130 (0.028)	-0.2630 (0.035)	0.1840 (0.050)	0.0411 (0.030)	-0.5070 (0.079)	0.0097 (0.025)
V	-0.0900 (0.029)	-0.3830 (0.032)	0.1060 (0.048)	0.0123 (0.034)	-0.2800 (0.077)	0.0682 (0.021)
VI	-0.1160 (0.037)	-0.2150 (0.035)	0.0794 (0.052)	0.0707 (0.043)	-0.3600 (0.074)	0.1110 (0.031)
VII	0.0000*	-0.2810 (0.027)	0.1780 (0.047)	0.0000*	-0.2990 (0.073)	0.0465 (0.026)

SIERRA NEVADA REGION, CALIFORNIA

I	-0.7580 (0.648)	-0.4350 (0.239)	0.0754 (0.099)	-1.6901 (0.541)	-0.8380 (0.222)	-0.0856 (0.085)
II	0.1160 (0.189)	-0.0933 (0.080)	0.1420 (0.086)	0.2660 (0.220)	-0.1180 (0.141)	-0.0483 (0.058)
III	0.0479 (0.087)	-0.0927 (0.055)	0.1700 (0.082)	-0.0198 (0.071)	-0.2010 (0.127)	-0.0013 (0.044)
IV	-0.0506 (0.579)	-0.1830 (0.209)	-0.4040 (0.179)	-0.4310 (0.419)	-0.3590 (0.175)	0.2220 (0.227)
V	0.0000*	-0.1830 (0.047)	0.0822 (0.100)	0.0000*	-0.2150 (0.120)	0.1410 (0.089)
VI	0.1590 (0.243)	-0.1160 (0.076)	0.0980 (0.076)	0.2890 (0.290)	-0.0358 (0.131)	-0.0284 (0.027)
VII	0.0563 (0.033)	-0.1920 (0.042)	0.1280 (0.075)	0.1120 (0.045)	-0.1040 (0.118)	-0.0297 (0.026)

PARKFIELD REGION, CALIFORNIA

I	0.2190 (0.122)	-0.1550 (0.077)	0.1130 (0.112)	0.1320 (0.139)	-0.0402 (0.167)	-0.0495 (0.064)
II	0.1450 (0.053)	-0.1100 (0.028)	0.1470 (0.100)	0.0590 (0.062)	-0.1620 (0.150)	-0.0162 (0.029)
III	0.0000*	-0.2500 (0.037)	0.2500 (0.105)	0.0000*	-0.4210 (0.154)	0.0796 (0.043)
IV	0.5840 (0.137)	-0.4260 (0.088)	-0.0174 (0.111)	0.0814 (0.190)	-0.5590 (0.191)	0.0607 (0.071)
V	-1.2500 (0.228)	0.6380 (0.170)	0.0000*	-0.4990 (0.319)	0.0000*	0.0000*
VI	-0.0581 (0.096)	0.0119 (0.074)	0.0041 (0.123)	-0.3930 (0.138)	-0.1210 (0.183)	0.0061 (0.101)
VII	0.0444 (0.072)	-0.1980 (0.061)	0.1460 (0.111)	-0.0505 (0.065)	-0.1940 (0.163)	-0.1320 (0.080)

MONTEREY REGION, CALIFORNIA

I	0.0000*	-0.1820 (0.021)	0.2070 (0.089)	0.0000*	-0.2900 (0.136)	-0.0315 (0.021)
II	-0.0901 (0.005)	-0.4410 (0.043)	0.5890 (0.098)	-0.1480 (0.006)	-0.4110 (0.145)	0.2820 (0.053)
III	-0.1320 (0.021)	-0.2540 (0.076)	0.3630 (0.101)	-0.2760 (0.034)	-0.2150 (0.173)	0.1810 (0.054)
IV	-0.3210 (0.015)	-0.2020 (0.031)	0.3030 (0.093)	-0.2770 (0.016)	-0.3280 (0.137)	0.0560 (0.036)
V	-0.2050 (0.042)	-0.8390 (0.097)	0.2380 (0.252)	-0.5690 (0.052)	0.2470 (0.180)	-1.8100 (0.292)
VI	-0.4260 (0.066)	-0.0814 (0.073)	0.2300 (0.097)	-0.3120 (0.092)	-0.2470 (0.154)	-0.0269 (0.069)

YOSEMITE REGION, CALIFORNIA

I	1.2300 (1.050)	-0.0761 (0.119)	0.0015 (0.143)	-0.1300 (0.880)	0.0546 (0.249)	0.0062 (0.089)
II	-1.1900 (1.512)	-0.3570 (0.048)	0.2540 (0.100)	-5.0800 (0.579)	-0.1160 (0.153)	0.5640 (0.061)
III	0.9080 (0.606)	-0.1970 (0.091)	0.0151 (0.106)	-1.6600 (0.501)	-0.1460 (0.173)	0.1490 (0.051)
IV	-3.7000 (1.900)	-0.9360 (0.086)	0.5040 (0.204)	-0.4020 (0.214)	0.0227 (0.144)	0.0000*

*Parameter constrained to this value.

Table C.2.--Continued

V	-1.2700 (0.403)	0.0295 (0.064)	0.2330 (0.089)	0.0006 (0.425)	-0.1120 (0.145)	-0.0601 (0.049)
VI	-0.4120 (0.257)	-0.1830 (0.024)	0.1480 (0.085)	-0.2730 (0.243)	-0.0866 (0.135)	-0.0251 (0.024)
VII	0.1700 (0.200)	-0.0432 (0.027)	0.0401 (0.091)	-0.4530 (0.199)	-0.0468 (0.140)	0.0001 (0.036)
VIII	0.9120 (0.791)	0.0000*	0.0000*	-0.5190 (0.238)	0.0000*	0.0000*
IX	0.0000*	-0.0874 (0.036)	0.1020 (0.108)	0.0000*	-0.0882 (0.144)	-0.1560 (0.062)
SAN FRANCISCO REGION, CALIFORNIA						
I	0.2880 (0.055)	-0.1730 (0.041)	-0.0564 (0.064)	0.0595 (0.055)	-0.0283 (0.104)	0.0036 (0.035)
II	0.0565 (0.033)	-0.2370 (0.029)	0.0832 (0.064)	-0.0166 (0.034)	-0.0376 (0.097)	0.0760 (0.035)
III	0.1290 (0.061)	-0.4920 (0.047)	0.1680 (0.068)	-0.1720 (0.073)	0.0103 (0.112)	0.1220 (0.044)
IV	0.0000*	-0.1110 (0.053)	-0.0762 (0.079)	0.0000*	0.0992 (0.104)	-0.1200 (0.049)
V	-0.0386 (0.039)	-0.1710 (0.048)	0.0864 (0.085)	-0.0706 (0.039)	-0.0173 (0.107)	0.0832 (0.073)
VI	-0.0479 (0.027)	-0.0913 (0.039)	0.0027 (0.067)	-0.0524 (0.025)	-0.0906 (0.103)	0.1360 (0.044)
VII	-0.2120 (0.122)	-0.1080 (0.084)	0.1500 (0.079)	-0.6130 (0.182)	-0.1520 (0.126)	0.3960 (0.086)
VIII	0.3490 (0.029)	0.0000*	0.0000*	0.0701 (0.096)	0.0000*	0.0000*
IX	-0.0833 (0.044)	-0.1520 (0.097)	0.0797 (0.106)	-0.0918 (0.049)	0.2470 (0.159)	-0.2740 (0.112)
SACRAMENTO REGION, CALIFORNIA						
I	-0.0925 (0.069)	0.0016 (0.062)	0.4600 (0.408)	-0.2940 (0.105)	-0.7980 (0.663)	-0.0791 (0.076)
II	0.0000*	-0.0534 (0.077)	0.5010 (0.403)	0.0000*	-0.7640 (0.657)	0.1600 (0.089)
III	0.7190 (0.374)	-0.0468 (0.036)	0.6280 (0.416)	-0.2120 (0.385)	-0.8650 (0.667)	-0.0382 (0.050)
SANTA ROSA REGION, CALIFORNIA						
I	-0.0368 (0.062)	-0.3330 (0.051)	0.3050 (0.097)	-0.0774 (0.081)	-0.4110 (0.151)	0.2540 (0.061)
II	-0.0828 (0.022)	-0.3490 (0.032)	0.3830 (0.090)	0.0422 (0.031)	-0.3430 (0.143)	0.1140 (0.039)
III	0.0436 (0.432)	-0.3140 (0.049)	0.2300 (0.099)	-0.1640 (0.057)	-0.5880 (0.157)	0.3410 (0.079)
IV	0.0697 (0.127)	-0.2520 (0.100)	0.4010 (0.120)	-0.0135 (0.148)	-0.1980 (0.178)	-0.2830 (0.128)
V	0.0000*	-0.2070 (0.027)	0.2950 (0.090)	0.0000*	-0.3690 (0.145)	-0.0115 (0.032)
VI	-0.0930 (0.071)	-0.2670 (0.062)	0.2310 (0.097)	0.1380 (0.082)	-0.2590 (0.155)	0.0044 (0.064)
VII	0.0036 (0.078)	-0.1720 (0.065)	0.2810 (0.103)	0.1040 (0.100)	-0.4090 (0.159)	-0.1050 (0.068)
VIII	0.5240 (0.076)	-2.8900 (0.280)	0.2290 (0.358)	0.8120 (0.092)	1.3600 (0.348)	-4.2000 (0.415)
IX	0.0083 (0.134)	-0.2680 (0.074)	0.2760 (0.121)	0.1520 (0.186)	-0.2850 (0.162)	-0.0337 (0.121)
UKIAH REGION, CALIFORNIA						
I	-2.5200 (1.820)	-0.2210 (0.441)	1.2000 (0.597)	-3.5400 (1.430)	-1.2600 (0.971)	0.9320 (0.448)
II	-0.0137 (0.194)	0.0161 (0.116)	0.0825 (0.095)	-0.0244 (0.293)	-0.2530 (0.199)	-0.0309 (0.093)
III	-1.3700 (0.441)	-0.4720 (0.219)	0.7740 (0.186)	1.3100 (0.664)	-0.1450 (0.275)	-0.4980 (0.256)
IV	-1.9800 (0.140)	-2.9000*	0.2300*	8.2000 (0.134)	1.3600*	-4.2000*
V	-0.1400 (0.102)	-0.2610 (0.041)	0.2870 (0.091)	-0.1600 (0.116)	-0.1780 (0.151)	0.0587 (0.048)
VI	-0.1160 (0.313)	0.1190 (0.186)	0.1970 (0.147)	-0.6280 (0.467)	-0.4540 (0.251)	0.3100 (0.190)
VII	0.0000*	0.0128 (0.052)	0.1470 (0.081)	0.0000*	-0.1630 (0.132)	0.0399 (0.057)
VIII	-0.0701 (0.046)	0.0853 (0.063)	0.1000 (0.088)	0.0182 (0.058)	-0.2230 (0.134)	-0.0034 (0.076)
IX	0.1160 (0.255)	0.0729 (0.105)	0.1850 (0.128)	0.0367 (0.307)	-0.2340 (0.169)	0.0485 (0.140)
REDDING REGION, CALIFORNIA						
I	0.2890 (3.100)	-0.4650 (0.382)	0.5970 (0.249)	0.0000*	0.5010 (0.413)	-0.0284 (0.387)
II	0.0000*	0.2420 (0.145)	-0.0806 (0.237)	0.0000*	0.2740 (0.393)	0.3330 (0.157)
III	-0.1070 (0.092)	0.2120 (0.118)	-0.1360 (0.234)	-0.1230 (0.143)	0.1970 (0.402)	0.2750 (0.117)
IV	0.0388 (0.199)	0.2050 (0.119)	-0.0944 (0.236)	-0.4900 (0.261)	0.1550 (0.404)	0.1770 (0.116)

*Parameter constrained to this value.

Table C.2.--Continued

ALTURAS REGION, CALIFORNIA

I	-0.0631 (0.412)	0.0567 (0.051)	0.0135 (0.105)	0.0519 (0.513)	0.0182 (0.062)	0.0272 (0.132)
II	-0.0231 (0.033)	0.0312 (0.040)	-0.0321 (0.043)	0.0116 (0.045)	-0.0199 (0.061)	0.0265 (0.057)
III	0.0000*	-0.0567 (0.035)	0.0026 (0.039)	0.0000*	-0.0238 (0.055)	-0.0221 (0.060)
IV	0.0427 (0.240)	-0.0229 (0.072)	0.0086 (0.068)	-0.0568 (0.323)	0.0124 (0.108)	-0.0397 (0.087)

FALLON REGION, NEVADA

I	0.0000*	-0.2010 (0.115)	0.1030 (0.210)	0.0000*	-0.0995 (0.325)	-0.0987 (0.113)
---	---------	-----------------	----------------	---------	-----------------	-----------------

HILO REGION, HAWAII

I	0.1840 (0.089)	0.6650 (0.204)	0.3540 (1.880)	0.2320 (0.103)	0.6160 (2.090)	0.8960 (0.167)
II	-2.0900 (0.599)	0.0000*	0.0000*	-0.3950 (1.330)	0.0000*	0.0000*
III	-0.0928 (0.375)	0.0000*	0.0000*	0.0941 (1.260)	0.0000*	0.0000*
IV	-0.7820 (0.375)	0.1630 (0.630)	1.9700 (1.890)	-1.3200 (0.501)	-1.6500 (2.240)	0.4380 (0.300)
V	0.0000*	0.2760 (0.165)	0.1940 (1.870)	0.0000*	-0.2560 (2.080)	0.1650 (0.177)

*Parameter constrained to this value.

Table C.3.--Parameters for modeled earthquakes in California and Nevada

Earthquake	Fault Center		Fault Orientation		Fault Size		Slip (standard error)		
	Year/Day (Magnitude)	Latitude (N)	Longitude (W)	Strike	Dip	Depth Range, km	Length, km	Strike, m	Dip, m
SAN DIEGO REGION									
1940/138(M7.1)	32° 34' 31"	115 14' 44"	N 37° W	75° NE	0-10	22		-2.62(.31)	0.03(.47)
1940/138(M7.1)	32° 41' 00"	115 20' 30"	N 37° W	75° NE	0-10	8		-4.33(.16)	2.58(.16)
1940/138(M7.1)	32° 44' 38"	115 23' 35"	N 37° W	75° NE	0-10	8		-1.34(.12)	-0.99(.20)
1940/138(M7.1)	32° 50' 31"	115 28' 59"	N 37° W	75° NE	0-10	20		-1.03(.06)	0.20(.09)
1940/138(M7.1)	32° 55' 15"	115 29' 12"	N 03° W	90°	0-10	12		-0.01(.08)	0.28(.10) (east up)
1942/294(M6.5)	32° 59' 23"	115 54' 00"	N 50° W	90°	0-10	30		-0.11(.07)	0.23(.12) (east up)
1954/078(M6.2)	33° 19' 30"	116 13' 30"	N 46° W	90°	0-10	30		-0.26(.09)	0.48(.21) (west up)
1968/100(M6.2)	33° 08' 06"	116 05' 10"	N 40° W	83° NE	1-10	30		-0.70(.14)	0.39(.16)
1979/288(M6.6)	32° 46' 49"	115 25' 21"	N 37° W	90°	0-10	40		-0.65(.03)	0.45(.07) (west up)
SAN BERNARDINO REGION									
1948/338(M6.5)	33° 54' 00"	116 25' 00"	N 50° W	70° NE	5-25	18		0.09(.46)	0.09(.48)
BARSTOW REGION									
1952/203(M7.7)	Three rectangles. See Bakersfield region for parameters.								
1971/040(M6.4)	Four rectangles. See Los Angeles region for parameters.								
CHANNEL ISLANDS REGION									
1933/070(M6.3)	One rectangle. See Los Angeles region for parameters.								
1941/181(M5.9)	34° 20' 53"	119 23' 22"	N 70° W	80° N	0-10	18		-0.18(.02)	-0.13(.09)
1971/040(M6.4)	Four rectangles. See Los Angeles region for parameters.								
LOS ANGELES REGION									
1933/070(M6.3)	33° 43' 01"	118 02' 57"	N 47° W	90°	5-15	38		-0.09(.17)	0.04(.15) (east up)
1971/040(M6.4)	34° 16' 54"	118 22' 40"	N 70° W	35° N	1-12	12		-0.43(.07)	-2.23(.17)
1971/040(M6.4)	34° 19' 10"	118 21' 40"	N 70° W	52° N	8.7-20	12		1.76(.37)	-2.69(.57)
1971/040(M6.4)	34° 15' 52"	118 29' 22"	N 70° W	35° N	3-24	6		-0.29(.08)	-0.86(.10)
1971/040(M6.4)	34° 20' 30"	118 25' 15"	N 20° E	90°	1.7-6.9	10		1.86(.27)	0.00(fixed)
BAKERSFIELD REGION									
1952/203(M7.7)	35° 00' 53"	119 02' 27"	N 73° E	75° S	5-27	25		-0.07(.23)	-3.40(.29)
1952/203(M7.7)	35° 10' 33"	118 47' 51"	N 58° E	35° S	6-26	25		1.36(.07)	-1.20(.12)
1952/203(M7.7)	35° 17' 41"	118 35' 53"	N 43° E	20° S	6-26	25		0.19(.04)	-0.39(.06)
1971/040(M6.4)	Four rectangles. See Los Angeles region for parameters								
SIERRA NEVADA REGION									
1952/203(M7.7)	Three rectangles. See Bakersfield region for parameters.								
PARKFIELD REGION									
1934/158(M6.0)	35° 48' 00"	120 20' 58"	N 43° W	90°	0-10	20		-0.01(.04)	0.15(.06) (west up)
1966/179(M5.5)	35° 47' 25"	120 20' 25"	N 43° W	90°	5-20	32		-0.15(.07)	0.41(.08) (west up)
1966/179(M5.5)	35° 41' 23"	120 14' 37"	N 43° W	90°	0-5	4		-0.03(.01)	0.02(.03) (west up)
1966/179(M5.5)	35° 48' 12"	120 21' 19"	N 43° W	90°	0-5	28		-0.21(.01)	0.25(.03) (west up)
FALLON REGION									
1954/350(M7.1)	39° 13' 16"	118 07' 55"	N 13° E	63° E	0-5	36		-4.11(.08)	2.99(.09)
1954/350(M7.1)	39° 23' 38"	118 05' 02"	N 07° W	70° W	0-5	24		-1.44(.08)	0.27(.12)
1954/350(M6.9)	39° 38' 00"	118 11' 00"	N 08° E	60° E	0-6	38		-0.91(.22)	2.26(.29)
1954/350(M7.1)	39° 18' 25"	118 08' 53"	N 16° W	54° E	3.5-41	51.5		-0.40(.10)	1.29(.15)
1954/187(M6.6)	39° 28' 08"	118 27' 54"	N 12° E	63° E	0-10	40		0.68(.43)	-1.10(1.25)
1954/235(M6.8)	39° 28' 08"	118' 27' 54"	N 12° E	63° E	0-10	40		-1.48(.43)	3.15(1.21)

Comments: Depth measured along dip of rectangle.
 Positive strike slip denotes left-lateral slip.
 Positive dip slip denotes normal slip.

Table C.4.--District boundaries

San Diego region			San Bernardino region					
REGION BOUNDARY 32 30 114 0 32 30 117 30 33 30 117 30 33 30 114 0			REGION BOUNDARY 33 30 114 0 33 30 117 30 34 30 117 30 34 30 114 0					
DISTRICTS ***** DISTRICT 1 DISTRICT BOUNDARY 32 42 30.00 114 50 0.00 33 0 0.00 115 14 0.00 33 20 0.00 115 45 0.00 33 25 30.00 115 49 0.00 33 37 30.00 116 2 0.00 33 42 0.00 116 10 15.00 33 50 0.00 116 30 0.00 34 10 0.00 117 6 0.00 40 0 0.00 121 0 0.00 40 0 0.00 90 0 0.00 20 0 0.00 90 0 0.00 2 0 0.00 114 50 0.00			DISTRICT 7 DISTRICT BOUNDARY 33 0 0.00 115 49 0.00 32 58 0.00 115 51 0.00 33 9 0.00 116 6 30.00 33 15 0.00 116 14 0.00 33 30 0.00 116 38 0.00 33 51 30.00 117 2 0.00 33 54 30.00 117 9 30.00 34 10 0.00 117 9 30.00 33 50 0.00 116 30 0.00 33 42 0.00 116 10 15.00 33 37 30.00 116 2 0.00 33 25 30.00 115 49 0.00 33 20 0.00 115 45 0.00 33 10 16.00 115 40 0.00 33 1 0.00 115 36 0.00 32 55 30.00 115 34 0.00 33 0 0.00 115 39 0.00			DISTRICTS ***** DISTRICT 1 DISTRICT BOUNDARY 33 13 0.00 116 46 15.00 33 19 0.00 116 55 0.00 33 37 0.00 117 20 30.00 34 0 0.00 117 45 0.00 34 45 0.00 118 45 0.00 32 40 0.00 118 30 0.00 32 40 0.00 116 10 0.00		
DISTRICT 2 DISTRICT BOUNDARY 32 41 0.00 115 21 0.00 32 0 0.00 115 25 0.00 32 40 0.00 115 31 0.00 32 47 0.00 115 34 0.00 32 50 40.00 115 35 0.00 33 0 0.00 115 49 0.00 33 0 0.00 115 39 0.00 32 55 30.00 115 34 0.00			DISTRICT 2 DISTRICT BOUNDARY 33 13 0.00 116 46 15.00 33 19 0.00 116 55 0.00 33 37 0.00 117 20 30.00 34 0 0.00 117 45 0.00 34 45 0.00 118 45 0.00 34 25 0.00 117 45 0.00 34 13 0.00 117 29 10.00 34 6 15.59 117 19 36.00 34 6 17.00 117 19 51.00 34 4 41.00 117 19 51.00 34 4 41.00 117 18 14.00 34 4 37.00 117 18 8.00 33 54 30.00 117 9 30.00 33 51 30.00 117 2 0.00 33 30 0.00 116 38 0.00 33 15 0.00 116 14 0.00 32 45 0.00 115 45 0.00 32 40 0.00 116 10 0.00					
DISTRICT 3 DISTRICT BOUNDARY 32 41 29.00 115 8 0.00 32 55 0.00 115 25 0.00 32 58 0.00 115 31 30.00 33 10 0.00 115 37 0.00 33 20 0.00 115 45 0.00 33 0 0.00 115 14 0.00 32 42 30.00 114 50 0.00 20 0 0.00 114 50 0.00 20 0 0.00 115 8 0.00			DISTRICT 8 DISTRICT BOUNDARY 32 55 30.00 115 34 0.00 33 1 0.00 115 36 0.00 33 10 16.00 115 40 0.00 33 20 0.00 115 45 0.00 33 10 0.00 115 37 0.00 32 58 0.00 115 31 30.00					
DISTRICT 4 DISTRICT BOUNDARY 32 41 0.00 115 21 0.00 32 55 30.00 115 34 0.00 32 58 0.00 115 31 30.00 32 55 0.00 115 25 0.00 32 41 29.00 115 8 0.00 20 0 0.00 115 8 0.00 20 0 0.00 115 21 0.00			DISTRICT 9 DISTRICT BOUNDARY 32 54 0.00 116 17 5.00 33 6 0.00 116 34 40.00 33 13 0.00 116 46 15.00 33 19 0.00 116 55 0.00 33 37 0.00 117 20 30.00 34 0 0.00 117 45 0.00 35 0 0.00 118 30 0.00 33 54 30.00 117 9 30.00 33 51 30.00 117 2 0.00 33 30 0.00 116 38 0.00 33 15 0.00 116 14 0.00 33 9 0.00 116 6 30.00					
DISTRICT 5 DISTRICT BOUNDARY 32 38 30.00 115 50 0.00 32 47 0.00 116 0 0.00 32 54 0.00 116 17 5.00 33 9 0.00 116 6 30.00 32 58 0.00 115 51 0.00 32 57 0.00 115 49 24.00 32 50 40.00 115 35 0.00 32 47 0.00 115 34 0.00 32 40 0.00 115 31 0.00 20 0 0.00 115 21 0.00 20 0 0.00 115 50 0.00			DISTRICT 10 DISTRICT BOUNDARY 20 0 0.00 115 50 0.00 32 38 30.00 115 50 0.00 32 47 0.00 116 0 0.00 32 54 0.00 116 17 5.00 33 6 0.00 116 34 40.00 33 13 0.00 116 46 15.00 33 19 0.00 116 55 0.00 33 37 0.00 117 20 30.00 34 0 0.00 117 45 0.00 35 0 0.00 118 30 0.00 35 0 0.00 120 0 0.00 20 0 0.00 120 0 0.00					
DISTRICT 6 DISTRICT BOUNDARY 32 50 40.00 115 34 40.00 32 57 0.00 115 49 24.00 32 58 0.00 115 51 0.00 33 0 0.00 115 49 0.00			DISTRICT 3 DISTRICT BOUNDARY 33 15 0.00 116 14 0.00 33 30 0.00 116 38 0.00 33 51 30.00 117 2 0.00 33 54 30.00 117 9 30.00 34 4 37.00 117 18 8.00 33 57 0.00 116 44 0.00 33 56 0.00 116 37 0.00 33 52 30.00 116 25 30.00 33 52 0.00 116 21 50.00 33 38 10.00 116 2 0.00 33 38 3.00 116 2 0.00 33 10 0.00 115 45 0.00					
			DISTRICT 4 DISTRICT BOUNDARY 33 57 0.00 116 44 0.00 34 4 37.00 117 18 8.00 34 4 32.00 117 18 2.00 34 6 4.00 117 18 2.00 34 6 4.63 117 19 34.00 34 6 15.59 117 19 36.00 34 13 0.00 117 29 10.00 34 25 0.00 117 45 0.00 34 16 25.00 117 27 9.00 34 16 0.00 117 25 24.00 34 12 2.00 117 19 56.00 34 11 56.00 117 19 44.00 34 4 30.00 116 58 6.00					
			DISTRICT 5 DISTRICT BOUNDARY 33 38 10.00 116 2 0.00 33 52 0.00 116 21 50.00 33 52 10.00 116 21 30.00 34 1 54.00 116 40 9.00 34 2 0.00 116 30 0.00 34 9 0.00 115 58 0.00 33 33 40.00 115 21 0.00 33 15 0.00 115 5 0.00 32 30 0.00 115 5 0.00 32 30 0.00 115 45 0.00 33 10 0.00 115 45 0.00 33 38 3.00 116 2 0.00					

(continued)

Table C.4.--Continued

San Bernardino
region

Barstow region

DISTRICT 6

DISTRICT BOUNDARY

34 2 0.00 116 30 0.00
34 1 54.00 116 40 9.00
34 4 50.00 116 58 6.00
34 11 56.00 117 19 44.00
34 12 2.00 117 19 56.00
34 16 0.00 117 25 24.00
34 16 25.00 117 27 9.00
34 25 0.00 117 45 0.00
35 0 0.00 118 45 0.00
35 0 0.00 117 15 0.00
34 45 0.00 117 15 0.00

DISTRICT 7

DISTRICT BOUNDARY

34 2 0.00 116 30 0.00
34 45 0.00 117 15 0.00
34 45 0.00 116 35 0.00
34 31 0.00 116 21 0.00
34 18 0.00 116 7 0.00
34 9 0.00 115 58 0.00

DISTRICT 8

DISTRICT BOUNDARY

34 45 0.00 115 30 0.00
33 15 0.00 114 30 0.00
32 30 0.00 114 30 0.00
32 30 0.00 115 5 0.00
33 15 0.00 115 5 0.00
33 33 40.00 115 21 0.00
34 9 0.00 115 58 0.00
34 18 0.00 116 7 0.00
34 31 0.00 116 21 0.00
34 45 0.00 116 35 0.00

DISTRICT 9

DISTRICT BOUNDARY

34 4 33.00 117 18 8.00
34 4 41.00 117 18 14.00
34 4 41.00 117 19 51.00
34 6 17.00 117 19 51.00
34 6 15.59 117 19 36.00
34 6 4.63 117 19 34.00
34 6 4.00 117 18 2.00
34 4 32.00 117 18 2.00

DISTRICT 10

DISTRICT BOUNDARY

33 52 0.00 116 21 50.00
33 52 30.00 116 25 30.00
33 56 0.00 116 37 0.00
33 57 0.00 116 44 0.00
34 4 30.00 116 58 6.00
34 11 56.00 117 19 44.00
34 4 50.00 116 58 6.00
34 1 54.00 116 40 9.00
33 52 10.00 116 21 30.00

DISTRICT 11

DISTRICT BOUNDARY

34 45 0.00 115 30 0.00
33 15 0.00 114 30 0.00
32 0 0.00 112 0 0.00
36 0 0.00 112 0 0.00

REGION BOUNDARY

34 30 114 0
34 30 118 30
35 30 118 30
35 30 114 0

DISTRICTS *****

DISTRICT 1

DISTRICT BOUNDARY

34 0 0.00 117 45 0.00
34 13 26.00 118 3 38.00
34 16 9.00 118 14 17.00
34 21 10.00 118 25 43.00
34 26 10.00 118 32 31.00
34 28 44.00 118 33 57.00
34 45 0.00 119 0 0.00
34 45 3.00 118 44 32.00
34 44 0.00 118 36 0.00
34 40 34.00 118 27 17.50
34 40 32.89 118 26 57.00
34 34 52.00 118 11 12.40
34 34 44.00 118 10 54.80
34 32 0.00 118 4 16.00
34 31 59.00 118 4 3.00
34 29 29.00 117 55 5.20
34 25 0.00 117 45 0.00
34 22 4.00 117 39 36.00
34 22 2.50 117 39 26.00
34 13 0.00 117 29 10.00

DISTRICT 2

DISTRICT BOUNDARY

34 13 0.00 117 29 10.00
34 22 2.50 117 39 26.00
34 22 4.00 117 39 36.00
34 25 0.00 117 45 0.00
34 29 29.00 117 55 5.20
34 31 59.00 118 4 3.00
34 32 0.00 118 4 16.00
34 34 44.00 118 10 54.00
34 34 52.00 118 11 12.40
34 40 32.89 118 26 57.00
34 40 34.00 118 27 17.50
34 44 0.00 118 36 0.00
34 51 30.00 118 25 0.00
34 59 0.00 118 7 0.00
35 6 0.00 117 45 0.00
35 5 0.00 117 34 45.00
34 42 0.00 117 12 0.00
34 30 0.00 117 0 0.00
34 0 0.00 116 30 0.00

DISTRICT 3

DISTRICT BOUNDARY

34 45 0.00 119 0 0.00
34 56 45.00 119 50 0.00
35 1 39.00 118 59 45.00
35 21 0.00 118 35 11.00
35 47 13.00 118 25 36.00
36 6 38.00 118 23 25.00
36 8 0.00 115 0 0.00
35 25 27.00 117 43 39.00
35 7 14.80 118 17 0.00
35 3 25.00 118 23 31.00

DISTRICT 4

DISTRICT BOUNDARY

34 56 45.00 119 50 0.00
35 1 39.00 118 59 45.00
35 21 0.00 118 35 11.00
35 47 13.00 118 25 36.00
36 6 38.00 118 23 25.00

DISTRICT 5

DISTRICT BOUNDARY

35 25 27.00 117 43 39.00
36 6 0.00 115 0 0.00
35 4 14.00 116 19 21.00
34 56 7.00 116 39 57.00

DISTRICT 6

DISTRICT BOUNDARY

34 0 0.00 116 30 0.00
34 30 0.00 117 0 0.00
34 42 0.00 117 12 0.00
35 5 0.00 117 34 45.00
35 6 0.00 117 45 0.00
35 7 14.80 118 17 0.00
35 25 27.00 117 43 39.00
34 56 7.00 116 39 57.00
34 45 0.00 116 25 0.00
34 9 0.00 115 47 10.00

DISTRICT 7

DISTRICT BOUNDARY

36 6 0.00 115 0 0.00
36 6 0.00 112 40 0.00
33 0 0.00 112 40 0.00
33 0 0.00 114 30 0.00
33 30 0.00 115 45 0.00
34 9 0.00 115 47 10.00
34 45 0.00 116 25 0.00
34 56 7.00 116 39 57.00
35 4 14.00 116 19 21.00

DISTRICT 8

DISTRICT BOUNDARY

34 45 0.00 119 0 0.00
34 45 3.00 118 44 32.00
34 44 0.00 118 36 0.00
34 51 30.00 118 25 0.00
34 59 0.00 118 7 0.00
35 6 0.00 117 45 0.00
35 7 14.80 118 17 0.00
35 3 25.00 118 23 31.00

DISTRICT 9

DISTRICT BOUNDARY

34 0 0.00 117 45 0.00
34 13 26.00 118 3 38.00
34 16 9.00 118 14 17.00
34 21 10.00 118 25 43.00
34 26 10.00 118 32 31.00
34 28 44.00 118 33 57.00
34 45 0.00 119 0 0.00
34 0 0.00 119 0 0.00

Table C.4.--Continued

Channel Islands region				Los Angeles region							
REGION BOUNDARY				DISTRICT 9				REGION BOUNDARY			
32 30	117	30		34 48	15.00	118	48 52.60	33 30	117	30	
32 30	121	0		DISTRICT BOUNDARY				33 30	118	45	
34 30	121	0		34 46	38.19	118	48 52.60	34 30	118	45	
34 30	118	45		34 37	46.00	118	44 6.50	34 30	117	30	
33 30	118	45		34 28	17.87	118	34 9.00				
33 30	117	30		34 27	0.59	118	32 27.69	DISTRICTS *****			
DISTRICTS *****				34 18	31.90	118	13 59.39	DISTRICT 1			
DISTRICT 1				34 4	57.52	118	3 38.69	DISTRICT BOUNDARY			
DISTRICT BOUNDARY				34 4	57.52	118	34 25.24	33 37	13.58	117	17 46.42
33 30	0.00	118	0 0.00	34 6	0.00	118	35 0.00	33 37	13.58	116	0 0.00
32 45	0.00	118	0 0.00	34 2	26.50	118	46 31.16	32 30	0.00	116	0 0.00
33 30	0.00	119	0 0.00	34 1	52.00	118	50 39.89	32 30	0.00	118	0 0.00
DISTRICT 2				33 30	0.00	119	0 0.00	33 30	0.00	118	0 0.00
DISTRICT BOUNDARY				34 9	0.00	119	16 0.00	33 36	26.00	117	55 6.58
33 30	0.00	119	0 0.00	34 9	0.00	120	0 0.00	33 36	44.18	117	55 6.58
32 45	0.00	118	0 0.00	34 9	0.00	121	0 0.00	33 47	19.57	118	8 16.46
32 45	0.00	119	0 0.00	35 5	0.00	121	0 0.00	33 51	38.85	118	13 55.00
DISTRICT 3				35 5	0.00	118	48 52.60	33 56	16.84	118	18 12.00
DISTRICT BOUNDARY								34 0	15.91	118	21 44.50
33 30	0.00	119	0 0.00					34 4	2.25	118	29 9.61
33 0	0.00	119	0 0.00					34 8	13.09	118	23 31.00
33 30	0.00	119	45 0.00					34 7	7.40	118	19 29.67
DISTRICT 4								34 8	28.21	118	15 37.33
DISTRICT BOUNDARY								34 9	48.85	118	7 7.87
33 30	0.00	119	45 0.00					33 56	26.53	117	53 25.85
33 0	0.00	119	0 0.00					33 56	6.28	117	51 39.37
33 0	0.00	119	45 0.00					33 50	41.89	117	36 36.19
DISTRICT 5								33 50	29.60	117	33 12.00
DISTRICT BOUNDARY								33 46	22.50	117	31 59.00
33 30	0.00	119	0 0.00					DISTRICT 2			
34 9	0.00	119	16 0.00					DISTRICT BOUNDARY			
34 9	0.00	120	0 0.00					32 30	0.00	118	0 0.00
33 55	0.00	119	55 0.00					33 30	0.00	118	0 0.00
33 30	0.00	119	45 0.00					33 30	0.00	119	15 0.00
DISTRICT 6								34 1	52.00	119	15 0.00
DISTRICT BOUNDARY								34 15	0.00	119	15 0.00
33 30	0.00	119	45 0.00					34 15	0.00	119	30 0.00
33 55	0.00	119	55 0.00					32 30	0.00	119	30 0.00
34 9	0.00	120	0 0.00					DISTRICT 3			
34 9	0.00	121	0 0.00					DISTRICT BOUNDARY			
33 30	0.00	121	0 0.00					33 37	13.58	116	0 0.00
DISTRICT 7								34 0	0.00	116	0 0.00
DISTRICT BOUNDARY								34 9	21.67	117	20 45.66
34 1	52.00	118	50 39.89					34 11	29.90	118	0 59.46
34 2	26.50	118	46 31.16					34 9	48.85	118	7 7.87
34 6	0.00	118	35 0.00					33 56	26.53	117	53 25.85
34 4	57.52	118	34 25.24					33 56	6.28	117	51 39.37
34 4	57.52	118	3 38.69					33 50	41.89	117	36 36.19
34 4	57.52	116	0 0.00					33 50	29.60	117	33 12.00
32 0	0.00	116	0 0.00					33 46	22.50	117	31 59.00
32 0	0.00	118	0 0.00					33 37	13.58	117	17 46.42
33 30	0.00	118	0 0.00					DISTRICT 4			
33 30	0.00	119	0 0.00					DISTRICT BOUNDARY			
DISTRICT 8								34 9	48.85	118	7 7.87
DISTRICT BOUNDARY								34 16	8.39	118	14 17.08
34 4	57.52	116	0 0.00					34 18	31.90	118	13 59.39
34 48	15.00	116	0 0.00					34 27	0.59	118	32 27.69
34 48	15.00	118	48 52.60					34 28	17.87	118	34 9.00
34 46	38.19	118	48 52.60					34 37	46.00	118	44 6.50
34 37	46.00	118	44 6.50					34 46	38.19	118	48 52.60
34 28	17.87	118	34 9.00					34 46	38.19	119	15 0.00
34 27	0.59	118	32 27.69					34 15	0.00	119	15 0.00
34 18	31.90	118	13 59.39					34 1	52.00	119	15 0.00
34 4	57.52	118	3 38.69					34 1	52.00	118	50 39.89
								34 2	26.50	118	46 31.16
								34 6	0.00	118	35 0.00
								34 4	57.52	118	34 25.24
								34 4	2.25	118	29 9.61
								34 8	13.09	118	23 31.00
								34 7	7.40	118	19 29.67
								34 8	28.21	118	15 37.33
								34 18	31.90	118	13 59.39
								34 18	31.90	118	0 59.46
								34 14	53.00	117	40 27.24
								34 17	12.70	117	29 9.79

(continued)

Table C.4.--Continued

Los Angeles region	Bakersfield	region
DISTRICT 5	REGION BOUNDARY	
DISTRICT BOUNDARY	34 30 118 30	34 53 0.00 118 46 6.00
34 22 2.52 117 39 26.10	34 30 121 30	34 49 30.00 118 52 0.00
34 22 4.29 117 39 36.10	35 30 121 30	34 46 13.30 118 45 47.00
34 31 59.05 118 4 3.16	35 30 118 30	34 46 11.90 118 45 30.19
34 32 0.69 118 4 16.25		34 40 34.00 118 27 17.50
34 32 51.07 118 6 28.03	DISTRICTS *****	34 40 32.89 118 26 57.00
34 34 44.78 118 10 54.85	DISTRICT 1	34 38 16.00 118 20 35.50
34 34 52.39 118 11 12.46	DISTRICT BOUNDARY	34 30 0.00 118 0 0.00
34 37 53.80 118 21 11.17	36 0 0.00 121 30 0.00	34 0 0.00 117 0 0.00
34 39 14.83 118 22 20.87	33 30 0.00 121 30 0.00	35 0 0.00 117 0 0.00
34 46 38.21 118 44 32.54	33 30 0.00 118 30 0.00	DISTRICT 6
34 46 38.19 118 48 52.60	34 0 0.00 118 30 0.00	DISTRICT BOUNDARY
34 37 46.00 118 44 6.50	34 50 0.00 119 35 0.00	34 30 0.00 118 0 0.00
34 28 17.87 118 34 9.00	35 0 0.00 119 59 0.00	34 38 16.00 118 20 35.50
34 27 0.59 118 32 27.69	35 22 9.00 120 33 35.00	34 40 32.89 118 26 57.00
	35 24 6.20 120 37 6.90	34 40 34.00 118 27 17.50
	35 24 20.35 120 37 14.50	34 46 11.90 118 45 30.19
	35 28 30.00 120 44 50.00	34 46 13.30 118 45 47.00
		34 49 30.00 118 52 0.00
		34 30 0.00 118 37 0.00
		34 0 0.00 118 0 0.00
DISTRICT 6	DISTRICT 2	
DISTRICT BOUNDARY	DISTRICT BOUNDARY	
34 18 31.90 118 13 59.39	36 0 0.00 121 30 0.00	DISTRICT 7
34 18 31.90 118 0 59.46	35 28 30.00 120 44 30.00	DISTRICT BOUNDARY
34 14 53.00 117 40 27.24	35 24 20.39 120 37 14.50	34 50 0.00 119 35 0.00
34 17 12.70 117 29 9.79	35 24 6.20 120 37 6.90	34 57 32.00 119 25 0.00
34 15 53.75 117 27 51.10	35 22 9.00 120 33 35.00	34 50 33.00 118 58 51.00
34 15 10.59 117 27 48.59	35 0 0.00 119 59 0.00	34 49 30.00 118 52 0.00
34 9 21.67 117 20 45.66	34 50 0.00 119 35 0.00	34 30 0.00 118 37 0.00
34 11 29.90 118 0 59.46	34 57 32.00 119 25 0.00	34 0 0.00 118 0 0.00
34 9 48.85 118 7 7.87	35 3 13.00 119 33 15.00	34 0 0.00 118 30 0.00
34 16 8.39 118 14 17.08	35 30 31.00 120 0 23.69	
DISTRICT 7	35 39 0.00 120 10 9.00	
DISTRICT BOUNDARY	36 0 0.00 120 36 0.00	
33 30 0.00 118 0 0.00	DISTRICT 3	
33 36 26.00 117 55 6.58	DISTRICT BOUNDARY	
33 36 44.18 117 55 6.58	36 0 0.00 120 10 9.00	
33 47 19.57 118 8 16.46	35 39 0.00 120 10 9.00	
33 51 32.85 118 13 55.00	35 30 31.00 120 0 23.69	
33 56 16.84 118 18 12.00	35 3 13.00 119 33 15.00	
34 0 15.91 118 21 44.50	34 57 32.00 119 25 0.00	
34 4 2.25 118 29 9.61	35 1 41.39 119 0 8.75	
34 4 57.52 118 34 25.24	35 2 41.00 118 57 44.00	
34 6 0.00 118 35 0.00	35 2 50.00 118 57 25.69	
34 2 26.50 118 46 31.16	35 7 26.00 118 49 17.80	
34 1 52.00 118 50 39.89	35 15 0.00 118 35 11.50	
34 1 52.00 119 15 0.00	35 38 29.60 118 28 33.00	
33 30 0.00 119 15 0.00	36 0 0.00 118 28 0.00	
DISTRICT 8	DISTRICT 4	
DISTRICT BOUNDARY	DISTRICT BOUNDARY	
34 0 0.00 116 0 0.00	36 0 0.00 118 28 0.00	
34 9 21.67 117 20 45.66	35 38 29.60 118 28 33.00	
34 15 10.59 117 27 48.59	35 15 0.00 118 35 11.50	
34 15 53.75 117 27 51.10	35 7 26.00 118 49 17.80	
34 17 12.70 117 29 9.79	35 2 50.00 118 57 25.69	
34 22 2.52 117 39 26.10	35 2 41.00 118 57 44.00	
34 22 4.29 117 39 36.10	35 1 41.39 119 0 8.75	
34 31 59.05 118 4 3.16	34 57 32.00 119 25 0.00	
34 32 0.69 118 4 16.25	34 50 33.00 118 58 51.00	
34 32 51.07 118 6 28.03	34 49 30.00 118 52 0.00	
34 34 44.78 118 10 54.85	34 53 0.00 118 46 6.00	
34 34 52.39 118 11 12.46	34 53 5.70 118 45 58.70	
34 37 53.80 118 21 11.17	34 53 46.00 118 45 56.70	
34 39 14.83 118 22 20.87	34 57 21.00 118 34 20.60	
34 46 38.21 118 44 32.54	35 0 0.00 118 25 5.00	
34 46 38.19 118 48 52.60	35 16 51.60 118 0 0.00	
34 46 38.19 119 15 0.00	36 0 0.00 118 0 0.00	
35 0 0.00 119 15 0.00	DISTRICT 5	
35 0 0.00 116 0 0.00	DISTRICT BOUNDARY	
	35 16 51.60 118 0 0.00	
	35 0 0.00 118 25 5.00	
	34 57 21.00 118 34 20.60	
	34 53 46.00 118 45 56.70	
	34 53 5.70 118 45 58.70	

Table C.4.--Continued

Sierra Nevada region			Parkfield region								
REGION BOUNDARY			DISTRICT 7			REGION BOUNDARY					
35 30	115	0	DISTRICT BOUNDARY	37 20	0.00	118 13	30.00	35 30	119	30	
35 30	119	30	37 0	0.00	118 13	30.00	35 30	122	0		
37 0	119	30	36 48	24.00	118 7	30.00	36 30	122	0		
37 0	115	0	36 35	28.00	118 3	45.00	36 30	119	30		
DISTRICTS *****			36 23	26.50	118 1	0.00	DISTRICTS *****				
DISTRICT 1			36 20	0.00	118 6	0.00	DISTRICT 1				
DISTRICT BOUNDARY			35 59	13.00	117 56	0.00	DISTRICT BOUNDARY				
35 10	0.00	120 0	0.00	35 45	43.00	117 55	0.00	35 0	0.00	121 38	0.00
34 30	0.00	119 25	0.00	35 30	0.00	118 3	0.00	34 50	0.00	120 36	0.00
35 1	41.39	119 0	8.75	35 17	0.00	118 3	0.00	35 0	0.00	119 59	0.00
35 2	41.00	118 57	44.00	35 26	0.00	117 44	0.00	35 22	9.00	120 33	35.00
35 2	50.00	118 57	25.69	35 33	13.00	117 14	7.60	35 24	6.20	120 37	6.90
35 7	26.00	118 49	17.80	35 36	0.00	117 3	0.00	35 24	20.39	120 37	14.50
35 15	0.00	118 35	11.50	35 37	0.00	116 41	0.00	35 28	30.00	120 44	50.00
35 30	0.00	118 31	0.00	35 37	0.00	113 50	0.00	35 40	35.00	121 3	5.00
35 30	0.00	120 0	0.00	38 0	0.00	113 30	0.00	36 0	44.00	121 30	5.00
				38 0	0.00	114 3	0.00	36 7	32.00	121 38	0.00
DISTRICT 2							36 5	0.00	122 0	0.00	
DISTRICT BOUNDARY							DISTRICT 2				
35 30	0.00	119 31	0.00					DISTRICT BOUNDARY			
35 47	13.30	118 25	37.00					36 7	32.00	121 36	0.00
36 0	0.00	118 24	0.00					36 0	44.00	121 30	5.00
36 0	0.00	120 0	0.00					35 40	35.00	121 3	5.00
35 30	0.00	120 0	0.00					35 28	30.00	120 44	50.00
								35 24	20.39	120 37	14.50
DISTRICT 3							35 24	6.20	120 37	6.90	
DISTRICT BOUNDARY							35 22	9.00	120 33	35.00	
35 0	0.00	118 24	0.00					35 0	0.00	119 59	0.00
36 20	0.00	118 21	0.00					35 0	13.00	119 33	15.00
37 20	0.00	118 24	0.00					35 15	0.00	119 50	0.00
37 30	0.00	118 24	0.00					35 21	10.00	119 55	0.00
37 30	0.00	120 0	0.00					35 29	50.00	120 3	30.00
36 0	0.00	120 0	0.00					35 35	30.00	120 7	30.00
								35 38	0.00	120 10	50.00
								35 41	18.00	120 14	30.00
								35 41	29.00	120 14	43.00
								35 43	0.00	120 15	55.00
								35 44	50.00	120 17	50.00
								35 49	52.00	120 22	0.00
								35 53	30.00	120 26	30.00
								35 57	50.00	120 31	50.00
								35 59	0.00	120 34	0.00
								36 3	50.00	120 39	0.00
								36 11	0.00	120 45	6.00
								36 13	4.00	120 48	10.00
								36 24	0.00	121 0	0.00
								36 28	50.00	121 5	0.00
								36 37	0.00	121 12	0.00
								36 38	3.00	121 14	0.00
								36 38	11.00	121 14	13.00
								36 45	6.00	121 23	7.00
								36 55	51.00	121 38	35.10
								36 55	0.00	122 0	0.00
								36 5	0.00	122 0	0.00
								DISTRICT 3			
								DISTRICT BOUNDARY			
								35 13	13.00	119 33	15.00
								35 15	0.00	119 50	0.00
								35 21	10.00	119 55	0.00
								35 29	50.00	120 3	30.00
								35 35	30.00	120 7	30.00
								35 38	0.00	120 10	50.00
								35 41	18.00	120 14	30.00
								35 41	29.00	120 14	43.00
								35 43	0.00	120 15	55.00
								35 44	50.00	120 17	50.00
								35 49	52.00	120 22	0.00
								35 53	30.00	120 26	30.00
								35 57	0.00	120 20	0.00
								36 5	59.00	120 1	7.00
								35 51	0.00	119 49	0.00
								35 35	12.00	119 35	9.00
								35 21	19.00	119 16	17.00
								35 15	0.00	119 0	0.00
								35 0	0.00	119 0	0.00

(continued)

Table C.4.--Continued

Parkfield region	Monterey	region
DISTRICT 4	REGION BOUNDARY	DISTRICT 6
DISTRICT BOUNDARY	36 30 119 30	DISTRICT BOUNDARY
35 53 30.00 120 26 30.00	36 30 123 0	37 30 0.00 121 27 0.00
35 57 50.00 120 31 50.00	37 0 123 0	37 11 0.00 121 14 0.00
35 59 0.00 120 34 0.00	37 0 119 30	36 47 40.00 120 55 18.00
36 3 50.00 120 39 0.00	DISTRICTS *****	36 20 41.00 120 33 0.00
36 11 0.00 120 45 6.00	DISTRICT 1	35 59 0.00 120 34 0.00
36 13 4.00 120 48 10.00	DISTRICT BOUNDARY	35 59 0.00 118 30 0.00
36 12 49.00 120 42 57.00	36 24 0.00 121 0 0.00	37 30 0.00 118 30 0.00
36 16 55.00 120 34 8.00	36 28 50.00 121 5 0.00	
36 26 44.00 120 20 11.00	36 37 0.00 121 12 0.00	
36 11 52.00 120 8 0.00	36 38 3.00 121 14 0.00	
36 5 59.00 120 1 7.00	36 38 11.00 121 14 13.00	
35 57 0.00 120 20 0.00	36 45 6.00 121 23 7.00	
DISTRICT 5	36 45 10.00 121 23 17.00	
DISTRICT BOUNDARY	36 46 24.00 121 25 29.00	
36 28 50.00 121 5 0.00	36 46 35.00 121 25 40.00	
36 37 0.00 121 12 0.00	36 49 55.00 121 30 0.00	
36 38 3.00 121 14 0.00	36 55 51.00 121 38 35.10	
36 38 11.00 121 14 13.00	37 4 0.00 121 50 0.00	
36 45 6.00 121 23 7.00	37 19 12.00 122 8 45.00	
36 55 51.00 121 38 35.10	37 30 0.00 122 29 45.00	
36 59 0.00 121 26 0.00	37 30 0.00 122 30 0.00	
36 49 5.00 121 22 45.00	35 59 0.00 122 30 0.00	
36 43 0.00 121 16 0.00	35 59 0.00 120 34 0.00	
36 38 0.00 121 8 0.00	36 3 50.00 120 39 0.00	
	36 10 0.00 120 46 0.00	
	DISTRICT 2	
DISTRICT 6	DISTRICT BOUNDARY	
DISTRICT BOUNDARY	36 28 50.00 121 5 0.00	
36 13 4.00 120 48 10.00	36 38 0.00 121 8 0.00	
36 24 0.00 121 0 0.00	36 43 0.00 121 16 0.00	
36 28 50.00 121 5 0.00	36 45 0.00 121 23 0.00	
36 38 0.00 121 8 0.00	36 51 0.00 121 25 0.00	
36 43 0.00 121 16 0.00	36 54 0.00 121 27 0.00	
36 49 5.00 121 22 45.00	36 57 0.00 121 38 0.00	
36 59 0.00 121 26 0.00	37 4 0.00 121 50 0.00	
36 55 0.00 120 50 0.00	36 55 51.00 121 38 35.10	
36 33 4.00 120 29 17.00	36 49 55.00 121 30 0.00	
36 26 44.00 120 20 11.00	36 46 35.00 121 25 40.00	
36 16 55.00 120 34 8.00	36 46 24.00 121 25 29.00	
36 12 49.00 120 42 57.00	36 45 10.00 121 23 17.00	
	36 45 6.00 121 23 7.00	
	36 38 11.00 121 14 13.00	
	36 38 3.00 121 14 0.00	
	36 37 0.00 121 12 0.00	
	DISTRICT 3	
DISTRICT 7	DISTRICT BOUNDARY	
DISTRICT BOUNDARY	36 54 0.00 121 27 0.00	
35 0 0.00 119 0 0.00	36 59 0.00 121 26 0.00	
35 15 0.00 119 0 0.00	37 5 0.00 121 32 0.00	
35 21 19.00 119 16 17.00	37 5 0.00 121 32 0.00	
35 35 12.00 119 35 9.00	37 30 0.00 121 54 43.00	
35 51 0.00 119 49 0.00	37 30 0.00 122 29 45.00	
36 5 59.00 120 1 7.00	37 19 12.00 122 8 45.00	
36 11 52.00 120 8 0.00	37 4 0.00 121 50 0.00	
36 26 44.00 120 20 11.00	36 57 0.00 121 38 0.00	
36 33 4.00 120 29 17.00	DISTRICT 4	
36 55 0.00 120 50 0.00	DISTRICT BOUNDARY	
37 11 0.00 121 14 0.00	36 59 0.00 121 26 0.00	
37 32 0.00 121 21 0.00	37 5 0.00 121 32 0.00	
37 32 0.00 119 0 0.00	37 30 0.00 121 54 43.00	
37 11 0.00 119 0 0.00	37 30 0.00 121 27 0.00	
36 55 0.00 119 0 0.00	37 11 0.00 121 14 0.00	
	36 47 40.00 120 55 18.00	
	36 20 41.00 120 33 0.00	
	35 59 0.00 120 34 0.00	
	36 3 50.00 120 39 0.00	
	36 10 0.00 120 46 0.00	
	36 24 0.00 121 0 0.00	
	36 28 50.00 121 5 0.00	
	36 38 0.00 121 12 0.00	
	36 43 0.00 121 16 0.00	
	36 49 5.00 121 22 45.00	
	36 55 33.00 121 21 22.00	
	DISTRICT 5	
	DISTRICT BOUNDARY	
	36 43 0.00 121 16 0.00	
	36 49 0.00 121 23 0.00	
	36 51 0.00 121 25 0.00	
	36 54 0.00 121 27 0.00	
	36 59 0.00 121 26 0.00	
	36 55 33.00 121 21 22.00	
	36 49 5.00 121 22 45.00	

Table C.4.--Continued

Yosemite region				San Francisco region			
REGION BOUNDARY				REGION BOUNDARY			
37 0	117 0			37 42	25.00	121 41	20.00
37 0	121 45			37 42	0.00	120 30	0.00
38 0	121 45			38 6	0.00	120 50	0.00
38 0	117 0			38 35	0.00	121 13	0.00
DISTRICTS *****				DISTRICTS *****			
DISTRICT 1				DISTRICT 6			
DISTRICT BOUNDARY				DISTRICT BOUNDARY			
36 30	0.00	121 0	0.00	37 35	36.00	121 52	30.00
36 28	50.00	121 5	0.00	37 34	52.00	121 51	59.00
36 37	0.00	121 12	0.00	37 34	44.00	121 51	52.00
36 38	3.00	121 14	0.00	37 27	0.00	121 48	45.00
36 38	11.00	121 14	13.00	37 17	34.00	121 42	25.69
36 45	6.00	121 23	7.00	37 0	0.00	121 28	45.00
36 45	10.00	121 23	17.00	36 59	0.00	121 26	0.00
36 46	24.00	121 25	29.00	36 55	33.00	121 22	0.00
36 46	35.00	121 25	40.00	36 49	5.00	121 22	45.00
36 49	55.00	121 30	0.00	36 43	0.00	121 16	0.00
36 55	51.00	121 38	35.10	36 38	0.00	121 8	0.00
37 4	0.00	121 50	0.00	36 28	50.00	121 5	0.00
37 19	12.00	122 8	45.00	36 30	0.00	121 0	0.00
37 30	0.00	122 20	45.00	36 10	0.00	120 46	0.00
37 30	0.00	122 30	0.00	36 10	0.00	119 10	0.00
35 59	0.00	122 30	0.00	36 30	0.00	119 20	0.00
35 59	0.00	120 34	0.00	37 42	0.00	120 30	0.00
36 3	50.00	120 39	0.00	37 42	25.00	121 41	20.00
36 10	0.00	120 46	0.00	37 39	0.00	121 46	55.00
DISTRICT 2				DISTRICT 7			
DISTRICT BOUNDARY				DISTRICT BOUNDARY			
36 28	50.00	121 5	0.00	38 35	0.00	121 13	0.00
36 38	0.00	121 8	0.00	38 6	0.00	120 50	0.00
36 43	0.00	121 16	0.00	37 42	0.00	120 30	0.00
36 49	0.00	121 23	0.00	36 30	0.00	119 20	0.00
36 51	0.00	121 25	0.00	36 30	0.00	118 5	0.00
36 54	0.00	121 27	0.00	36 41	0.00	118 5	0.00
36 57	0.00	121 38	0.00	36 56	0.00	118 11	0.00
37 4	0.00	121 50	0.00	37 2	0.00	118 12	0.00
36 55	51.00	121 38	35.10	37 14	0.00	118 20	0.00
36 49	55.00	121 30	0.00	37 23	0.00	118 20	0.00
36 46	35.00	121 25	40.00	37 32	0.00	118 21	0.00
36 46	24.00	121 25	29.00	37 42	0.00	118 23	0.00
36 45	10.00	121 23	17.00	37 51	0.00	118 24	0.00
36 45	6.00	121 23	7.00	38 0	0.00	118 32	0.00
36 38	11.00	121 14	13.00	38 35	0.00	118 35	0.00
36 38	3.00	121 14	0.00	38 45	0.00	118 35	0.00
36 37	0.00	121 12	0.00	38 45	0.00	121 13	0.00
DISTRICT 3				DISTRICT 8			
DISTRICT BOUNDARY				DISTRICT BOUNDARY			
36 54	0.00	121 27	0.00	37 17	34.00	121 42	25.69
36 59	0.00	121 26	0.00	37 27	0.00	121 48	45.00
37 0	0.00	121 28	45.00	37 34	44.00	121 51	52.00
37 17	34.00	121 42	25.69	37 34	52.00	121 51	59.00
37 18	51.39	121 45	16.09	37 35	36.00	121 52	30.00
37 31	21.90	121 56	33.00	37 41	7.00	121 55	45.00
37 31	31.69	121 56	43.00	37 49	4.00	122 0	0.00
37 35	59.00	122 0	36.00	37 51	26.00	122 3	0.00
37 36	4.00	122 0	41.00	38 35	0.00	122 40	0.00
38 35	0.00	122 40	0.00	37 36	4.00	122 0	41.00
37 30	0.00	122 40	0.00	37 35	59.00	122 0	36.00
37 30	0.00	122 29	45.00	37 31	31.69	121 56	43.00
37 19	12.00	122 8	45.00	37 31	21.90	121 56	33.00
37 4	0.00	121 50	0.00	37 18	51.39	121 45	16.09
36 57	0.00	121 38	0.00	DISTRICT 9			
DISTRICT 4				DISTRICT BOUNDARY			
DISTRICT BOUNDARY				39 0	0.00	118 35	0.00
36 43	0.00	121 16	0.00	38 0	0.00	118 32	0.00
36 49	0.00	121 23	0.00	37 51	0.00	118 24	0.00
36 51	0.00	121 25	0.00	37 42	0.00	118 23	0.00
36 54	0.00	121 27	0.00	37 32	0.00	118 21	0.00
36 59	0.00	121 26	0.00	37 23	0.00	118 20	0.00
36 55	34.00	121 21	22.00	37 14	0.00	118 20	0.00
36 49	5.00	121 22	45.00	37 2	0.00	118 12	0.00
DISTRICT 5				36 56	0.00	118 11	0.00
DISTRICT BOUNDARY				36 41	0.00	118 5	0.00
38 35	0.00	122 40	0.00	36 30	0.00	118 5	0.00
37 51	26.00	122 3	0.00	36 30	0.00	116 20	0.00
37 49	4.00	122 0	0.00	39 0	0.00	116 20	0.00
37 41	7.00	121 55	45.00	DISTRICT 1			
37 35	36.00	121 52	30.00	DISTRICT BOUNDARY			
37 39	0.00	121 46	55.00	36 24	0.00	121 0	0.00
DISTRICT 6				36 43	0.00	121 16	0.00
DISTRICT BOUNDARY				36 49	0.00	121 23	0.00
37 35	36.00	121 52	30.00	36 51	0.00	121 25	0.00
37 34	52.00	121 51	59.00	36 54	0.00	121 27	0.00
37 34	44.00	121 51	52.00	36 59	0.00	121 26	0.00
37 27	0.00	121 48	45.00	37 5	0.00	121 32	0.00
37 17	34.00	121 42	25.69	37 17	34.00	121 42	25.69
37 0	0.00	121 28	45.00	37 20	0.00	121 42	25.69
36 59	0.00	121 26	0.00	37 34	44.00	121 51	52.00
36 55	33.00	121 22	0.00	37 34	52.00	121 51	59.00
36 49	5.00	121 22	45.00	37 35	36.00	121 52	30.00
36 43	0.00	121 16	0.00	37 39	0.00	121 46	55.00
36 38	0.00	121 8	0.00	37 42	25.00	121 41	20.00
36 28	50.00	121 5	0.00	37 32	0.00	121 21	0.00
36 30	0.00	121 0	0.00	37 11	0.00	121 14	0.00
36 10	0.00	120 46	0.00	36 55	0.00	121 0	0.00
36 10	0.00	119 10	0.00	38 10	0.00	120 45	0.00
36 30	0.00	119 20	0.00	38 45	0.00	120 0	0.00
37 42	0.00	120 30	0.00	38 45	0.00	119 0	0.00
37 42	25.00	121 41	20.00	37 50	0.00	119 0	0.00
37 39	0.00	121 46	55.00	DISTRICT 2			
DISTRICT BOUNDARY				DISTRICT BOUNDARY			
37 5	8.00	121 51	58.00	37 5	8.00	121 51	58.00
37 9	30.00	121 58	25.00	37 9	30.00	121 58	25.00
37 18	15.00	122 9	0.00	37 18	15.00	122 9	0.00
37 40	47.39	122 30	0.00	37 40	47.39	122 30	0.00
37 48	24.00	122 24	5.00	37 48	24.00	122 24	5.00
37 56	2.00	122 18	9.90	37 56	2.00	122 18	9.90
37 52	16.00	122 14	58.80	37 52	16.00	122 14	58.80
37 36	4.00	122 0	41.00	37 36	4.00	122 0	41.00
37 35	59.00	122 0	36.00	37 35	59.00	122 0	36.00
37 31	31.69	121 56	43.00	37 31	31.69	121 56	43.00
37 31	21.90	121 56	33.00	37 31	21.90	121 56	33.00
37 18	51.39	121 45	16.09	37 18	51.39	121 45	16.09
37 17	34.00	121 42	25.69	37 17	34.00	121 42	25.69
37 5	0.00	121 32	0.00	37 5	0.00	121 32	0.00
36 59	0.00	121 26	0.00	36 59	0.00	121 26	0.00
36 54	0.00	121 27	0.00	36 54	0.00	121 27	0.00
36 51	0.00	121 25	0.00	36 51	0.00	121 25	0.00
36 49	0.00	121 23	0.00	36 49	0.00	121 23	0.00
36 43	0.00	121 16	0.00	36 43	0.00	121 16	0.00
36 28	50.00	121 5	0.00	36 28	50.00	121 5	0.00

(continued)

Table C.4.--Continued

San Francisco region				Sacramento region							
DISTRICT 5				DISTRICT 9				REGION BOUNDARY			
DISTRICT BOUNDARY				DISTRICT BOUNDARY				38 0 118 30			
37 17	34.00	121	42 25.69	37 42	25.00	121 41	20.00	38 0 121 30			
37 18	51.39	121	45 16.09	37 43	2.00	121 41	27.00	39 0 121 30			
37 31	21.90	121	56 33.00	38 0	0.00	121 59	4.29	39 0 118 30			
37 31	31.69	121	56 43.00	38 6	55.00	122 7	0.00	DISTRICTS *****			
37 35	59.00	122	0 36.00	38 14	26.00	122 8	27.00	DISTRICT 1			
37 36	4.00	122	0 41.00	38 24	17.00	122 8	7.00	DISTRICT BOUNDARY			
37 52	16.00	122	14 58.80	38 35	35.00	122 16	0.00	37 30	0.00	120	0 0.00
37 56	2.00	122	18 9.90	38 45	0.00	122 16	0.00	37 23	0.00	118	20 0.00
37 56	7.00	122	18 17.00	38 45	0.00	120 55	0.00	37 32	0.00	118	21 0.00
37 58	52.00	122	22 14.00	38 35	35.00	120 55	0.00	37 42	0.00	118	23 0.00
38 6	55.00	122	24 25.00	36 55	0.00	121 0	0.00	37 51	0.00	118	24 0.00
38 3	23.00	122	13 32.00	37 11	0.00	121 14	0.00	38 0	0.00	118	32 0.00
38 0	46.00	122	10 31.00	37 32	0.00	121 21	0.00	39 30	0.00	119	30 0.00
37 55	2.00	122	4 24.00					40 0	0.00	119	30 0.00
37 53	12.00	122	3 45.00					40 0	0.00	120	51 15.00
37 48	51.00	121	58 54.00					39 6	38.00	120	51 15.00
37 41	7.00	121	55 45.00					38 58	25.00	120	47 50.00
37 35	36.00	121	52 30.00					38 49	30.00	120	46 30.00
37 34	52.00	121	51 59.00					38 38	25.00	120	47 45.00
37 34	44.00	121	51 52.00					38 25	25.00	120	48 10.00
37 20	0.00	121	42 25.69					38 18	0.00	120	43 30.00
								37 53	0.00	120	21 15.00
DISTRICT 6											
DISTRICT BOUNDARY								DISTRICT 2			
39 4	0.00	123	25 0.00					DISTRICT BOUNDARY			
38 35	35.00	122	49 50.00					37 0	0.00	123	0 0.00
38 26	12.00	122	42 23.00					37 0	0.00	120	0 0.00
38 20	35.00	122	35 40.00					37 30	0.00	120	0 0.00
38 19	46.00	122	35 40.00					37 53	0.00	120	21 15.00
38 13	43.00	122	31 18.00					38 18	0.00	120	43 30.00
38 6	55.00	122	24 25.00					38 25	25.00	120	46 10.00
38 3	23.00	122	13 32.00					38 38	25.00	120	47 45.00
38 0	46.00	122	10 31.00					38 49	30.00	120	48 30.00
37 55	2.00	122	4 24.00					38 58	25.00	120	47 50.00
37 53	12.00	122	3 45.00					39 6	38.00	120	51 15.00
37 48	51.00	121	58 54.00					40 0	0.00	120	51 15.00
37 41	7.00	121	55 45.00					40 0	0.00	123	0 0.00
37 35	36.00	121	52 30.00								
37 39	0.00	121	46 55.00					DISTRICT 3			
37 42	25.00	121	41 20.00					DISTRICT BOUNDARY			
37 43	2.00	121	41 27.00					37 23	0.00	118	20 0.00
38 0	0.00	121	59 4.29					37 23	0.00	117	0 0.00
38 6	55.00	122	7 0.00					40 0	0.00	119	30 0.00
38 14	26.00	122	8 27.00					40 0	0.00	119	30 0.00
38 24	17.00	122	8 7.00					39 30	0.00	119	30 0.00
38 35	35.00	122	16 0.00					38 0	0.00	118	32 0.00
39 11	29.00	122	29 49.00					37 51	0.00	118	24 0.00
39 12	0.00	120	0 0.00					37 42	0.00	118	23 0.00
39 30	0.00	120	0 0.00					37 32	0.00	118	21 0.00
39 30	0.00	123	25 0.00								
DISTRICT 7											
DISTRICT BOUNDARY											
38 35	35.00	123	15 0.00								
38 30	41.00	123	14 55.00								
38 19	0.00	123	2 21.00								
38 10	54.00	122	55 25.00								
38 3	8.00	122	48 11.00								
38 2	41.39	122	47 50.80								
38 2	40.50	122	47 50.39								
37 55	52.00	122	41 39.00								
37 40	47.39	122	30 0.00								
37 30	0.00	122	45 0.00								
37 30	0.00	123	15 0.00								
DISTRICT 8											
DISTRICT BOUNDARY											
37 5	8.00	121	51 58.00								
37 2	0.00	121	41 0.00								
36 57	0.00	121	38 0.00								
36 54	0.00	121	27 0.00								
36 51	0.00	121	25 0.00								
36 49	0.00	121	23 0.00								
36 43	0.00	121	16 0.00								
36 24	0.00	121	0 0.00								
36 45	10.00	121	23 17.00								
36 49	55.00	121	30 0.00								
36 55	51.00	121	38 35.10								

Table C.4.--Continued

Santa	Rosa	region	Ukiah	region
REGION BOUNDARY 38 0 121 30 38 0 124 0 39 0 124 0 39 0 121 30 DISTRICTS ***** DISTRICT 1 DISTRICT BOUNDARY 37 30 0.00 122 18 21.00 37 40 47.39 122 30 0.00 37 55 52.00 122 41 39.00 38 2 40.39 122 47 50.39 38 2 41.30 122 47 50.80 38 3 8.00 122 48 11.00 38 10 54.00 122 55 25.00 38 19 0.00 123 2 21.00 38 30 41.00 123 14 55.00 38 59 15.30 123 40 59.00 39 0 34.00 123 41 0.00 39 30 0.00 124 10 0.00 37 30 0.00 124 10 0.00 DISTRICT 2 DISTRICT BOUNDARY 37 30 0.00 122 18 21.00 37 40 47.39 122 30 0.00 37 55 52.00 122 41 39.00 38 2 40.39 122 47 50.39 38 2 41.30 122 47 50.80 38 3 8.00 122 48 11.00 38 10 54.00 122 55 25.00 38 19 0.00 123 2 21.00 38 20 48.00 122 55 47.60 38 18 52.70 122 47 21.00 38 20 35.00 122 35 40.00 38 19 46.00 122 35 40.00 38 13 43.00 122 31 18.00 38 6 55.00 122 24 25.00 37 58 52.00 122 22 14.00 37 56 7.00 122 18 17.00 37 56 2.00 122 18 10.00 37 50 1.00 122 12 55.00 37 30 0.00 121 54 0.00 DISTRICT 3 DISTRICT BOUNDARY 38 19 0.00 123 2 21.00 38 30 41.00 123 14 55.00 38 59 15.30 123 40 59.00 39 0 34.00 123 41 0.00 39 30 0.00 124 10 0.00 39 32 0.00 123 45 0.00 39 4 0.00 123 25 0.00 38 56 55.00 123 13 0.00 38 45 21.00 123 0 44.00 38 39 29.00 122 52 15.00 38 35 35.00 122 49 50.00 38 26 12.00 122 42 23.00 38 20 35.00 122 35 40.00 38 18 52.70 122 47 21.00 38 20 48.00 122 55 47.60 DISTRICT 4 DISTRICT BOUNDARY 37 30 0.00 121 54 0.00 37 50 1.00 122 12 55.00 37 56 2.00 122 18 10.00 37 56 7.00 122 18 17.00 37 58 52.00 122 22 14.00 38 6 55.00 122 24 25.00 38 3 23.00 122 13 32.00 38 0 46.00 122 10 31.00 37 55 2.00 122 4 24.00 37 53 12.00 122 3 45.00 37 48 51.00 121 58 54.00 37 42 25.00 121 55 30.00 37 30 0.00 121 48 4.00 DISTRICT 5 DISTRICT BOUNDARY 38 3 23.00 122 13 32.00 38 6 55.00 122 24 25.00 38 13 43.00 122 31 18.00 38 19 46.00 122 35 40.00 38 20 35.00 122 35 40.00 38 26 12.00 122 42 23.00 38 35 35.00 122 49 50.00 38 45 0.00 122 50 0.00 38 45 0.00 122 43 0.00 38 51 0.00 122 43 0.00 38 51 0.00 122 50 0.00 38 45 0.00 122 50 0.00 38 35 35.00 122 49 50.00 38 39 29.00 122 52 15.00 38 45 21.00 123 0 44.00 38 56 55.00 123 13 0.00 39 4 0.00 123 25 0.00 39 32 0.00 123 45 0.00 39 32 0.00 122 45 0.00 39 11 29.00 122 29 49.00 38 39 16.00 122 17 59.00 38 24 17.00 122 8 7.00 38 14 26.00 122 8 27.00 38 6 55.00 122 7 0.00 38 2 26.00 122 7 5.00 DISTRICT 6 DISTRICT BOUNDARY 37 30 0.00 121 48 4.00 37 42 25.00 121 55 30.00 37 48 51.00 121 58 54.00 37 53 12.00 122 3 45.00 37 55 2.00 122 4 24.00 38 0 46.00 122 10 31.00 38 3 23.00 122 13 32.00 38 2 26.00 122 7 5.00 38 6 55.00 122 7 0.00 38 14 26.00 122 8 27.00 38 24 17.00 122 8 7.00 38 31 55.00 121 48 3.00 38 21 31.00 121 47 7.00 38 9 9.00 121 40 6.00 37 52 36.00 121 36 0.00 37 46 0.00 121 33 11.00 37 32 0.00 121 21 0.00 37 30 0.00 121 20 0.00 DISTRICT 7 DISTRICT BOUNDARY 38 24 17.00 122 8 7.00 38 39 16.00 122 17 59.00 39 11 29.00 122 29 49.00 39 32 0.00 122 45 0.00 39 30 0.00 121 0 0.00 39 30 0.00 120 0 0.00 38 41 0.00 120 0 0.00 38 41 0.00 121 0 0.00 38 37 5.00 121 17 46.00 38 34 50.00 121 29 35.00 38 31 55.00 121 48 3.00 DISTRICT 8 DISTRICT BOUNDARY 38 51 0.00 122 50 0.00 38 45 0.00 122 50 0.00 38 45 0.00 122 43 0.00 38 51 0.00 122 43 0.00 DISTRICT 9 DISTRICT BOUNDARY 38 41 0.00 121 0 0.00 38 37 5.00 121 17 46.00 38 34 50.00 121 29 35.00 38 31 55.00 121 48 3.00 38 21 31.00 121 47 7.00 38 9 9.00 121 40 6.00 37 52 36.00 121 36 0.00 37 46 0.00 121 33 11.00 37 32 0.00 121 21 0.00 37 30 0.00 121 20 0.00 37 30 0.00 119 50 0.00 38 50 0.00 119 50 0.00 REGION BOUNDARY 39 0 120 0 39 0 124 30 40 0 124 30 40 0 120 0 DISTRICTS ***** DISTRICT 1 DISTRICT BOUNDARY 38 19 0.00 123 2 21.00 38 30 41.00 123 14 55.00 38 57 0.00 123 40 0.00 38 57 0.00 123 37 0.00 39 1 0.00 123 37 0.00 39 1 0.00 123 42 0.00 39 5 0.00 123 46 0.00 38 53 0.00 123 46 0.00 38 19 0.00 123 10 0.00 DISTRICT 2 DISTRICT BOUNDARY 40 30 0.00 122 15 0.00 40 15 0.00 121 25 0.00 40 0 0.00 121 5 50.00 39 57 0.00 121 2 30.00 39 56 15.00 120 57 35.00 39 52 30.00 120 49 50.00 39 49 35.00 120 45 0.00 39 44 50.00 120 36 5.00 39 38 50.00 120 27 30.00 39 34 30.00 120 21 20.00 39 30 50.00 120 16 20.00 39 23 30.00 120 4 50.00 39 22 15.00 120 0 0.00 39 1 17.00 119 28 0.00 38 40 0.00 119 0 0.00 40 30 0.00 119 0 0.00 40 50 0.00 121 57 0.00 DISTRICT 3 DISTRICT BOUNDARY 38 19 0.00 123 2 21.00 38 30 41.00 123 14 55.00 38 57 0.00 123 40 0.00 38 57 0.00 123 37 0.00 39 1 0.00 123 37 0.00 39 1 0.00 123 42 0.00 39 5 0.00 123 46 0.00 39 20 0.00 123 55 0.00 39 40 0.00 123 50 0.00 40 15 0.00 124 25 0.00 40 23 39.00 124 30 0.00 40 23 39.00 124 22 30.00 40 1 8.00 123 42 45.00 39 54 20.00 123 38 0.00 39 49 30.00 123 37 15.00 39 41 23.00 123 32 20.00 39 32 20.00 123 30 0.00 39 25 30.00 123 22 0.00 39 21 50.00 123 22 40.00 39 10 35.00 123 13 30.00 39 8 16.00 123 12 0.00 39 4 10.00 123 9 0.00 39 0 20.00 123 6 50.00 38 56 10.00 123 3 40.00 38 50 10.00 123 0 50.00 38 45 21.00 123 0 44.00 38 39 29.00 122 52 15.00 38 35 35.00 122 49 50.00 38 26 12.00 122 42 23.00 38 20 35.00 122 35 40.00 38 18 52.70 122 47 21.00 38 20 48.00 122 55 47.60 DISTRICT 4 DISTRICT BOUNDARY 38 45 0.00 122 50 0.00 38 45 0.00 122 43 0.00 38 51 0.00 122 43 0.00 DISTRICT 4 DISTRICT BOUNDARY 38 45 0.00 122 50 0.00 38 45 0.00 122 43 0.00 38 51 0.00 122 43 0.00 38 51 0.00 122 50 0.00				

(continued)

Table C.4.--Continued

Ukiah region		Redding region	
DISTRICT 5		DISTRICT 8	
DISTRICT BOUNDARY		DISTRICT BOUNDARY	
38 20 35.00	122 35 40.00	38 19 0.00	120 45 0.00
38 26 12.00	122 42 23.00	38 30 0.00	120 56 40.00
38 35 35.00	122 49 50.00	39 0 0.00	121 15 0.00
38 45 0.00	122 50 0.00	39 22 0.00	121 27 20.00
38 45 0.00	122 43 0.00	40 0 0.00	121 55 20.00
38 51 0.00	122 43 0.00	40 5 30.00	122 1 0.00
38 51 0.00	122 50 0.00	40 30 0.00	122 15 0.00
38 45 0.00	122 50 0.00	40 16 0.00	121 25 0.00
38 35 35.00	122 49 50.00	40 0 0.00	121 5 50.00
38 39 29.00	122 52 15.00	39 57 0.00	121 2 30.00
38 45 21.00	123 0 44.00	39 49 20.00	120 52 20.00
38 50 10.00	123 0 50.00	39 42 30.00	120 50 25.00
38 56 10.00	123 3 40.00	39 29 20.00	120 49 15.00
39 0 20.00	123 6 50.00	39 25 20.00	120 47 40.00
39 4 10.00	123 9 0.00	39 8 30.00	120 46 25.00
39 8 16.00	123 12 0.00	39 3 10.00	120 43 10.00
39 10 35.00	123 13 30.00	39 1 0.00	120 46 50.00
39 21 50.00	123 22 40.00	38 58 25.00	120 47 20.00
39 25 30.00	123 22 0.00	38 52 10.00	120 49 40.00
39 32 20.00	123 30 0.00	38 44 0.00	120 46 25.00
39 33 10.00	123 3 15.00	38 39 30.00	120 48 40.00
39 34 0.00	122 36 30.00	38 30 0.00	120 50 20.00
39 27 20.00	122 32 45.00	DISTRICT 9	
39 20 20.00	122 32 30.00	DISTRICT BOUNDARY	
39 11 29.00	122 29 49.00	38 19 0.00	120 45 0.00
38 39 16.00	122 17 59.00	38 30 0.00	120 50 20.00
38 24 17.00	122 8 7.00	38 39 30.00	120 48 40.00
38 19 0.00	122 8 10.00	38 44 0.00	120 46 25.00
DISTRICT 6		38 52 10.00	120 49 40.00
DISTRICT BOUNDARY		38 58 25.00	120 47 20.00
40 23 39.00	124 22 30.00	39 1 0.00	120 46 50.00
40 1 8.00	123 42 45.00	39 3 10.00	120 43 10.00
39 54 20.00	123 38 0.00	39 8 30.00	120 46 25.00
39 49 30.00	123 37 15.00	39 25 20.00	120 47 40.00
39 41 23.00	123 32 20.00	39 29 20.00	120 49 15.00
39 32 20.00	123 30 0.00	39 42 30.00	120 50 25.00
39 33 10.00	123 3 15.00	39 49 20.00	120 53 20.00
39 34 0.00	122 36 30.00	39 57 0.00	121 2 30.00
39 51 0.00	122 32 0.00	39 56 15.00	120 57 35.00
39 55 38.00	122 38 20.00	39 52 30.00	120 49 50.00
40 2 47.00	122 40 43.50	39 49 35.00	120 45 0.00
40 6 44.00	122 49 0.00	39 44 50.00	120 36 5.00
40 6 44.00	123 5 0.00	39 38 50.00	120 27 30.00
40 15 37.00	123 16 0.00	39 34 30.00	120 21 20.00
40 23 40.00	123 26 5.00	39 30 50.00	120 16 20.00
DISTRICT 7		39 23 30.00	120 4 50.00
DISTRICT BOUNDARY		39 22 15.00	120 0 0.00
38 19 0.00	122 8 10.00	39 1 17.00	119 28 0.00
38 24 17.00	122 8 7.00	38 40 0.00	119 0 0.00
38 39 10.00	122 17 59.00	38 19 0.00	119 0 0.00
39 11 29.00	122 29 49.00	DISTRICT 3	
39 20 20.00	122 32 30.00	DISTRICT BOUNDARY	
39 27 20.00	122 32 45.00	41 31 28.00	123 55 0.00
39 34 0.00	122 36 30.00	41 11 11.00	123 49 58.00
39 51 0.00	122 32 0.00	41 4 14.00	123 48 14.00
39 55 38.00	122 38 20.00	40 52 30.00	123 43 55.00
40 2 47.00	122 40 43.50	40 45 33.00	123 41 0.00
40 6 44.00	122 49 0.00	40 36 52.00	123 33 36.00
40 6 44.00	123 5 0.00	40 23 40.00	123 26 5.00
40 15 37.00	123 16 0.00	40 32 17.80	123 30 50.00
40 23 40.00	123 26 5.00	40 23 40.00	123 26 5.00
40 40 0.00	122 39 0.00	40 15 37.00	123 16 0.00
40 30 0.00	122 15 0.00	40 6 44.00	123 5 0.00
40 5 30.00	122 1 0.00	40 6 44.00	122 49 0.00
40 0 0.00	121 55 20.00	40 2 47.00	122 40 43.50
39 2 0.00	121 27 20.00	39 55 38.00	122 38 20.00
39 0 0.00	121 15 0.00	39 51 0.00	122 32 0.00
38 30 0.00	120 56 40.00	39 10 0.00	122 38 20.00
38 19 0.00	120 45 0.00	39 0 0.00	121 45 0.00
37 50 0.00	120 45 0.00	39 53 0.00	121 50 0.00
37 50 0.00	122 8 10.00	40 5 30.00	122 1 0.00
DISTRICT 4		40 30 0.00	122 15 0.00
DISTRICT BOUNDARY		40 37 0.00	122 15 0.00
39 20 0.00	121 45 0.00	40 43 35.00	122 2 28.00
39 53 0.00	121 50 0.00	40 57 47.00	121 52 10.00
40 5 30.00	122 1 0.00	41 12 8.00	121 58 45.00
40 30 0.00	122 15 0.00	41 14 40.00	122 17 26.00
40 37 0.00	122 15 0.00	41 35 0.00	122 33 0.00
40 43 35.00	122 2 28.00	41 35 0.00	119 40 0.00
40 57 47.00	121 52 10.00	39 20 0.00	119 40 0.00

Table C.4.--Continued

Alturas region, CA	Fallon region, NV	Hilo region, HI
<p>REGION BOUNDARY</p> <p>41 0 120 0</p> <p>41 0 124 30</p> <p>42 0 124 30</p> <p>42 0 120 0</p> <p>DISTRICTS *****</p> <p>DISTRICT 1</p> <p>DISTRICT BOUNDARY</p> <p>40 30 0.00 123 30 0.00</p> <p>40 45 33.00 123 41 0.00</p> <p>40 52 30.00 123 43 55.00</p> <p>41 4 14.00 123 48 14.00</p> <p>41 11 11.00 123 49 58.00</p> <p>41 31 28.00 123 55 0.00</p> <p>41 42 11.00 124 1 0.00</p> <p>41 50 37.00 124 5 12.00</p> <p>42 11 52.00 124 0 24.00</p> <p>42 26 30.00 124 13 0.00</p> <p>42 30 0.00 124 25 0.00</p> <p>42 30 0.00 125 0 0.00</p> <p>40 30 0.00 125 0 0.00</p> <p>DISTRICT 2</p> <p>DISTRICT BOUNDARY</p> <p>40 30 0.00 123 30 0.00</p> <p>40 45 33.00 123 41 0.00</p> <p>40 52 30.00 123 43 55.00</p> <p>41 4 14.00 123 48 14.00</p> <p>41 11 11.00 123 49 58.00</p> <p>41 31 28.00 123 55 0.00</p> <p>41 42 11.00 124 1 0.00</p> <p>41 50 37.00 124 5 12.00</p> <p>42 11 52.00 124 0 24.00</p> <p>42 26 30.00 124 13 0.00</p> <p>42 30 0.00 124 25 0.00</p> <p>42 30 0.00 125 0 0.00</p> <p>42 45 0.00 125 0 0.00</p> <p>42 45 0.00 123 0 0.00</p> <p>42 20 20.00 123 0 0.00</p> <p>42 10 53.00 122 42 30.00</p> <p>41 54 19.00 122 32 0.00</p> <p>41 42 51.00 122 31 0.00</p> <p>41 30 0.00 122 33 0.00</p> <p>41 14 40.00 122 17 26.00</p> <p>41 12 8.00 121 58 45.00</p> <p>40 57 47.00 121 52 10.00</p> <p>40 43 35.00 122 2 28.00</p> <p>40 37 0.00 122 15 0.00</p> <p>40 30 0.00 122 15 0.00</p> <p>DISTRICT 3</p> <p>DISTRICT BOUNDARY</p> <p>40 30 0.00 122 15 0.00</p> <p>40 37 0.00 122 15 0.00</p> <p>40 43 35.00 122 2 28.00</p> <p>40 57 47.00 121 52 10.00</p> <p>41 12 8.00 121 58 45.00</p> <p>41 14 40.00 122 17 26.00</p> <p>41 30 0.00 122 33 0.00</p> <p>41 42 51.00 122 31 0.00</p> <p>41 54 19.00 122 32 0.00</p> <p>42 10 53.00 122 42 30.00</p> <p>42 20 20.00 123 0 0.00</p> <p>42 45 0.00 123 0 0.00</p> <p>42 45 0.00 121 30 0.00</p> <p>42 19 35.00 121 16 3.00</p> <p>42 0 24.00 121 6 51.00</p> <p>41 41 39.00 120 54 7.00</p> <p>41 24 51.00 120 54 42.00</p> <p>41 9 25.00 120 32 4.00</p> <p>40 43 49.00 120 11 49.00</p> <p>40 30 0.00 120 5 0.00</p> <p>DISTRICT 4</p> <p>DISTRICT BOUNDARY</p> <p>42 45 0.00 121 30 0.00</p> <p>42 19 35.00 121 16 3.00</p> <p>42 0 24.00 121 6 5.09</p> <p>41 41 39.00 120 54 7.00</p> <p>41 24 51.00 120 54 42.00</p> <p>41 9 25.00 120 32 4.00</p> <p>40 43 49.00 120 11 49.00</p> <p>40 30 0.00 120 5 0.00</p> <p>40 30 0.00 119 0 0.00</p> <p>42 45 0.00 119 0 0.00</p>	<p>REGION BOUNDARY</p> <p>38 45 119 0</p> <p>40 0 119 0</p> <p>40 0 117 0</p> <p>38 45 117 0</p> <p>DISTRICTS *****</p> <p>DISTRICT 1</p> <p>DISTRICT BOUNDARY</p> <p>37 45 0.00 120 0 0.00</p> <p>41 0 0.00 120 0 0.00</p> <p>41 0 0.00 115 20 0.00</p> <p>37 45 0.00 115 20 0.00</p>	<p>REGION BOUNDARY</p> <p>18 45 154 45</p> <p>18 45 156 15</p> <p>20 17 156 15</p> <p>20 17 154 45</p> <p>DISTRICTS *****</p> <p>DISTRICT 1</p> <p>DISTRICT BOUNDARY</p> <p>19 0 0.00 155 15 0.00</p> <p>19 12 50.17 155 22 7.00</p> <p>19 15 0.00 155 22 30.00</p> <p>19 16 11.07 155 22 35.00</p> <p>19 18 25.00 155 23 0.00</p> <p>19 21 20.00 155 16 45.09</p> <p>19 22 30.00 155 12 29.49</p> <p>19 27 0.00 154 56 50.00</p> <p>19 27 0.00 154 56 0.00</p> <p>19 25 45.00 154 54 0.00</p> <p>19 15 0.00 154 45 0.00</p> <p>19 15 0.00 154 0 0.00</p> <p>19 38 45.00 154 56 0.00</p> <p>19 36 0.00 155 3 30.00</p> <p>19 31 15.00 155 18 6.00</p> <p>19 29 30.00 155 23 15.00</p> <p>19 22 6.00 155 25 30.00</p> <p>19 14 27.00 155 27 10.00</p> <p>19 8 52.00 155 28 0.00</p> <p>DISTRICT 2</p> <p>DISTRICT BOUNDARY</p> <p>19 21 13.59 155 16 45.30</p> <p>19 21 13.75 155 16 45.30</p> <p>19 21 13.75 155 16 45.07</p> <p>19 21 13.59 155 16 45.07</p> <p>DISTRICT 3</p> <p>DISTRICT BOUNDARY</p> <p>19 22 28.10 155 12 29.50</p> <p>19 22 28.30 155 12 29.50</p> <p>19 22 28.30 155 12 29.40</p> <p>19 22 28.10 155 12 29.40</p> <p>DISTRICT 4</p> <p>DISTRICT BOUNDARY</p> <p>19 0 0.00 155 15 0.00</p> <p>19 12 50.17 155 22 7.00</p> <p>19 15 0.00 155 22 30.00</p> <p>19 16 11.07 155 22 35.00</p> <p>19 18 25.00 155 23 0.00</p> <p>19 21 20.00 155 16 45.09</p> <p>19 22 30.00 155 12 29.49</p> <p>19 27 0.00 154 56 50.00</p> <p>19 27 0.00 154 56 0.00</p> <p>19 25 45.00 154 54 0.00</p> <p>19 15 0.00 154 45 0.00</p> <p>19 0 0.00 154 45 0.00</p> <p>DISTRICT 5</p> <p>DISTRICT BOUNDARY</p> <p>19 38 45.00 154 56 0.00</p> <p>19 36 0.00 155 3 30.00</p> <p>19 31 15.00 155 18 6.00</p> <p>19 29 30.00 155 23 15.00</p> <p>19 22 6.00 155 25 30.00</p> <p>19 14 27.00 155 27 10.00</p> <p>19 8 52.00 155 28 0.00</p> <p>19 0 0.00 155 15 0.00</p> <p>18 0 0.00 155 15 0.00</p> <p>18 0 0.00 156 30 0.00</p> <p>21 0 0.00 156 30 0.00</p> <p>21 0 0.00 154 0 0.00</p>

APPENDIX D.--SECULAR STRAIN RATES

For a point P that is interior to the i-th district, let $\dot{\omega}$ denote the rotation rate at P in the clockwise direction and let $\dot{\epsilon}_{EE}$, $\dot{\epsilon}_{NN}$, and $\dot{\epsilon}_{EN}$ denote the components of the two-dimensional, second rank, strain rate tensor at P referred to a coordinate system whose positive coordinate axes are oriented in the east (E) and north (N) directions. Then using a spherical approximation of the Earth, the values of these parameters may be obtained from model parameters by the following relations:

$$\dot{\omega} = 0.5(a_{\phi\lambda}(i) \sec \tilde{\phi} - a_{\lambda\phi}(i) \cos \tilde{\phi}) \quad (D.1)$$

$$\dot{\epsilon}_{EE} = a_{\lambda\lambda}(i) \quad (D.2)$$

$$\dot{\epsilon}_{NN} = a_{\phi\phi}(i) \quad (D.3)$$

$$\dot{\epsilon}_{EN} = -0.5(a_{\phi\lambda}(i) \sec \tilde{\phi} + a_{\lambda\phi}(i) \cos \tilde{\phi}) \quad (D.4)$$

where positive strain corresponds to extension. In (D.1) and (D.4), $\tilde{\phi}$ represents the latitude at P. A single set of values for $\dot{\omega}$, $\dot{\epsilon}_{EE}$, $\dot{\epsilon}_{NN}$, and $\dot{\epsilon}_{EN}$ that is representative of the deformation throughout the i-th district is obtained by substituting an average value for the latitude in (D.1) and (D.4).

Auxiliary strain-rate parameters, derivable from the above tensor components, include

$$\dot{\Delta} = \dot{\epsilon}_{EE} + \dot{\epsilon}_{NN} \quad (D.5)$$

$$\dot{\gamma}_1 = \dot{\epsilon}_{EE} - \dot{\epsilon}_{NN} \quad (D.6)$$

$$\dot{\gamma}_2 = 2\dot{\epsilon}_{EN} \quad (D.7)$$

The parameter $\dot{\Delta}$, called the dilatation rate, measures the areal expansion rate of a district. The quantities $\dot{\gamma}_1$ and $\dot{\gamma}_2$ denote (right-lateral) shear-strain rates in specific directions. Shear strain for a given direction measures the change in magnitude of a right angle, the angle being considered in the clockwise sense, whose initial side coincides with this direction. The formula

$$\dot{\gamma}_\alpha = -(\dot{\gamma}_1 \sin 2\alpha + \dot{\gamma}_2 \cos 2\alpha) \quad (D.8)$$

gives the shear strain rate for the direction α as measured clockwise from north. Consequently, $\dot{\gamma}_1$ corresponds to the shear strain rate for the northwest or southeast direction, $\dot{\gamma}_2$ the rate for the east or west direction. The direction of maximum shear strain rate, denoted ψ and measured clockwise from north, is given by the relation

$$\tan 2\psi = \dot{\gamma}_1 / \dot{\gamma}_2 \quad (D.9)$$

The quantity

$$\dot{\gamma} = (\dot{\gamma}_1^2 + \dot{\gamma}_2^2)^{1/2} \quad (D.10)$$

represents the magnitude of shear strain rate for the direction ψ . The auxiliary parameters $\dot{\Delta}$, $\dot{\gamma}$, and ψ present the same information as that contained in the strain rate tensor. Numerical estimates for these quantities are thus tabulated in table D.1 along with estimates for the rotation rate ω . Numbers in parentheses denote formal standard errors.

Table D.1.--Secular strain-rate estimates

District	Rotation rate $\dot{\omega}$, $\mu\text{rad/yr}$	Dilatation rate $\dot{\Delta}$, $\mu\text{strain/yr}$	Max. shear rate $\dot{\gamma}$, $\mu\text{rad/yr}$	Direction of max. dextral shear, deg.
SAN DIEGO REGION				
I	0.064 (0.124)	-0.070 (0.036)	0.196 (0.035)	N 26 W (5)
II	0.201 (0.128)	0.393 (0.083)	0.966 (0.070)	N 59 W (2)
III	0.145 (0.127)	-0.031 (0.083)	0.455 (0.067)	N 45 W (5)
IV	0.117 (0.134)	-0.997 (0.151)	1.450 (0.113)	N 23 W (3)
V	0.099 (0.125)	-0.021 (0.045)	0.374 (0.042)	N 63 W (3)
VI*				
VII	0.106 (0.125)	-0.111 (0.045)	0.357 (0.053)	N 51 W (4)
VIII*				
IX	0.020 (0.126)	-0.140 (0.043)	0.190 (0.047)	N 59 W (8)
X	-0.046 (0.122)	-0.242 (0.038)	0.059 (0.027)	N 24 W (14)
SAN BERNARDINO REGION				
I*				
II	0.114 (0.066)	-0.174 (0.035)	0.164 (0.031)	N 48 W (5)
III	0.146 (0.069)	-0.111 (0.044)	0.276 (0.051)	N 57 W (5)
IV	0.220 (0.086)	-0.042 (0.122)	0.289 (0.113)	N 53 W (11)
V	0.179 (0.072)	-0.087 (0.044)	0.324 (0.051)	N 42 W (5)
VI	0.174 (0.072)	-0.171 (0.063)	0.100 (0.051)	N 51 W (16)
VII	0.107 (0.071)	-0.089 (0.056)	0.123 (0.050)	N 47 W (12)
VIII	-0.044 (0.075)	-0.089 (0.081)	0.033 (0.057)	N 44 E (51)
IX*				
X*				
XI	-0.088 (0.081)	-0.079 (0.102)	0.172 (0.045)	N 39 E (8)
BARSTOW REGION				
I	0.203 (0.125)	-0.067 (0.036)	0.386 (0.031)	N 63 W (3)
II	0.120 (0.124)	-0.117 (0.029)	0.215 (0.027)	N 54 W (4)
III	0.078 (0.124)	-0.191 (0.071)	0.248 (0.046)	N 62 W (5)
IV	1.360 (0.222)	1.130 (0.601)	2.960 (0.380)	N 74 E (5)
V	0.032 (0.123)	-0.210 (0.074)	0.122 (0.060)	N 36 W (15)
VI	0.026 (0.123)	0.423 (0.054)	0.128 (0.044)	N 75 W (8)
VII	0.006 (0.134)	-0.378 (0.124)	0.037 (0.044)	N 67 W (34)
VIII	0.062 (0.126)	-0.108 (0.059)	0.351 (0.050)	N 59 W (4)
IX	0.091 (0.137)	-0.431 (0.070)	0.434 (0.060)	N 31 W (4)
CHANNEL ISLANDS REGION				
I	0.167 (0.211)	0.514 (0.816)	0.269 (0.114)	N 41 W (12)
II	0.336 (0.146)	-0.039 (1.990)	0.423 (0.075)	N 37 W (5)
III*				
IV	-0.353 (0.225)	0.315 (0.797)	0.690 (0.136)	N 33 W (7)
V	0.111 (0.109)	-0.380 (0.240)	0.191 (0.105)	N 15 W (18)
VI	0.410 (0.130)	0.022 (0.463)	1.110 (0.141)	N 82 E (4)
VII	-0.223 (0.101)	-0.295 (0.241)	0.082 (0.039)	N 43 W (13)
VIII	0.136 (0.118)	-0.171 (0.346)	0.604 (0.244)	N 51 W (13)
IX	-0.033 (0.051)	-0.233 (0.043)	0.107 (0.026)	N 26 W (7)

* District contains an insufficient number of geodetic stations to compute parameters.

Table D.1.--Continued

LOS ANGELES REGION				
I	0.114 (0.057)	-0.023 (0.040)	0.122 (0.019)	N 42 W (5)
II	0.115 (0.219)	0.390 (0.915)	0.363 (0.592)	N 19 W (18)
III	0.088 (0.058)	-0.041 (0.045)	0.152 (0.034)	N 58 W (7)
IV	0.009 (0.062)	-0.439 (0.051)	0.239 (0.037)	N 23 W (4)
V	0.122 (0.060)	-0.087 (0.041)	0.303 (0.033)	N 64 W (4)
VI	0.005 (0.089)	-0.499 (0.122)	0.582 (0.121)	N 43 W (8)
VII	0.074 (0.111)	-0.319 (0.186)	0.151 (0.172)	N 24 W (39)
VIII	0.000 (0.064)	-0.033 (0.053)	0.076 (0.053)	N 80 W (19)
BAKERSFIELD REGION				
I	0.233 (0.050)	-0.070 (0.050)	0.181 (0.027)	N 56 W (5)
II	0.302 (0.054)	-0.115 (0.057)	0.236 (0.038)	N 48 W (5)
III	0.277 (0.054)	-0.301 (0.047)	0.341 (0.034)	N 58 W (3)
IV	0.320 (0.060)	-0.254 (0.047)	0.334 (0.038)	N 63 W (3)
V	0.180 (0.058)	-0.314 (0.041)	0.462 (0.035)	N 51 W (2)
VI	0.196 (0.058)	-0.105 (0.053)	0.382 (0.038)	N 61 W (3)
VII	0.231 (0.056)	-0.235 (0.036)	0.329 (0.038)	N 48 W (3)
SIERRA NEVADA REGION				
I	0.382 (0.121)	-0.520 (0.283)	0.672 (0.216)	N 74 W (8)
II	0.136 (0.104)	-0.142 (0.117)	0.095 (0.069)	N 14 W (23)
III	0.187 (0.099)	-0.094 (0.081)	0.105 (0.057)	N 30 W (15)
IV	-0.109 (0.138)	0.039 (0.397)	0.890 (0.262)	N 76 W (6)
V	0.137 (0.100)	-0.042 (0.109)	0.332 (0.094)	N 51 W (8)
VI	0.076 (0.092)	-0.144 (0.086)	0.129 (0.072)	N 21 W (18)
VII	0.122 (0.092)	-0.222 (0.057)	0.180 (0.039)	N 32 W (6)
PARKFIELD REGION				
I	0.086 (0.132)	-0.204 (0.116)	0.150 (0.086)	N 22 W (14)
II	0.156 (0.122)	-0.126 (0.052)	0.106 (0.025)	N 31 W (7)
III	0.325 (0.125)	-0.170 (0.068)	0.331 (0.044)	N 48 W (4)
IV	0.217 (0.131)	-0.365 (0.122)	0.681 (0.114)	N 67 W (5)
V*				
VI	0.052 (0.135)	0.018 (0.133)	0.094 (0.130)	N 88 E (36)
VII	0.168 (0.131)	-0.330 (0.123)	0.069 (0.071)	N 36 W (25)
MONTEREY REGION				
I	0.246 (0.109)	-0.213 (0.032)	0.153 (0.026)	N 40 W (6)
II	0.532 (0.114)	-0.159 (0.067)	0.829 (0.070)	N 30 W (3)
III	0.312 (0.120)	-0.073 (0.098)	0.518 (0.101)	N 29 W (6)
IV	0.321 (0.111)	-0.146 (0.044)	0.283 (0.054)	N 33 W (5)
V	0.050 (0.198)	-2.650 (0.287)	1.090 (0.300)	N 31 E (9)
VI	0.242 (0.120)	-0.108 (0.129)	0.104 (0.053)	N 16 W (16)
YOSEMITE REGION				
I	-0.021 (0.160)	-0.070 (0.136)	0.094 (0.169)	N 31 W (60)
II	0.206 (0.116)	0.207 (0.081)	0.948 (0.074)	N 38 W (3)
III	0.068 (0.122)	-0.048 (0.106)	0.359 (0.095)	N 53 W (10)
IV*				
V	0.191 (0.108)	-0.031 (0.094)	0.222 (0.069)	N 12 E (9)
VI	0.128 (0.106)	-0.208 (0.033)	0.197 (0.039)	N 27 W (5)
VII	0.044 (0.110)	-0.043 (0.046)	0.045 (0.044)	N 37 W (29)
VIII*				
IX	0.099 (0.120)	-0.244 (0.080)	0.090 (0.066)	N 25 E (23)

* District contains an insufficient number of geodetic stations to compute parameters.

Table D.1.--Continued

SAN FRANCISCO REGION				
I	-0.024 (0.077)	-0.170 (0.055)	0.200 (0.047)	N 59 W (9)
II	0.067 (0.075)	-0.160 (0.048)	0.321 (0.044)	N 38 W (4)
III	0.102 (0.084)	-0.371 (0.078)	0.652 (0.047)	N 35 W (2)
IV	-0.087 (0.084)	-0.231 (0.067)	0.020 (0.081)	N 76 E (100)
V	0.061 (0.084)	-0.088 (0.089)	0.271 (0.088)	N 35 W (10)
VI	0.038 (0.078)	0.045 (0.064)	0.238 (0.055)	N 53 W (7)
VII	0.155 (0.092)	0.288 (0.155)	0.509 (0.071)	N 41 W (4)
VIII*				
IX	-0.048 (0.110)	-0.426 (0.141)	0.321 (0.141)	N 11 E (14)
SACRAMENTO REGION				
I	0.606 (0.519)	-0.077 (0.128)	0.089 (0.054)	N 58 E (16)
II	0.619 (0.513)	0.107 (0.159)	0.218 (0.050)	N 40 W (8)
III	0.740 (0.526)	-0.085 (0.051)	0.125 (0.062)	N 02 W (16)
SANTA ROSA REGION				
I	0.356 (0.116)	-0.079 (0.094)	0.591 (0.063)	N 42 W (4)
II	0.379 (0.112)	-0.234 (0.063)	0.514 (0.034)	N 32 W (2)
III	0.377 (0.116)	0.027 (0.106)	0.676 (0.085)	N 52 W (4)
IV	0.334 (0.126)	-0.535 (0.185)	0.360 (0.150)	N 03 E (11)
V	0.333 (0.113)	-0.218 (0.046)	0.215 (0.036)	N 33 W (5)
VI	0.249 (0.118)	-0.262 (0.097)	0.287 (0.079)	N 35 W (7)
VII	0.340 (0.124)	-0.277 (0.115)	0.078 (0.068)	N 29 W (23)
VIII	-0.385 (0.306)	7.090 (0.407)	1.890 (0.440)	N 22 E (9)
IX	0.288 (0.126)	-0.301 (0.163)	0.268 (0.118)	N 30 W (14)
UKIAH REGION				
I	1.260 (0.758)	0.711 (0.881)	1.290 (0.118)	N 32 W (3)
II	0.151 (0.126)	-0.015 (0.179)	0.101 (0.103)	N 76 E (36)
III	0.556 (0.186)	-0.969 (0.402)	0.887 (0.270)	N 01 E (9)
IV	-0.385 (0.306)	7.090 (0.407)	1.890 (0.440)	N 22 E (9)
V	0.254 (0.112)	-0.202 (0.056)	0.396 (0.076)	N 27 W (5)
VI	0.303 (0.187)	0.429 (0.363)	0.214 (0.093)	N 58 W (13)
VII	0.158 (0.102)	0.053 (0.103)	0.068 (0.034)	N 12 W (15)
VIII	0.151 (0.104)	0.082 (0.123)	0.099 (0.065)	N 58 E (19)
IX	0.210 (0.133)	0.121 (0.202)	0.063 (0.137)	N 11 E (66)
REDDING REGION				
I	0.204 (0.313)	-0.494 (0.760)	1.250 (0.134)	N 10 W (3)
II	-0.157 (0.302)	0.575 (0.288)	0.136 (0.092)	N 24 W (20)
III	-0.164 (0.306)	0.487 (0.230)	0.070 (0.045)	N 58 W (18)
IV	-0.121 (0.308)	0.382 (0.230)	0.029 (0.049)	N 52 E (51)
ALTURAS REGION				
I	0.002 (0.073)	0.084 (0.135)	0.043 (0.140)	N 21 E (104)
II	-0.014 (0.039)	0.058 (0.078)	0.058 (0.068)	N 88 E (30)
III	0.011 (0.039)	-0.079 (0.079)	0.038 (0.059)	N 56 W (40)
IV	0.001 (0.077)	-0.063 (0.135)	0.027 (0.074)	N 20 E (96)
FALLON REGION (NEVADA)				
I	0.105 (0.270)	-0.300 (0.221)	0.117 (0.054)	N 31 W (14)
HILO REGION (HAWAII)				
I	-0.10 (1.97)	1.56 (0.29)	0.98 (0.23)	N 07 W (7)
II*				
III*				
IV	1.82 (2.01)	0.60 (0.71)	0.61 (0.84)	N 13 W (34)
V	0.22 (2.00)	0.44 (0.31)	0.12 (0.15)	N 36 E (36)

* District contains an insufficient number of geodetic stations to compute parameters.

APPENDIX E.--ANCHORAGE REGION

The crustal motion model for Alaska's Anchorage region differs in form from models for the other 18 REDEAM regions. The Anchorage model was developed to characterize horizontal displacements associated with the 1964 Prince William Sound earthquake ($M=8.4$). At some locations, these displacements exceed 10 m in magnitude (fig. E.1).

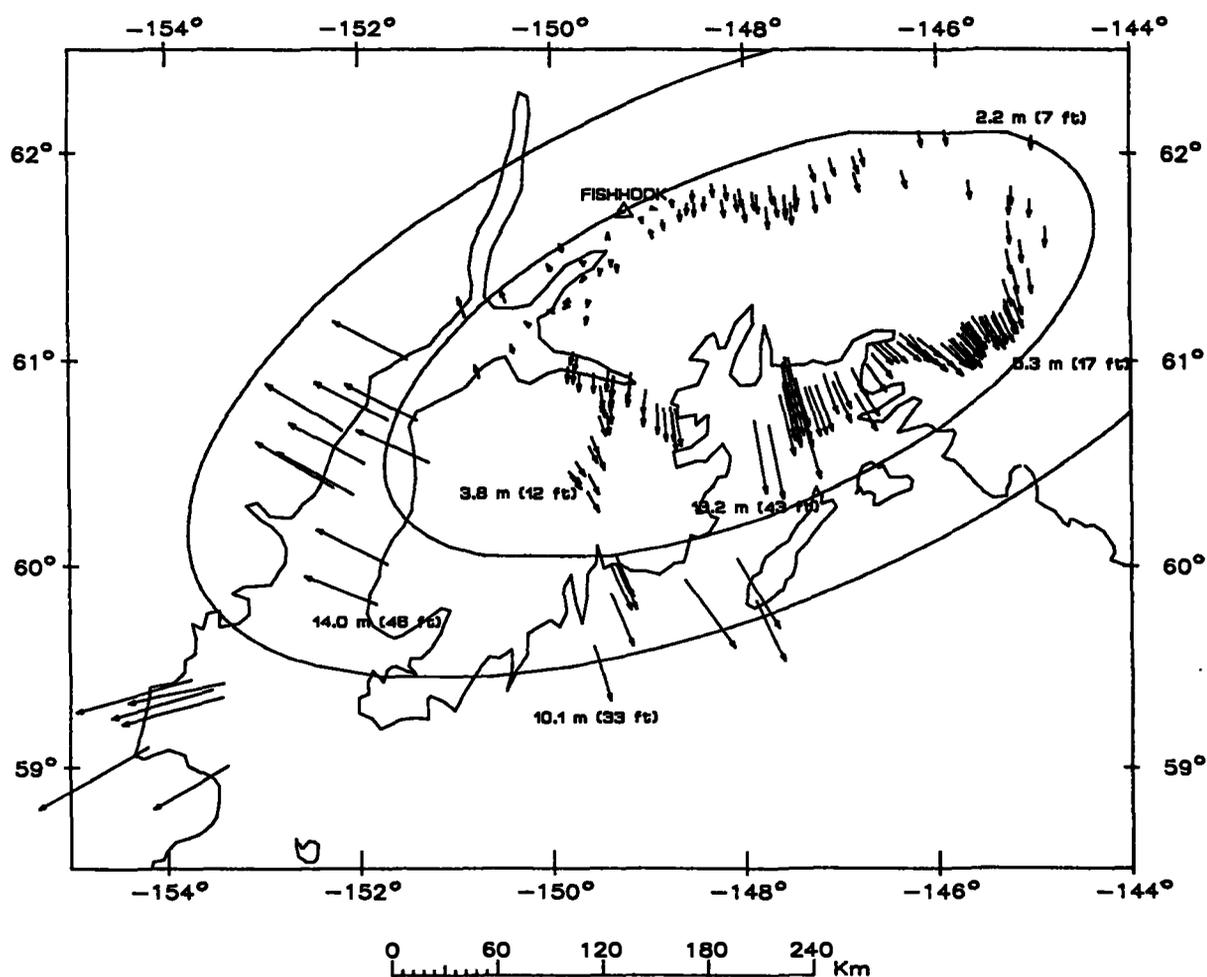


Figure E.1.--Displacement vectors for 1964 Prince William Sound earthquake as derived by differencing pre- and post-earthquake positions. Vectors are relative to the assumption of no motion at station FISHHOOK. Vectors west of the 151°W meridian are highly suspect because of poor network geometry. Displacement magnitudes are given for selected stations. These vectors were used to estimate polynomial coefficients for our model. Ellipses delineate the area where the polynomials are scaled from their full value to zero as described in the text.

The dislocation formulation that was used to model earthquakes in California and Nevada proved relatively unsatisfactory for the Prince William Sound earthquake. Our colleague, Steven Musman, found that the associated coseismic displacements were better represented by second order polynomials in two variables. With the polynomial model, the eastward displacement d_x and the northward displacement d_y at a point is expressed as

$$\begin{aligned}d_x &= a_0 + a_1x + a_2y + a_3x^2 + a_4xy + a_5y^2 \\d_y &= b_0 + b_1x + b_2y + b_3x^2 + b_4xy + b_5y^2\end{aligned}\tag{E.1}$$

where x and y denote coordinates of the point east and north of some reference origin, and a_i and b_i (for $i = 0,1,2,\dots,5$) are parameters to be estimated from the data. Indeed, when observations (directions, distances, and azimuths) in the vicinity of the earthquake were updated to a common time (see chapter 5) using the best polynomial model, a network adjustment of these updated observations yielded a variance of unit weight that is five times better than the variance of unit weight for a network adjustment of the same observations when updated by our best dislocation model.

Coefficients for the polynomial model were estimated using the positional coordinates and displacements for 198 geodetic stations (table E.1). These displacements were computed by differencing two sets of horizontal positions. The first set was obtained by a free adjustment of pertinent pre-earthquake horizontal data, and the second set by a free adjustment of pertinent postearthquake horizontal data. For a yet to be explained reason, these displacement vectors differ from those published by Parkin (1969) as presented in fig. E.2.

The following FORTRAN routine documents the derived polynomial coefficients. This routine may be used to calculate modeled displacements for the 1964 Alaska earthquake. The definition of two concentric ellipses (fig. E.1) is embedded within this routine. Inside the inner ellipse modeled displacements are as given by the polynomial expressions. For modeled displacements between the two ellipses, values from the polynomial expressions are multiplied by a function that decreases monotonically from 1 to 0 along any line segment starting at the perimeter of the inner ellipse and ending on the perimeter of the outer ellipse. Outside the outer ellipse, modeled displacements are defined as zero.

Table E.1.--Coseismic displacements for Prince William Sound earthquake

Station	Latitude		Longitude		Displacement	
	deg	min	deg	min	East, m	North, m
ANCHOR POINT 1908	59	48.8	151	49.7	-13.05	4.94
ANCHOR 1947	60	53.1	147	15.1	2.95	-7.74
AUDRY 1961	60	30.8	151	16.6	-13.23	5.74
AXIS USE 1942	60	27.1	149	36.6	1.78	-3.30
BAHIA 1947	60	55.0	147	27.8	2.42	-8.06
BAKER 1905	59	51.7	149	23.3	3.98	-9.06
BEAR USE 1942	60	52.9	149	29.7	.56	-3.63
BELUGA 1909	61	12.7	150	53.5	-1.27	3.19
BENCH 1941	61	5.3	145	45.3	3.92	-4.26
BERNARD 1941	61	34.3	145	8.4	.42	-3.74
BIRCH HILL 1942	60	55.3	150	45.0	-.86	2.07
BIRCH 1947	60	59.5	147	36.0	1.53	-7.41
BIRD POINT 1912	60	55.6	149	21.7	-.32	-3.30
BLUFF 1905	60	.5	149	20.0	3.61	-7.92
BONE USE 1942	60	22.1	149	37.9	1.94	-3.38
BOU 1944	61	50.8	147	43.5	.58	-2.93
BOULDER 1914	60	46.4	148	50.3	.49	-5.57
BRUIN 1941	61	10.8	145	43.8	3.42	-4.30
BUNCH 1901	61	4.9	146	39.9	3.56	-3.37
BURR 1913	59	25.1	153	25.2	-17.46	-3.53
BUSBY 1942	60	53.8	146	48.9	3.67	-5.51
BUTTE 1908	60	40.4	152	10.6	-13.83	7.85
CAA 1944	61	47.3	147	40.3	.56	-3.37
CALL 1947	60	52.8	147	19.4	3.20	-8.17
CAPE 1912	60	58.1	149	46.0	-.17	-2.46
CASCADE 1944	61	49.8	148	2.6	.55	-2.83
CASTLE 1944	61	49.3	148	31.5	-.05	-1.70
CENTRAL 1912	60	51.9	149	1.4	-.52	-4.25
CHISWELL 1905	59	36.1	149	33.9	2.85	-9.72
CHUGACH 1941	61	13.7	145	24.8	2.62	-4.81
CLEO USE 1942	60	48.1	149	2.6	-.18	-3.94
COOPER USE 1942	60	28.0	149	49.0	2.09	-1.97
COST 1961	60	43.3	151	42.2	-13.35	6.46
CRAG USE 1944	61	40.3	148	50.3	.31	-1.60
CROMBIE 1941	61	4.4	146	12.9	4.23	-3.17
CUB USE 1942	60	51.0	149	29.1	.31	-3.37
DARK 1912	61	2.4	149	46.7	-.50	-1.99
DIVIDE 1941	61	7.8	145	46.4	3.78	-4.17
DIVIDE 1947	60	52.6	147	22.0	2.75	-8.40
DORF 1961	61	15.7	149	49.7	-.59	-.40
DROP 1941	61	10.9	145	39.8	3.26	-4.43
EKLUTNA 1922	61	27.5	149	19.0	.10	-1.20
ELBOW 1941	61	12.5	145	25.7	2.70	-4.58
ELF 1947	60	56.6	147	3.3	2.91	-5.96
ERMI USE 1942	60	49.8	149	28.9	.48	-3.38
ERNESTINE 1941	61	26.0	145	3.1	.63	-3.98
EVANS 1905	60	2.2	148	4.7	7.66	-12.19
EVEN 1947	60	58.7	147	33.1	1.79	-7.45
FAIR 1947	60	52.6	147	27.6	2.66	-8.92
FALL 1941	61	20.5	145	11.2	1.34	-4.29
FALLS USE 1942	60	52.9	149	22.3	-.05	-3.47
FINSKI 1947	60	54.0	147	4.5	3.17	-6.73
FIRE 1944	61	58.9	146	51.4	.90	-2.63

Table E.1.--Coseismic displacements for Prince William Sound earthquake (cont'd)

Station	Latitude		Longitude		Displacement	
	deg	min	deg	min	East, m	North, m
FISHOOK 1944	61	43.0	149	14.0	0.00	0.00
FISHLAKE 1922	61	31.5	149	52.8	-.63	1.13
FLIENT 1947	60	56.1	147	10.3	2.57	-6.54
FLOW 1901	60	57.7	146	52.8	2.99	-4.50
FOOT 1941	61	10.9	145	32.3	3.14	-4.59
FORKS 1941	61	15.4	145	16.1	2.04	-4.49
FOSSIL 1944	61	16.8	149	36.6	-.29	-.81
GILPATRICK USE 1942	60	35.2	149	36.8	1.54	-2.70
GLACIER 1944	61	44.6	147	45.7	.48	-3.80
GLEN EAST BASE 1944	61	49.0	147	28.0	.38	-3.42
GLOBE B I E USE 1961	61	17.0	149	49.4	-.60	-.30
GOOSE GFLAT 1909	61	.6	151	30.0	-13.36	6.36
GRANT 1947	60	54.1	147	33.0	2.28	-8.90
GULL ROCK 2 USE 1942	60	58.2	149	49.8	-.13	-2.38
HARRIET 1908	60	23.5	152	16.1	-14.17	7.94
HELD 1901	61	7.7	146	22.7	3.88	-2.87
HELEN 1961	60	43.2	151	24.3	-13.03	6.27
HICKS 1944	61	48.9	147	54.6	.51	-3.00
HOG BACK 1941	61	2.4	145	58.9	4.46	-3.62
HOPE USE 1942	60	56.3	149	42.8	-.02	-2.66
HORN 1944	62	1.0	146	47.1	.80	-2.57
HUDSON 1941	61	52.1	145	40.3	.42	-3.00
INDIA 1947	60	52.6	147	33.1	2.39	-9.52
INLET 1947	60	56.8	147	32.2	2.05	-7.99
ISLE 1912	61	1.0	149	44.1	-.50	-2.28
JACK 1901	61	1.9	146	40.3	3.62	-3.79
JETTY 1947	61	.9	147	32.9	1.81	-6.42
JOHN USE 1942	60	34.2	149	31.6	1.30	-3.30
JONAH 1947	61	1.1	147	35.0	1.83	-6.92
KALGIN 1944	60	21.5	152	4.0	-13.81	7.43
KAMISHAK 1946	59	20.8	153	26.0	-18.31	-4.74
KENNY 1941	61	46.6	145	2.1	0.00	-3.12
KEY 1941	61	4.3	145	51.7	4.26	-3.90
KING 1944	61	45.6	148	30.5	.06	-2.28
KINIK 1947	60	50.3	147	37.6	2.62	-10.42
KLAWASI 1941	62	4.9	145	.4	-.09	-2.20
KNIFE 1941	61	6.5	146	11.9	4.12	-3.02
KNOB 1928	59	56.0	148	37.1	9.12	-11.97
KNOWLES 1941	61	15.3	145	18.0	2.12	-4.49
LAZY USE 1944	61	37.4	148	57.9	.39	-1.24
LEAN USE 1942	60	30.8	149	44.8	1.65	-2.21
LEILA 1944	61	48.8	147	16.7	.78	-3.53
LINA 1944	61	54.2	146	51.1	1.04	-2.98
LITTLE TONSINA 1941	61	31.6	145	16.8	.90	-3.88
LONG 1944	61	25.9	149	59.8	-.76	.93
LOON 1944	61	47.5	148	24.4	-.02	-2.22
LOWE 1941	61	3.1	146	8.1	4.28	-3.31
MAC 1941 RM 3	61	14.3	149	59.1	-.73	-.01
MAT 1944	61	46.5	148	13.3	.51	-2.92
MC NEIL 1946	59	6.1	154	12.0	-20.02	-11.13
MIDDLE 2 1927	60	3.4	149	20.5	3.25	-7.38
MISERY 3 1944	61	16.6	150	28.2	-.99	2.04

Table E.1.--Coseismic displacements for Prince William Sound earthquake (cont'd)

Station	Latitude		Longitude		Displacement	
	deg	min	deg	min	East, m	North, m
MOOSE USE 1942	60	45.0	149	25.7	.53	-3.49
MOOSE 1944	61	40.6	149	3.2	.26	-.62
NEL 1944	61	51.4	147	9.3	.90	-3.27
NELSON USE 1942	60	48.3	149	23.8	.39	-3.65
NINILCHIK 19082	60	.6	151	42.8	-13.02	6.13
NORTH KALGIN 1908	60	30.5	151	56.7	-13.82	7.12
NUSKA 1944	61	45.6	147	30.5	.37	-3.76
O BRIEN 1944	61	45.8	148	0.0	.54	-3.32
ODESSEY 1941	61	8.5	145	41.7	3.53	-4.37
OLSEN 1947	60	51.8	147	33.1	2.52	-9.77
PARK 1944	61	45.6	148	34.9	-.12	-1.98
PELLEW 1947	61	5.1	146	36.4	3.73	-3.24
PERRY 1912	60	43.1	147	53.9	2.37	-12.60
PETERS EAST BASE 1922	61	28.2	149	22.5	.04	-.72
PETERS WEST BASE 1922	61	25.7	149	29.3	-.12	-.72
POINT 1912	60	57.1	149	24.6	-.46	-3.17
POSSESSION 1909	61	2.3	150	23.7	-.55	1.29
POWER 1941	61	5.1	146	18.2	4.06	-3.05
PTARMIGAN 1941	61	12.5	145	37.6	3.08	-4.32
PUDD 1944	61	51.2	148	19.7	.36	-2.00
PURIN 1944	61	48.3	148	5.2	.51	-2.84
QUOTE 1947	60	50.9	147	27.0	2.85	-9.43
RACE POINT 1909	61	10.1	150	13.4	-.88	.58
RAM USE 1944	61	43.5	148	39.4	.14	-1.96
RED 1944	61	48.2	147	51.9	.43	-3.21
REEF 1942	60	50.7	146	50.6	3.81	-6.21
RICE 1941	61	22.8	145	18.4	1.52	-4.25
RIGID 1947	60	58.8	147	36.5	1.62	-7.65
RIM 1941	61	12.7	145	29.7	2.86	-4.51
ROCK SPUR 1941	61	9.5	145	44.3	3.55	-4.31
ROSE 1914	61	28.4	149	40.8	-.35	.21
RUSH 1944	61	50.9	148	11.4	.48	-2.40
SALMON 1914	60	47.6	148	54.2	-.04	-4.96
SAW 1901	61	5.4	146	24.5	4.17	-3.05
SCARP 1950	61	38.2	144	52.4	-.16	-3.23
SECOND 1912	60	55.6	149	21.2	-.37	-3.29
SHAKESPEARE 1914	60	45.0	148	45.1	.87	-6.46
SHAW 1946	59	.5	153	22.4	-13.63	-7.70
SHEEP ASTRO AZ MK 1943	61	47.1	147	34.5	.46	-3.60
SHEEP ASTRO 1943	61	48.0	147	34.4	.46	-3.48
SHEPARD 1941	62	6.7	145	54.4	.54	-2.36
SHIP 1944	61	12.5	149	37.6	-.14	-1.15
SIDE 1901	60	58.9	146	43.6	3.50	-4.26
SIGHT 1944	61	50.6	147	27.9	.51	-3.19
SIWASH 1947	60	56.6	147	35.0	1.89	-8.28
SLIDE 1914	60	46.1	148	43.0	1.11	-6.28
SNOW 1941	61	4.9	145	56.6	4.32	-3.64
SPONGE 1941	61	4.4	146	4.1	4.32	-3.37
SQUARE 1941	61	11.3	145	28.5	3.25	-4.71
SQUAW 1944	61	56.4	147	18.2	.84	-2.60

Table E.1.--Coseismic displacements for Prince William Sound earthquake (cont'd)

Station	Latitude		Longitude		Displacement	
	deg	min	deg	min	East, m	North, m
STAIR 1905	59	49.9	147	53.1	5.20	-10.76
STEP 1913	59	26.1	153	45.7	-20.97	-5.85
STETSON USE 1942	60	26.8	149	47.9	1.96	-2.16
STONE 1941	61	5.0	145	49.3	4.10	-4.06
STOREY 1947	60	44.6	147	24.1	3.47	-11.05
STUART NORTH BASE 1941	61	16.4	145	16.7	1.99	-4.37
STUART 1941	61	16.8	145	12.1	1.74	-4.33
STUCK 1941	61	47.5	145	15.2	.12	-3.20
SUGAR 1941	61	5.0	146	17.0	4.07	-2.92
SUMMIT USE 1942	60	38.2	149	35.3	1.30	-2.89
SUMMIT 1941	61	26.0	145	12.7	1.03	-3.98
TAHNETA 1944	61	58.3	147	6.0	.94	-2.66
TAZ 1944	61	55.0	146	21.7	.96	-2.75
THOMPSON PASS 1941	61	7.7	145	43.8	3.69	-4.27
TIEKEL 1941	61	17.0	145	18.0	1.95	-4.39
TOLSONA 1944	62	6.3	146	10.3	.56	-2.50
TONSINA 1941	61	39.7	145	15.7	.30	-3.63
TRIB USE 1942	60	44.1	149	30.5	1.20	-3.15
TRIPLE 1912	60	53.4	149	11.2	-.48	-3.76
TSINA EAST BASE 1941	61	12.2	145	31.6	2.99	-4.49
TSINA WEST BASE 1941	61	12.1	145	33.0	3.03	-4.46
TSINA 1941	61	13.2	145	20.4	2.41	-4.64
TURN 1905	59	59.7	149	23.1	3.51	-7.51
UNAK 1947	60	53.6	147	31.0	2.44	-8.81
VALDEZ NORTH BASE 1901	61	7.4	146	16.8	3.25	-3.70
VISIT 1947	61	5.0	146	31.7	3.83	-3.12
VISTA 1914	60	46.7	148	45.9	.64	-5.79
WALKER USE 1942	60	51.5	149	23.2	.14	-3.38
WASILLA 1922	61	36.6	149	24.5	.05	.05
WAVER 1947	60	41.9	147	44.9	2.83	-12.92
WEST AUGUSTINE 1913	59	23.1	153	32.5	-18.34	-5.08
WHITNEY 1922	61	23.3	149	40.0	-.31	-.32
WILLOW CRK S BASE 1941	61	50.3	145	13.5	.12	-3.08
WINDY 1912	60	56.2	149	33.8	-.08	-2.93
WISH USE 1944	61	44.2	148	57.0	.97	-.26
WOOD USE 1941	60	56.4	149	10.6	-.79	-4.59
WORTHINGTON NE BASE 1941	61	10.6	145	41.2	3.39	-4.38
WORTHINGTON SW BASE 1941	61	9.8	145	42.4	3.44	-4.35
WORTMAN NE BASE 1941	61	6.4	145	50.0	4.05	-4.03
WORTMAN SW BASE 1941	61	6.0	145	51.9	4.08	-3.94
YEAST 1947	60	53.1	147	35.3	2.18	-9.51
YOUNG 1944	61	46.7	148	44.5	.22	-1.23
ZINC 1947	60	54.1	147	35.6	2.07	-9.16

SUBROUTINE INTER(XLA,XLO,DX,DY)

```

C
C This subroutine interpolates to find the displacement at
C an arbitrary position for the 1964 Alaska earthquake.
C INPUT:
C   XLA is the (north) latitude of the position in degrees.
C   XLO is the (west) longitude of the position in degrees.
C OUTPUT:
C   DX is the eastward displacement in meters at the position.
C   DY is the northward displacement in meters at the position.
C R quantifies the distance of the position from the centroid
C of the observations. In the region where R has a value
C between 1.8 and 2.8, interpolation is somewhat suspect, and
C all displacements in this region are reduced. All displacements
C in the region where R is greater than 2.8 are set to zero.
C
C
C IMPLICIT DOUBLE PRECISION (A-H,O-Z)
C DIMENSION CX(5), CY(5)
C PI = 3.14159265D0
C HALFPI = PI/2.D0
C XBAR = 58.5D0
C YBAR = 64.9D0
C XX = 3772.D0
C XY = 1202.D0
C YY = 1240.D0
C DELTA = XX*YY - XY*XY
C DX0 = 0.25D0
C DY0 = -3.60D0
C CX(1) = +.5153D-01
C CY(1) = -.4312D-01
C CX(2) = -.4471D-01
C CY(2) = +.6828D-01
C CX(3) = -.4415D-03
C CY(3) = +.6686D-03
C CX(4) = -.6772D-03
C CY(4) = -.6842D-03
C CX(5) = +.6147D-03
C CY(5) = -.3730D-03
C X = DCOS(60.D0*PI/180.D0)*60.D0*(150.D0 - XLO)
C X = X - XBAR
C Y = 60.D0*(XLA - 60.D0)
C Y = Y - YBAR
C R = X*X*YY - 2.D0*X*Y*XY + Y*Y*XX
C R = R/DELTA
C R = DSQRT(R)
C F = X*X - XX
C G = X*Y - XY
C H = Y*Y - YY
C DX = DX0 + CX(1)*X + CX(2)*Y + CX(3)*F + CX(4)*G + CX(5)*H
C DY = DY0 + CY(1)*X + CY(2)*Y + CY(3)*F + CY(4)*G + CY(5)*H
C IF(R.GE.1.8D0 .AND. R.LT.2.8D0)THEN
C   FACR = (DSIN((2.8D0 - R)*HALFPI))**2
C   DX = FACR*DX
C   DY = FACR*DY
C ELSEIF (R.GE.2.8D0) THEN
C   DX = 0.D0
C   DY = 0.D0
C ENDIF
C RETURN
C END

```

APPENDIX F.--HILO REGION

The Hilo region covers the Island of Hawaii (fig. F.1). This island is often referred to as the "Big Island" to distinguish it from the entire group of islands which is also known as Hawaii. The REDEAM model for the Hilo region differs in form from models for the other 18 regions. The dislocation approach used successfully for modeling earthquakes in California and Nevada proved relatively unsatisfactory for representing horizontal displacements associated with the 1975 Kalpana earthquake ($M=7.1$). We found that coseismic displacements for this Hawaiian event were better represented by second order polynomials in two variables. These polynomials are equivalent in form to those used to model displacements for the Prince William Sound earthquake of Alaska's Anchorage region (Appendix E). The Hilo model differs from the Anchorage model, however, in that the Hilo region contains sufficient data to enable a rough estimate of secular motion. These data include EDM measurements from a deformation network regularly monitored by the USGS since 1970.

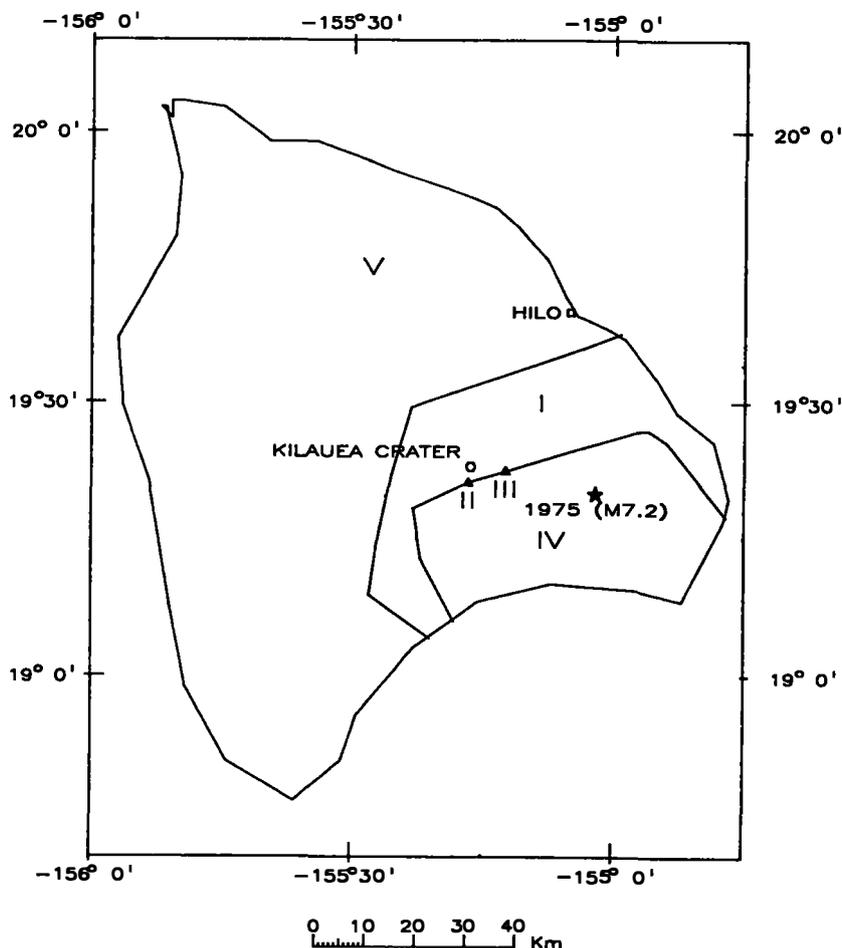


Figure F.1.--The Hilo region covers the Island of Hawaii. Roman numerals identify the various districts in the REDEAM model. The star represents the epicenter of the 1975 Kalpana earthquake.

Hawaii is the southeasternmost island of a 6,000 km long chain of islands and seamounts across the central Pacific Ocean. According to the theory of plate tectonics, these islands and seamounts were formed as a result of the relative northwesterly motion of the Pacific lithospheric plate over a thermal plume in the Earth's mantle (a "hot spot"). The theory contends that lavas of basalt composition emanate from the hot spot, forming volcanic islands or seamounts, which subsequently become inactive as plate motion continues. The northwestern end of this island-seamount chain is about 70 million years old, while the southeastern extremity of the chain, Hawaii's Kilauea volcano, is less than 1 million years in age.

Kilauea, like other volcanoes in the Hawaiian chain, is a "shield" volcano. Shield volcanoes are gently sloping and much broader than they are high. The extruded material from a shield volcano consists almost exclusively of lava flowing out in quiet eruptions from a central vent or from closely related fissures, called rifts, that extend laterally for several kilometers from the central vent. This type of volcano contrasts with the steeper "cone" volcanoes, like Mount Saint Helens, which are characterized by violent, explosive eruptions of cinders and pyroclastic materials which may or may not be accompanied by lava flows. The absence of violent eruptions, however, does not preclude the occurrence of destructive earthquakes near Kilauea as manifested by the 1975 Kalpana event on the south flank of the volcano. An earthquake of similar magnitude is thought to have occurred at Kilauea around 1868 (Wyss et al. 1981). Since 1868, repeated episodes of magma intrusion into the rifts have compressed the south flank--in effect, pushing it away from the rest of the island (fig. F.2). The large 1975 rupture is a result of this 100+ years of stress loading.

Ando (1979) published a detailed study of the mechanism of the Kalpana earthquake, including a dislocation fault model. Ando's dislocation parameters, describing the size and orientation of a fault plane and amount and sense of slip on this fault plane, were estimated from a comprehensive set of body wave and surface wave data, the aftershock distribution, tsunami effects, and geodetically derived crustal deformation data. Note that Ando's use of the geodetic data, however, was strictly qualitative—a yardstick to compare the coseismic models suggested by the other data types or to support the hypothesized strain pattern on the south flank of Kilauea. In Ando's model the dislocation surface is thought to represent the boundary between the original ocean floor and the overlying layer of solidified lava.

Table F.1--Dislocation parameters (Ando 1979)

Strike	= N70°E
Dip	= 20° SSE
Fault length	= 40 km
Fault width	= 20-30 km
Upper depth	= 10 km (measured vertically)
Dip slip	= 3.7 - 5.5 m (normal slip)

We used Ando's dislocation parameters (table F.1) as a "first approximation" dislocation plane to model the 1975 earthquake with the geodetic data. The location of this fault plane--the "ramp" in figure F.2--was computed from the seismically determined epicenter of the earthquake. A computer program, called JIGGL, was next used to vary the dislocation model parameters until a best fit to the geodetic data was found.

Given "observed" displacement vectors, JIGGL first computes corresponding horizontal shear strains for a collection of prespecified station groupings (for example, shear

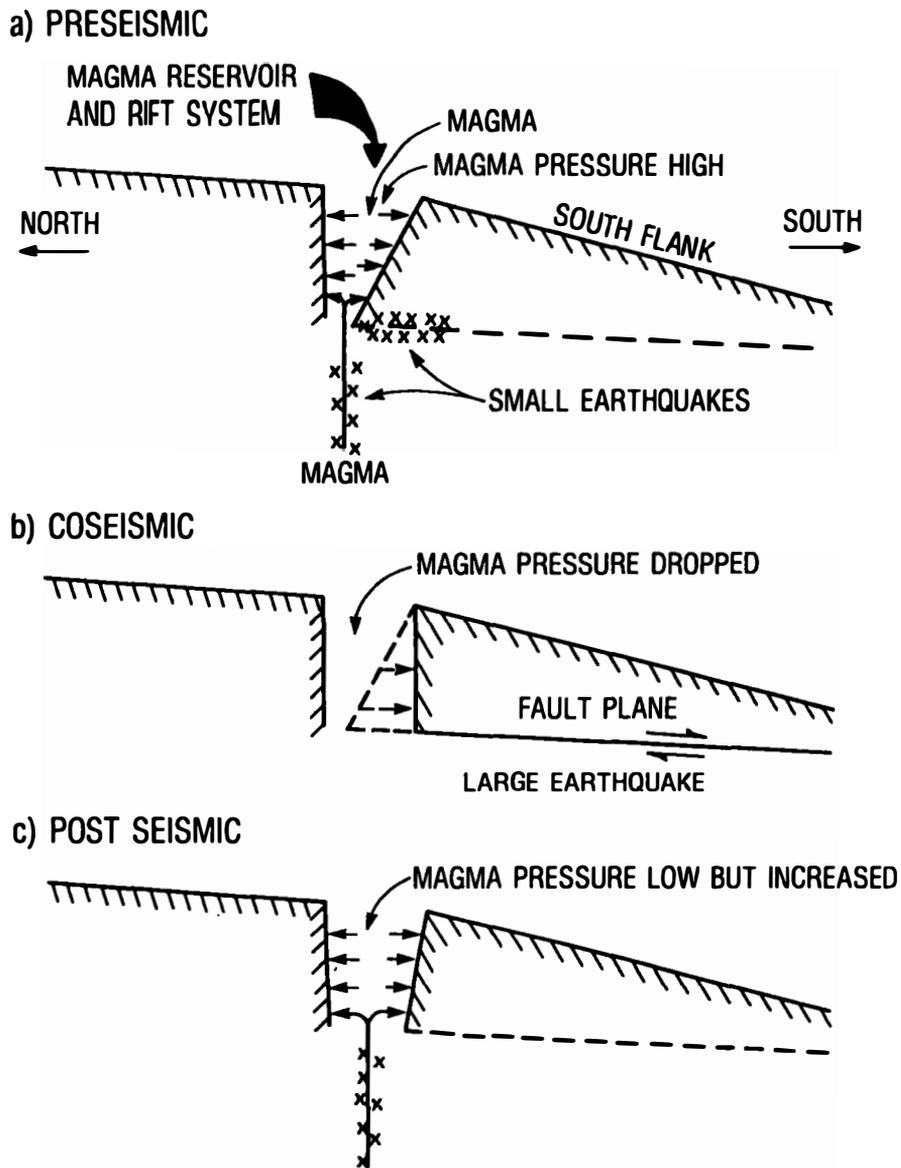


Figure F.2.--A schematic model for strain (a) accumulation, (b) release (earthquake), and (c) reaccumulation in the south flank of Kilauea. The sliding plane corresponds to the 1975 fault plane and is also inferred from background seismicity (from Ando 1979).

strain is computed for selected triangles in the network). Next, the software computes, for these same groupings, theoretical shear strains corresponding to the input dislocation model. The rms difference between the observed and theoretical strains is computed. The software then varies the values of the fault parameters seeking the best rms fit. JIGGL tests each dislocation parameter, one-by-one, to see if changing its current value by plus-or-minus some user-specified increment will decrease the rms. Parameter values are then changed, if necessary, to reflect the best fit. After testing the last parameter, the procedure is iterated until either a convergence criteria is met or a specified number of iterations are performed.

In the case of the Kalpana earthquake the observed displacements were computed at 68 geodetic stations by differencing the adjusted positions resulting from a minimally constrained adjustment of the 1974 USGS EDM data with adjusted positions from such an adjustment of the 1976 USGS EDM data.

The JIGGL results for Ando's dislocation model were poor. The best fitting set of parameters had an rms strain of 70 microradians (about 10σ). Furthermore, this solution required over 18 meters of normal dip slip on a plane dipping only 2° from the horizontal. These results are geophysically implausible. An additional fault plane, representing the vertical magma chamber on Kilauea (see fig. F.2), was added to Ando's model and the JIGGL process was repeated. This new model produced an even poorer fit to the geodetic displacements--an rms strain of over 112 microradians. Moreover, the resultant fault plane parameters were still unreasonable with more than 10 meters of dip slip on the "ramp" plane and the "magma chamber" plane dipping 8° to the northwest. There are several possible explanations for our inability to model the 1975 Kalpana earthquake with a dislocation model:

- (a) anomalous local offsets may be clouding the coseismic displacements;
- (b) extensive aseismic motions occurring over the 18-month period between the preseismic and postseismic geodetic surveys may bias the results;
- (c) the mathematics of dislocation theory may be inappropriate for modeling the displacements associated with this event.

Evidence in support of items (a) and (b) is discussed later in this appendix. Relating to item (c), Eissler and Kanamori (1987) have proposed that the 1975 Kalpana earthquake was due to a massive landslide and not to elastic dislocation. Concern for item (c) caused us to adopt an alternate modeling approach. In this approach, the coefficients of a second-order polynomial in two variables (latitude and longitude) were estimated from the displacements. In other words, the observed geodetic displacements (table F.2) were fit by a quadratic function of position. The same technique had been introduced by Steven Musman to model the 1964 Prince William Sound, Alaska, earthquake (appendix E). With the exception of the stations noted below the polynomial model fit the Hawaiian data rather well. Indeed for the 21 sites shown in figure F.3, the discrepancy between observed and modeled displacements has an rms value of only 0.33 m whereas observed displacements range between 0.74 m and 6.43 m.

Six stations on the summit of Kilauea volcano experienced anomalous coseismic motion. These stations were treated as distinct points--"before" and "after" the 1975 earthquake. These stations are:

HVO 113 USGS
KEANAKAKOI 1949
OHAIKEA HTS 1949
SAND HILL USGS
UWEKAHUNA HGS 1896
VOLCANO HOUSE FLAG

Similarly, the stations listed below were also treated as before-and-after stations. Station descriptions and written correspondence between the USGS and NGS suggest that mark stability is questionable on the entire south flank of Kilauea. This isn't surprising since the marks are set in young, solidified lava. Many points have been reset owing either to the geological instability, underlying magma intrusion, or nearby lava flows.

Table F.2.--Coseismic displacements for 1975 Kalpana earthquake

Station	Latitude		Longitude		Displacement	
	deg	min	deg	min	East, m	North, m
1 APUA PT BM	19	15.8	155	11.7	3.42	-5.45
2 GOAT 2 USGS	19	20.0	155	13.7	1.41	-2.58
3 HAKUMA 1914	19	20.9	154	58.9	1.38	-1.91
4 HEIHEIAHULU 1914	19	25.6	154	59.6	0.23	-1.22
5 HILINA USGS	19	17.9	155	18.6	1.64	-3.42
6 HOLEI HVO162 USGS	19	19.0	155	08.2	1.42	-3.03
7 HVO 147 USGS	19	24.1	154	55.5	0.71	-0.54
8 IILEWA USGS	19	26.3	154	57.6	0.01	-0.74
9 KAENA POINT USGS 1977	19	17.1	155	07.5	2.73	-3.97
10 KALALUA USGS	19	24.2	155	04.1	0.49	-1.50
11 KAMDAMOA HVO158 USGS	19	19.3	155	03.9	1.65	-2.56
12 KAU USGS	19	17.3	155	19.6	1.57	-3.55
13 KUPAPAU 1914	19	19.9	155	01.2	1.62	-2.48
14 LAEAPUKI 1914	19	18.2	155	05.4	2.04	-3.10
15 MAKAHANU PALI USGS	19	18.2	155	15.5	1.75	-4.40
16 MOANA HAUAE USGS 1978	19	22.5	154	57.2	1.07	-1.12
17 OHALE HGS 1897	19	21.2	155	16.8	1.10	-2.17
18 PANAU 1914	19	19.4	155	06.5	1.46	-2.55
19 PILAU 3 USGS	19	19.3	155	11.4	1.53	-2.41
20 PULAMA 1914	19	21.3	155	02.6	1.15	-1.89
21 PUU HULUHULU HGS 1891	19	22.5	155	12.5	0.90	-1.71

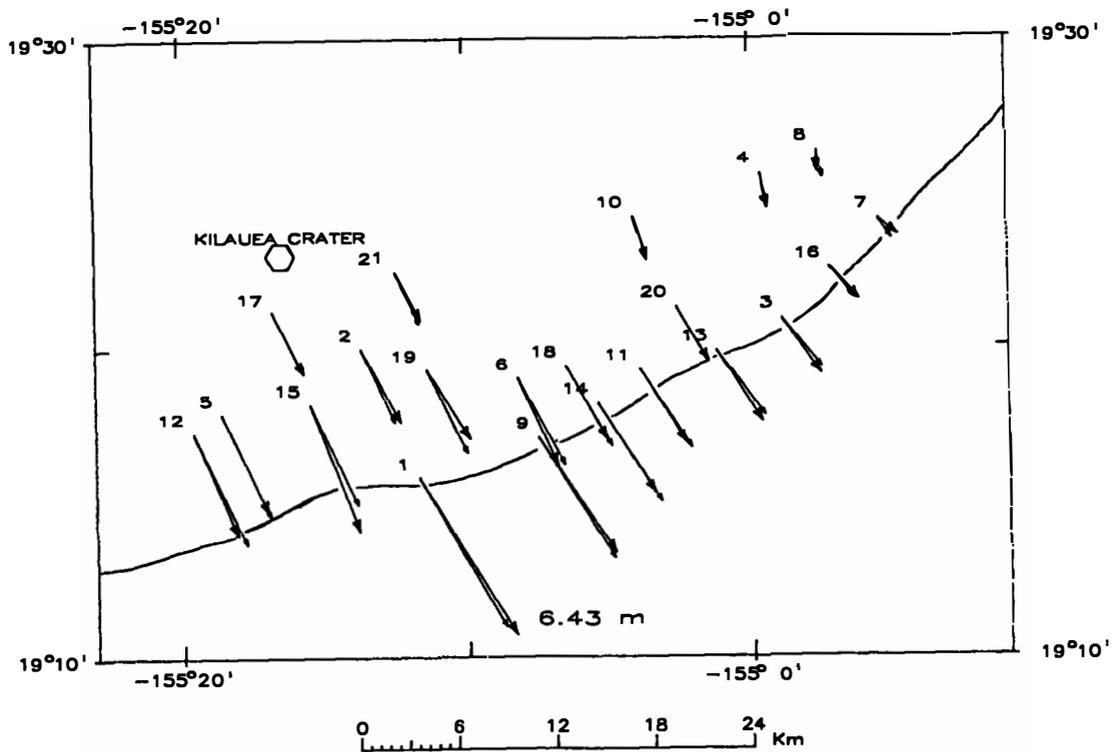


Figure F.3.--Comparison between observed displacements (large arrowheads) and modeled displacements (smaller arrowheads) for the 1975 Kalpana earthquake. Numerals identify stations as keyed in table F.2.

HVO 128 USGS
KAAHA USGS
KALIU HGS 1891
KOAE USGS 1949
NALI USGS
PAPAYA RESET USGS
PUU KAPUKAPU 1914

The MOTION III software (chapter 4) was used to derive a model for the secular movement of the stations in the Hilo region. While the formulation of the model allows for both secular and episodic movement, only the secular parameters were estimated in the Hawaii application of MOTION III. To achieve this, "front end" software was written to correct the geodetic observations for coseismic motion attributed to the 1975 earthquake. Using the results from the empirical coseismic model, corrections were applied to all observations involving stations in the "earthquake zone" (described later in this appendix). The pre-earthquake observations were corrected to be compatible with the post-earthquake observations. This allowed us to model the secular motion as if the earthquake had not occurred.

The Hilo region was divided into five districts (fig. F.1). This districting scheme differs from those used for most other REDEAM models in that district boundaries do not correspond to geologic faults. The rationale for the Hawaii districting scheme follows:

- District I -- A buffer between the earthquake zone (District IV) and the rest of the island (District V).
- District II -- Station OHALE HGS 1897 (and a no-check spur) only. The reasons for this district are given below.
- District III -- Station PUU HULUHULU HGS 1891 (and a no-check spur) only. The reasons for this district are given below.
- District IV -- The earthquake zone. District II and III are embedded in, but separate from, this district.
- District V -- The rest of the "Big Island."

Note, that there are several fault systems on the flanks of Kilauea. There is, however, insufficient temporal and spatial distribution of geodetic data in these areas to justify the finer districting scheme required to resolve movements associated with these fault systems.

Swanson et al. (1976) published a study of the 20th century triangulation, trilateration, and leveling data in the vicinity of Kilauea. These data (and the preparation of their report) precede the 1975 earthquake. Swanson et al. examined the geodetic data epoch-by-epoch and related observed displacements to the geology and to episodes of forceful magma intrusion. Their report is a landmark study on south flank tectonics and offers a concise picture of secular deformation patterns. In summarizing the geodetic data from 1914 to 1971, Swanson et al. showed that the summit area undergoes uplift between eruptions. In addition, the filling and emptying of the magma reservoir system beneath the summit of Kilauea are generally accompanied by largely reversible ground deformation. Those authors document short term deformation events keyed to specific eruptions or episodes of ground cracking associated with intrusions. A good example of extensive aseismic movement on the south flank is seen by comparing the 1970 and 1974 USGS trilateration data. These data span a time period characterized

by the onset of swarms of intrusions and eruptions which culminated in the 1975 earthquake. The picture, then, is that of strain accumulation on the south flank that is nonlinear in time and spatially heterogeneous. Obviously a more sophisticated model than ours (which has the assumption of linear strain accumulation over time and spatial homogeneity within districts) would be required to describe such complex deformation. It is not clear that a single model--rather than several models keyed to specific episodes of eruptions or intrusions--is possible.

Swanson et al. also noted large, atypical displacements at stations CHALE and PUU HULUHULU for the 1949-70 data epoch. In our early adjustments, these two stations had large residuals and we therefore designated them as districts II and III in our final districting scheme. These two "ad hoc" districts are embedded in district IV, which encompasses the earthquake zone.

The following FORTRAN routine documents the derived polynomial coefficients that characterize the coseismic displacements associated with the 1975 Kalpana earthquake. Note that nonzero displacements are obtained only for districts II, III, and IV.

```

SUBROUTINE HAMOV(XLA,XLO, IDIST,DX,DY)
C
C This subroutine interpolates to find the displacement at
C an arbitrary position for the 1975 Hawaii earthquake.
C INPUT:
C   XLA denotes the (north) latitude of the position in degrees.
C   XLO denotes the (west) longitude of the position in degrees.
C   IDIST is a district identifier.
C OUTPUT:
C   DX is the eastward displacement in meters at the position.
C   DY is the northward displacement in meters at the position.
C
IMPLICIT DOUBLE PRECISION (A-H,O-Z)
DIMENSION CX(5), CY(5)
PI = 3.14159265D0
IF(IDIST.EQ.1 .OR. IDIST.EQ.5) THEN
  DX = 0.0D0
  DY = 0.0D0
ELSEIF(IDIST.EQ.2 .OR. IDIST.EQ.3 .OR. IDIST.EQ.4) THEN
  XBAR = 7.2154D0
  YBAR = 5.4478D0
  XX = 45.9756D0
  XY = 11.6674D0
  YY = 7.5701D0
  DX0 = 1.397D0
  DY0 = -2.5D0
  CX(1) = .03058D0
  CY(1) = .03280D0
  CX(2) = -.3087D0
  CY(2) = .3881D0
  CX(3) = .003245D0
  CY(3) = .003092D0
  CX(4) = -.02123D0
  CY(4) = .002082D0
  CX(5) = .04132D0
  CY(5) = -.04603D0
  X = DCOS(19.25D0*PI/180.D0)*60.D0*(155.25D0 -XLO)
  X = X - XBAR

```

```

      Y = 60.D0*(XLA - 19.25D0)
      Y = Y - YBAR
      F = X*X - XX
      G = X*Y - XY
      H = Y*Y - YY
      DX = DX0 + CX(1)*X + CX(2)*Y + CX(3)*F + CX(4)*G + CX(5)*H
      DY = DY0 + CY(1)*X + CY(2)*Y + CY(3)*F + CY(4)*G + CY(5)*H
ELSE
      WRITE(6,10)
10    FORMAT(' IMPROPER DISTRICT IDENTIFIER')
ENDIF
RETURN
END

```

ACKNOWLEDGMENTS

We, the authors, express our gratitude to the Federal, State, local, and private organizations that provided data for the development of the REDEAM models. We especially thank those individuals who participated directly in model development: Steven A. Musman (NGS), Kenneth J. Hurst (Columbia Univ./Lamont-Doherty), and Christiana Mitsakaki (National Technical Univ. of Athens, Greece). Anna-Mary B. Miller (NGS) authored the software that applied the developed models to the redefinition of the North American horizontal reference system. Robert H. Hanson (NGS) designed graphics software that has enhanced both interpretation and presentation of the models. We also thank several other programmers who developed helpful software, supervisors who encouraged our efforts, scientists who shared their ideas, and numerous NGS employees who assisted with data processing and with publication preparation; in particular:

Eleanor Z. Andree	(NGS)	David C. McAdoo	(NGS)
Peter Bird	(UCLA)	Dennis G. Milbert	(NGS)
John D. Bossler	(NGS/Ohio State Univ.)	F. Foster Morrison	(NGS)
Bernard Chovitz	(NGS/retired)	Allen J. Pope	(NGS/deceased)
Cindra S. Craig	(NGS)	William J. Prescott	(USGS)
Jimmie L. David	(NGS)	W. I. Reilly	(New Zealand)
Sue Dietterle	(NGS)	David T. Sandwell	(NGS/Univ. of Texas)
William H. Dillinger	(NGS)	James C. Savage	(USGS)
Bruce C. Douglas	(NGS)	Tomas Soler	(NGS)
Nancy G. Doyle	(NGS)	Rodney K. Spinks	(NGS)
Stephen J. Frakes	(NGS)	Ross S. Stein	(USGS)
Rudolf J. Fury	(NGS)	William E. Strange	(NGS)
John G. Gergen	(NGS)	Wayne Thatcher	(USGS)
Soren W. Henriksen	(NGS/retired)	George A. Thompson	(Stanford Univ.)
Edward Herbrechtsleier	(NGS)	Carl A. Wagner	(NGS)
Sandford R. Holdahl	(NGS)	Robert E. Wallace	(NGS)
Vasanthi Kammula	(NGS)	Karen E. Weathers	(NGS)
William M. Kaula	(NGS/UCLA)	James E. Whitcomb	(U. Colorado)
James R. Lucas	(NGS)	Charles A. Whitten	(NOAA/retired)

REFERENCES

- Aki, K., 1984: Asperities, barriers, characteristic earthquakes and strong motion prediction. J. Geophys. Res., 89, 5867-5872.
- Ando, M., 1979: The Hawaii earthquake of November 29, 1975: low dip angle fault due to forceful injection of magma, J. Geophys. Res., 84, 7616-7626.
- Bennett, J. H., 1980: Geodimeter measurements of strain and slip along the northern San Andreas fault system, Spec. Rep. Calif. Div. Mines Geol., 140, 99-109.
- Bonilla, M. G., and Buchanan, J. M., 1970: Interim report on worldwide historic surface faulting. U.S. Geological Survey, Open-File Report, 32 pp.
- California Department of Water Resources, 1968: Geodimeter fault movement investigations in California. Calif. Dept. Water Resources Bull., 116 (6), 1-183.
- Chinnery, M. A., 1961: The deformation of the ground around surface faults, Bull. Seismol. Soc. Am., 51, 355-372.
- Cline, M. W., Snay, R.A., and Timmerman, E. L., 1985: Geodetically derived strain from San Francisco Bay to the Mendocino triple junction, California. NOAA Tech. Rep., NOS 109 NGS 31, National Geodetic Information Center, NOAA, Rockville, MD, 17 pp.
- Cline, M. W., Snay, R. A., and Timmerman, E. L., 1984: Regional deformation of the earth model for the Los Angeles region, California. Tectonophysics, 107, 279-314.
- Dunbar, W. S., 1977: The determination of fault models from geodetic data. Ph.D dissertation, Stanford Univ., Stanford, California, 221 pp.
- Eissler, H. K., and Kanamori, H., 1987: A single-force model for the 1975 Kalpana, Hawaii, earthquake, J. Geophys. Res., 92, 4827-4836.
- Gergen, J. G., 1975: The new adjustment of the North American Datum: the observables, ACSM Bulletin, (Nov. issue), 9.
- Hastie, L. M., and Savage, J. C., 1970: A dislocation model for the 1964 Alaska earthquake. Bull. Seismol. Soc. Am., 60, 1389-1392.
- Hayford, J. F., and Baldwin, A. L., 1908: Geodetic measurements of earth movements: in, A. C. Lawson (Chairman), The California Earthquake of April 18, 1906--Report of the State Earthquake Commission, vol. 1, pp. 114-145, Carnegie Institution of Washington, Washington, D.C.
- King, N. E., and Savage, J.C., 1983: Strain-rate profile across the Elsinore, San Jacinto, and San Andreas faults near Palm Springs, California, 1973-1981, Geophys. Res. Lett., 10, 55-57.
- Mansinha, L., and Smylie, D. E., 1971: The displacement fields of inclined faults, Bull. Seismol. Soc. Am., 61, 1433-1440.

- Meade, B. K., 1971: Report of the Subcommittee on Recent Crustal Movements in North America: in, Reports on Geodetic Measurements of Crustal Movement, 1906-71, National Geodetic Information Center, NOAA, Rockville, MD.
- Minster, J. B., and Jordan, T. H., 1978: Present-day plate motion. J. Geophys. Res., 83, 5331-5354.
- National Oceanic and Atmospheric Administration, 1971: Triangulation Diagrams - State of California, (special edition with geologic faults), National Geodetic Information Center, NOAA, Rockville, MD.
- Okada, Y., 1985: Surface deformation due to shear and tensile faults in a half-space. Bull. Seismol. Soc. Am., 75, 1135-1154.
- Parkin, E. J., 1969: Horizontal crustal movements determined from surveys after the Alaskan earthquake of 1964, in The Prince William Sound, Alaska Earthquake of 1964 and Aftershocks, Vol. III, L. E. Leipold, Editor, U.S. Government Printing Office, Washington, D.C., 35-98.
- Savage, J. C., 1983: Strain accumulation in the western United States. Ann. Rev. Earth Planet. Sci., 11, 11-43.
- Savage, J. C., and Prescott, W. H., 1973: Precision of geodolite distance measurements for determining fault movement. J. Geophys. Res., 78, 6001-6008.
- Savage, J. C., and Hastie, L. M., 1966: Surface deformation associated with dip-slip faulting. J. Geophys. Res., 71, 4897-4904.
- Schwarz, C. R., 1978: The TRAV-10 horizontal network adjustment program. NOAA Tech. Mem., NOS NGS-12, National Geodetic Information Center, NOAA, Rockville, MD, 52 pp.
- Snay, R. A., Cline, M. W., and Timmerman, E. L., 1986: Horizontal crustal deformation models for California from historical geodetic data. Royal Soc. New Zealand Bull., 24, 131-140.
- Snay, R. A., Cline, M. W., and Timmerman, E. L., 1985: Dislocation models for the 1954 earthquake sequence in Nevada: in, Proc. of Workshop XXVIII: the Borah Peak, Idaho, Earthquake, U.S. Geol. Surv. Open-File Report 85-290, 531-555.
- Snay, R. A., Cline, M. W., and Timmerman, E. L., 1983: Regional deformation of the earth model for the San Diego region, California. J. Geophys. Res., 88, 5009-5024.
- Snay, R. A., Cline, M. W., and Timmerman, E. L., 1982: Horizontal deformation in the Imperial Valley, California, between 1934 and 1980. J. Geophys. Res., 87, 3959-3968.
- Snay, R. A., 1976: Reducing the profile of sparse symmetric matrices. Bull. Geodesique, 50, 341-352.

- Stein, R. S., and Thatcher, W., 1981: Seismic and aseismic deformation associated with the 1952 Kern County, California, earthquake and relationship to the quaternary history of the White Wolf fault, J. Geophys. Res., 86, 4913-4928.
- Steketee, J. A., 1958a: On Volterra's dislocations in a semi-infinite elastic medium, Can. J. Phys., 36, 192-205.
- Steketee, J. A., 1958b: Some geophysical applications of the elasticity theory of dislocations, Can. J. Phys., 36, 1168-1198.
- Swanson, D. A., Duffield, W. A., and Fiske, R. S., 1976: Displacement of the south flank of Kilauea volcano: the result of forceful intrusion of magma into the rift zones, U.S. Geol. Surv. Prof. Paper 963, 39 pp.
- Wyss, M., Johnson, A. C., and Klein, F. W., 1981: Multiple asperity model for earthquake prediction, Nature, 289, 231-234.